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Evaluation of Amino Acid Composition as a Geochronometer in Buried Soils on Mount Kenya, East Africa Détermination de la composition de l'acide aminé pour le géochronologie des sols enfouis du mont Kenya, Afrique de l'Est

Bestimmung der Amino-Säure als Geochronometer der begrabenen Böden auf dem Mount Kenya, Ost-Afrika

William C. Mahaney, Michael G. Boyer et Nathaniel W. Rutter

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Résumé de l'article

On a identifié et daté, au radiocarbone et à l'aide de techniques de datations à l'acide aminé, des paléosols enfouis et de surface afin de comprendre l'évolution quaternaire de la région. Les datations au radiocarbone des paléosols enfouis varient de 900 à plus de 40 000 BP. Les sols se sont développés dans des dépôts glaciaires et périglaciaires de différentes textures, constitués de fragments de roches détritiques, de phonolite, de basalte et de syénite. Tous les paléosols, sauf deux, sont situés dans la zone afroalpine (au-dessus de 3200 m). On a déterminé par racémisation les rapports D/L des acides aminés dans les horizons Ab en vue d'évaluer leur fiabilité pour la datation des âges relatifs. On a fait l'analyse de l'alaline, de l'acide aspartique, de l'acide glutamique, de la leucine, de la valine et du phénylalaline. L'acide aspartique, comme dans d'autres cas, a donné des résultats particulièrement satisfaisants, les quotients plus élevés correspondant aux âges les plus anciens. D'autres acides montraient des tendances bien distinctes, mais moins convaincantes que dans le cas de l'acide aspartique. Ainsi, dans la plupart des cas, les relations quotients/âges de l'acide aspartique étaient corroborées par les datations au radiocarbone. Les rapports D/L de l'acide aspartique variaient d'environ 0,07 pour les échantillons modernes à environ 0,45 pour les échantillons de plus de 40 000 ans.

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EVALUATION OF AMINO ACID COMPOSITION AS A GEOCHRONOMETER IN BURIED SOILS ON MOUNT KENYA, EAST AFRICA*

William C. MAHANEY, Michael G. BOYER and Nathaniel W. RUTTER, respectively: Geography Department, Atkinson College, York University, North York, Ontario M3J 1P3; Biology Department, York University, North York, Ontario M3J 1P3; Geology Department, University of Alberta, Edmonton, Alberta T6G 2E3.

ABSTRACT A sequence of surface and buried paleosols from the slopes of Mount Kenya, East Africa, has been identified and dated by radiocarbon and amino acid dating techniques in order to elucidate the Quaternary history of the area. Buried paleosols vary in radiocarbon age from 900 to $> 40,000$ yrs BP. They have developed in glacial and periglacial deposits of variable texture, consisting of a high percentage of clasts of phonolite, basalt and syenite. All but two paleosols are located in the Afroalpine zone (above 3200 m). D/L ratios of amino acids in Ab horizons were determined in order to establish their reliability for relative age dating. Alanine, aspartic acid, glutamic acid, leucine, valine, and phenylalanine were routinely analyzed. Aspartic acid, as in other cases, proved reliable yielding remarkably consistent results, with higher ratios corresponding to increasing age. Other acids analyzed showed distinct trends, although not as convincing as aspartic acid. In most cases, the aspartic acid ratio/ age relationships were supported by radiocarbon dates. D/L ratios of aspartic acid varied from approximately 0.07 for modern samples, to approximately 0.45 in samples $>$ 40,000 years old.

RÉSUMÉ Détermination de la composition de l'acide aminé pour le géochronologie des sols enfouis du mont Kenya, Afrique de l'Est. On a identifié et daté, au radiocarbone et à l'aide de techniques de datations à l'acide aminé, des paléosols enfouis et de surface afin de comprendre l'évolution quaternaire de la région. Les datations au radiocarbone des paléosols enfouis varient de 900 à plus de 40 000 BP. Les sols se sont développés dans des dépôts glaciaires et périglaciaires de différentes textures, constitués de fragments de roches détritiques, de phonolite, de basalte et de syénite. Tous les paléosols, sauf deux, sont situés dans la zone afroalpine (au-dessus de 3200 m). On a déterminé par racémisation les rapports D/L des acides aminés dans les horizons Ab en vue d'évaluer leur fiabilité pour la datation des âges relatifs. On a fait l'analyse de l'alaline, de l'acide aspartique, de l'acide glutamique, de la leucine, de la valine et du phénylalaline. L'acide aspartique, comme dans d'autres cas, a donné des résultats particulièrement satisfaisants, les quotients plus élevés correspondant aux âges les plus anciens. D'autres acides montraient des tendances bien distinctes, mais moins convaincantes que dans le cas de l'acide aspartique. Ainsi, dans la plupart des cas, les relations quotients/âges de l'acide aspartique étaient corroborées par les datations au radiocarbone. Les rapports D/L de l'acide aspartique variaient d'environ 0,07 pour les échantillons modernes à environ 0,45 pour les échantillons de plus de 40 000 ans.

ZUSAMMENFASSUNG Bestimmung der Amino-Sàure als Geochronometer der begrabenen Bôden auf dem Mount Kenya, Ost-Afrika. Eine Serie von an der Oberfläche liegenden und begrabenen Palàobôden von den Hàngen des Mount Kenya, Ost-Afrika, wurde mittels Radiokarbon- und Aminosâuredatierungs-techniken identifiziert und datiert, um die Geschichte dieses Gebiets im Quaternàr zu erhellen. Das durch Radiokarbon bestimmte Alter der begrabenen Palàobôden variiert von 900 bis $>$ 40,000 Jahren v.u.Z. Diese Bôden haben sich in glazialen und periglazialen Ablagerungen verschiedener Beschaffenheit entwickelt, welche zu einem hohen Prozentsatz aus Trùmmern von Phonolith, Basalt und Syenit bestehen. Außer zweien befinden sich aile Palàobôden in der afroalpinen Zone (oberhalb 3200 m). Die D/L Anteile der Aminosâuren in den Ab-Horizonten wurden bestimmt, um ihre Verlâpiichkeit bei der relativen Altersbestimmung festzustellen. Alamin, aspartische Sàure, Glutamin-Sàure, Leuzin, Valin und Phenylalanin wurden laufend analysiert. Wie in anderen Fällen erwies sich die aspartische Säure als verläßlich, indem sie bemerkenswert bestàndige Ergebnisse ergab, bei denen die hôheren Quotienten dem hôheren Alter entsprachen. Andere analysierte Sàuren zeigten ausgepràgte Trends, wenn auch nicht so ùberzeugend wie die aspartische Sàure. In den meisten Fallen wurden die Beziehungen Quotient/Alter der aspartischen Sàure durch Radiokarbondatierungen gestùtzt. Die D/L Anteile der aspartischen Sàure variierten von ungefähr 0.07 für moderne Proben bis ungefâhr 0.45 in Proben, die > 40,000 Jahre alt sind.

Contribution du premier symposium de la CANQUA, sous la direction de René W. Barendregt

INTRODUCTION

The fate of amino acids in buried paleosols is of considerable interest to Quaternary geologists, particularly since the racemization of aspartic acid and others may provide a means of dating organic-rich horizons. Dated paleosols ranging from 960 to > 42,000 yrs BP have been described on Mount Kenya (Fig. 1a and 1b; 0°10'S; 37°15'E) in several drainage basins between 2990 and 4050 m elevation (MAHANEY, 1979, 1981, 1982a, 1982b, 1983, 1984a, 1984b and 1985). Most soils have developed in transported regolith with different lithologie compositions. We consider this paleosol sequence to provide

a convenient field setting in which to test the reliability of using amino acid racemization for age determination, and where the deposits can be calibrated by radiocarbon. Moreover, time stratigraphy refinement of amino acid racemization assists in increasing our understanding of the timing of geologic events before, during, and following the last (Liki) glaciation on Mount Kenya (Fig, 2).

Many varieties of fossils are now routinely analyzed to determine the D/L ratios of amino acids for relative and absolute age dating of sediments. The method has been particularly useful when fossils have experienced similar temperature

FIGURE 1. Paleosols sites on the western flank (a) and eastern flank (b) of Mount Kenya, East Africa.

Sites des paléosols sur les versants ouest (a) et est (b) du mont Kenya, Afrique de l'Est.

GLACIAL
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advance $pos1 - L1$ $10(2)$ \sim 501 \overline{C} Darwin 50115 ADJACENT NUMBERS) Formation $\overline{1}$ $\overline{0}$ $\frac{1}{\sigma}$ e lek i 12.5 Ш \overline{c} $\begin{array}{c} \n\zeta_{\mathcal{F}_{\mathcal{F}_{\mathcal{F}}}}\n\end{array}$ \overline{a} $\mathop{\rm III}\nolimits$ Stages II-V \vdash \overline{a} **Proccorps** λ Z \circ $\ddot{}$ \mathfrak{a} $_{\rm II}$ ω \bar{z} Stage IB-D ω \circ α BETWEEN Ш $\mathbf I$ Stage IA $DOS1$ $\frac{k}{2}$ Teleki/Liki $100(2)$ $\frac{1}{\omega}$ \circ Interglaciation \vdash LINEAR \mathbf{I} ω \circ Teleki Older Glaciation Ω Glaciation $\overline{0}$ \vdash BP(SCALE Teleki/ pre-Telek ω Interglaciation YEARS pre-Teleki Ш \int_{0}^{3} Glaciation $500(2)$ Ω

FIGURE 2. Correlation diagram for Pleistocene and Holocene deposits on Mount Kenya. Unit boundaries are based on maximum and minimum dates for glacial and nonglacial sediments. Dashes denote boundaries for which there is no radiometric control. Arrows indicate directions in which boundaries are likely to be shifted when additional dates become available.

Diagramme de corrélation des dépôts pléistocènes et holocènes sur le mont Kenya. Les limites des unités sont fondées sur les dates maximales et minimales de sédiments glaciaires et non glaciaires. Les tiretés identifient les limites qui ne sont pas datées selon des méthodes radiométriques. Les flèches pointent vers les directions que les limites pourront prendre lorsque de nouvelles datations s'ajouteront.

histories and diagenetic conditions. Marine molluscs have been widely utilized and are probably the most successful for age determination, whereas freshwater and terrestrial molluscs offer promise (ANDREWS et al., 1979; BELKNAP and WEHMILLER, 1980; BRIGHAM, 1982,1983; EMERSON et al., 1981; McCOY, 1980; MILLER, 1982; MILLER and HARE, 1974; MILLER ef a/., 1977, 1983; MITTERER, 1974, 1975; RUTTER ef a/., 1980; WEHMILLER, 1982; WEHMILLER et al., 1977). Bone and wood have been useful, although problems have been encountered (BADA, 1972; BADA and HELFMAN, 1975; BADA et al., 1973, 1974, 1979; BENDER, 1974; HARE, 1974; HO, 1965, 1967; LEE et al., 1976; RUTTER and CRAWFORD, 1984). In addition, investigations involving amino acid racemization age dating have been carried out with marine sediments (BADA et al., 1970; KVENOLDEN et al., 1970), foraminifera (KING, 1980; WEHMILLER and HARE, 1971) coral (WEHMILLER et al., 1976, WEHMILLER and HARE, 1970) and with soils (LIMMER and WILSON, 1980).

Racemization (or epimerization) of amino acids in geological specimens is a complex process involving a number of factors. The rates of interconversion processes are dependent upon the species of amino acid, ambient temperatures, types of materials in which the amino acids are embedded (shells, bones, wood, etc.), and the state in which the amino acids occur, i.e. whether they are terminally bound, internally bound, or free amino acids. In addition, other external diagenetic factors such as moisture content, acidity, and oxidation-reduction conditions can play a role. The accuracy of D/L determinations rests with the care of preparation and the reliability of the analytical methods employed. These factors have contributed to the apparent lack of interest in amino acid D/L ratio determinations of paleosols. The complex nature of soil formation and diagenetic changes that take place at the surface and upon burial compound the problems already inherent in the method. It will be seen, however, that in the present work there is a trend for D/L ratios, of aspartic acid, to increase with age of the paleosol.

FIELD AREA

A number of major deposits of Pleistocene and Holocene age are found in the Afroalpine area of Mount Kenya (Fig. 1 a and 1b for location; Fig. 2 for stratigraphy). These glacial and nonglacial deposits have been assigned ages on the basis of topographic position (BAKER, 1967), as well as weathering features, degree of soil expression, and radiocarbon (MA-HANEY, 1979, 1980, 1981, 1982a, 1982b).

CLIMATE

The Afroalpine area rises from \sim 3200 m to the summit of Mount Kenya at 5199 m (Fig. 1a). Precipitation derived from the Indian Ocean is delivered by the Tropical Easterly Wind Belt as well as the classical and Atlantic monsoon system. The southeastern slopes are wettest with rainfall reaching 2500 mm (at 2250 m); the northern slopes are driest with \sim 1000 mm annually. On the western flank of the mountain, precipitation rises from \sim 500 mm at Naro Moru (2000 m), to \sim 1000 mm on the high moorland (\sim 3500 m), and then drops in the central peaks areas (\sim 4300 m) to 750-800 mm (COETZEE, 1967). Temperature extrapolations (COE, 1967; THOMPSON, 1966) suggest that temperatures at 4200 m average 3.1 to 3.6°C (from observations taken during two periods in 1948; COETZEE, 1967; HEDBERG, 1964). The depth of frost penetration in Afroalpine sites while limited to a few centimetres, is diurnal in nature, and controlled largely by variations in thermal conductivity of different cover sediments. Variations in subsurface (-1.0 m depth) temperatures, $\frac{1}{2}$ metric. Vanations in subsurface $($ \cdots of a copyrigremperatures, as they might allect racemization rates or all no actus are

warmer Holocene climate. Frost penetration in surface deposits in the Bamboo Forest, while nil at present, may have occurred to depths of a few centimetres during the last glaciation when ice limits reached \sim 3100 m. As in the Afroalpine zone, however, paleoclimatic variations in surface temperature did not affect buried paleosols insofar as racemization rates are concerned.

VEGETATION

Below timberline the subalpine forest consists of Hagenia Woodland dominated by Hagenia abyssinica and Bamboo Forest (Arundinaria alpina). Above the heather zone at timberline (\sim 3200 m) the lower Afroalpine zone, dominated by Senecio brassica, extends upwards to \sim 3700 m where it merges with the upper Afroalpine zone characterized by Senecio keniodendron. The Upper Afroalpine zone merges with vegetation of the nival zone at \sim 4500 m (COE, 1967; HED-BERG, 1964). The probability that different species produce variable ratios of amino acids was considered to affect the amount of each L-form amino acid that might be present in a paleosol at time zero, although not the D/L ratios which should be close to zero.

METHODS

Selected horizons were sampled in sections in a variety of glacial and nonglacial deposits in several topographic situations ranging from well- to poorly-drained. Samples were collected from each paleosol horizon described in the field following the nomenclature of the SOIL SURVEY STAFF (1951, 1975; and BIRKELAND, 1984). Particle size analyses were performed on the < 2 mm fractions; coarse grade sizes $(> 63 \mu m)$ were calculated by sedimentation following procedures established by BOUYOUCOS (1962) and DAY (1965).

Amino acids were determined at the University of Alberta and at York University. Once in the laboratory, samples were prepared for D/L ratio determinations using the following procedures. Large samples of between 1 and 2.5 grams of organic material were weighed directly into clean vials. It was not possible to clean the samples by the usual method of rinsing and sonicating in HCI and double distilled water because whole-soil samples were utilized.

The samples were hydrolyzed for 24 hours at 108 $^{\circ}$ C in 5.5M HCI. The hydrolyzate was decanted, then dried under vacuum in a rotary biodryer. The dried residue was then taken up in double distilled water and placed on a cation exchange resin. Samples were eluted with 3M NH₄OH and again dried under vacuum. Esterification was accomplished by adding acidified isopropanol and heating for 15 minutes at 100° C. After drying the samples were acylated with pentafluoropropionic anhydride in methylene chloride. The samples were then heated for 5 minutes at 100° C and evaporated to remove any excess PFPA. After being washed with CH_2Cl_2 the samples were dried over potassium carbonate. Later samples were filtered and the filtrate injected into a 25 metre x 0.3 mm Chriasil-Val capillary column installed in a Hewlett-Packard 5840A Gas Chromatograph equipped with an FID detector. A chromatograph and report of a typical Mount Kenya soil are shown in Figures 3a and 3b respectively.

RESULTS AND DISCUSSION

PALEOSOLS

The samples were collected from dated paleosols (Appendices l-XI for descriptions; Fig. 2 for stratigraphy) in the Mount Kenya Afroalpine and Bamboo Forest zones. Radiocarbon dates for the buried soil horizons are from wholesoil samples collected in the center of each horizon and are given in Table I; particle size determinations are in Table II. The youngest buried soil (site TV42) was collected from a prominent debris flow in upper Teleki Valley (Fig. 1a; Appendix I). The buried soil formed in pebbly and cobbly material which was subsequently buried by debris from renewed mass wasting activity that probably coincided with the end of the Tyndall advance (MAHANEY 1984a, 1985).

TV33-Ab (lower)

FIGURE 3. a) Amino acid spectrum for sample TV33 — Ab (lower) (UA1474), and b) chromatograph report of selected amino acids.

a) Spectre des acides aminées de l'échantillon TV33—Ab (inférieur) (UAI474) et b) chromatographic de quelques acides aminés.

In order of increasing age the next oldest soil was described in a valley train deposit in upper Teleki Valley (site TV4a; Fig. 1a; Appendix II). This buried paleosol of clayey silty composition was covered by pebbly sandy alluvium approximately 1940 \pm 100 yr BP (Gak-8273; Table I) radiocarbon years ago; the Ab horizon in the buried profile has been dated by thermoluminescence at 1300 \pm 300 yr BP (MAHANEY, 1985; pers. comm., D. Huntley, 1985).

Alluvial fans representing the Darwin Formation (MAHANEY, 1984a, 1985) formed between 4000 and 7000 radiocarbon years ago. Profile TV33 (Fig. 1a; Appendix III) consists of an lnceptisol (ground soil) over three buried Entisols that radiocarbon date between 3400 ± 170 yr BP (Gak-11322) to 3610 \pm 180 yr BP (Gak-11320) to (Table I). The lowermost Ab horizon in this sequence radiocarbon dates at 3400 \pm 170 yr BP (Gak-11322) and is known to be contaminated by microorganisms (BOYER and MAHANEY, 1985). Profile TV39 (Fig. 1a for location; Appendix IV for morphology) contains an lnceptisol (ground soil) over a buried soil in slopewash. The Ab horizon was covered by slopewash 4540 ± 130 yr BP (Gak-11241; Table I).

On the eastern slopes profile GOR8 (Fig. 1b; Appendix V) in Liki I Till was covered with alluvium at 5110 ± 120 yr BP (Gak-11570; Table I). Profile GOR52 in Gorges Valley (Fig. 1b; Appendix Vl) contains alluvium deposited sometime in the late Pleistocene or early Holocene. An Entisol forming in this sediment was covered by younger alluvium 5100 ± 100 yr BP (Beta-10821; Table I).

The LE3 profile (Fig. 1b; Appendix VII) is a compound paleosol with the buried component formed in loess over till. The Ab horizon radiocarbon dates at $> 43,210$ yr BP (Beta-9085; Table I). This buried paleosol overlies a truncated unweathered deposit consisting of lake sediment which probably predates the Teleki Glaciation (Fig. 2). The surface paleosol is formed in outwash and loess of late Teleki age. Both the ground and buried paleosols are lnceptisols, which are common above timberline on Mount Kenya. Coarse textures prevail in the ground paleosol suggesting only slight weathering; however, in the buried soil a silty loam texture (loess) grades downward into till with sandy loam textures. The cemented lower paleosolum and subsoil in the buried soil are common in older pre-Pleniglacial paleosols on Mount Kenya. The coarsetextured surface paleosol and finer-textured and cemented buried paleosol probably result from wetter conditions following emplacement of Teleki Till. Since the paleosol formed in buried till and loess is weathered to a depth of 100 cm, at least several thousand years would be required to produce the buried paleosol. The surface paleosol must have formed under a dry climate throughout the late Pleistocene and Holocene.

Soil profile TV23 (Fig. 1a; Appendix VIII) consists of a compound buried paleosol overlain with a ground paleosol with somewhat similar characteristics (BOYER and MAHANEY, 1985). Both the buried and surface paleosols are Alfisols, which are more common below timberline on Mount Kenya. Heavy textures in the B2t horizons (Appendix VIII) are the result of downward movement of clay in profiles, which directly

TABLE I

DIL ratios for amino acids in buried soils and radiocarbon dates on Mount Kenya, East Africa

a. Sites are located on Figures 1a and 1b.

b. Not available

- - Nil

influences the development of blocky structures. Reddish colors in this profile result partly from the chemical composition of the parent material which is ultrabasic in nature and high in iron (MAHANEY 1982a, 1984b). The surface paleosol in Teleki Till, and the buried paleosol formed in older till of pre-Teleki age, are representative of paleosols formed in older tills of pre-Teleki age found elsewhere along the western flank of Mount Kenya. On the northern and eastern flanks of the mountain, tills of similar age occur in the same stratigraphie position, but the prevailing vegetation is Alpine Moorland, not Bamboo Forest. Despite a dark color (Appendix VII) in the INAb horizon, attempts at radiocarbon dating provided only a infinite age of $>$ 40,000 yr BP (Gak-11319). Dating by thermoluminescence places this horizon at $64 \pm 24 \times 10^3$ yrs ago (personal communication, G. Berger, January 1985).

Site LE4 (Fig. 1b; Appendix IX), located on a moraine ridge at 3520 m, contains pre-Teleki Till and outwash over bedrock. The ground paleosol forming in loess and outwash to a depth of 54 cm is an lnceptisol. Between 54 and 126 cm depth the buried paleosol represents weathered pre-Teleki Till, the Ab horizon of which gives an infinite radiocarbon age of $>$ 42,629 yr BP (Beta-9086) (Table I). The data (Appendix IX) suggest that following emplacement of pre-Teleki Till a paleosol formed in cobbly material upon which 18 cm of loess was deposited. Because this paleosol contains a finer texture throughout it must have been exposed at the surface for a considerable time following till emplacement and prior to the deposition of the outwash body.

Profile GOR48 (Fig. 1b; Appendix X), located in pre-Teleki Till in Gorges Valley does not contain any buried paleosols. However, cemented subsoil horizons together with its topographic position provide evidence for considerable age. BAKER (1967) considered this surface as representing the earliest glaciation on the mountain.

Site GOR41 (Fig. 1b; Appendix Xl) is situated on the highest strath terrace in the Nithi River Basin along the eastern slopes of Mount Kenya. The lowermost paleosol in the sequence was formed from bedrock emplaced \sim 2.0 Ma. If was later buried by stream activity, presumably during the pre-Teleki Glaciation ($> 0.7 \times 10^6$ yrs).

AMINO ACID RACEMIZATION

As mentioned earlier, soils are probably one of the least desirable materials to analyze for D/L ratios of amino acids because they are subject to so many biogeochemical influences. Indeed, some aspects of contamination are reported by LIMMER and WILSON (1980) in allophanous soils of New Zealand. Besides the complex nature of their formation, and later diagenetic changes that take place at the surface and upon burial, organic matter is difficult to isolate for analysis using gas chromatography. This is clearly illustrated for profile TV33-Ab (lower) (Fig. 3a and 3b) where some peaks of anomalous material are present. To determine the amount of biogeochemical contamination we studied the overall depth distributions of extractable Ca/Mg and pH to elicit information on downward movements. While some downward translocations occurred in profiles on the western flanks of the mountain, very little occurred on the eastern flanks, presumably in response to drier climate there. In all cases where we found downward translocations they occurred mainly in ground paleosols, leaving the buried paleosols unaffected and with apparently valid radiocarbon dates as well as amino acid values.

The D/L ratios for aspartic acid (Table I) were plotted against time (Fig. 4) revealing an overall trend from \sim 0.10 to \sim 0.45 over the late Pleistocene and Holocene. A factor that may contribute to consistent results is that samples are from the same geographical area with similar climatic (especially temperature) and vegetation histories. In other words, close proximity to the equator has tended to minimize the effect of the depth of frost penetration as a factor influencing the rate of racemization. Therefore, amino acid D/L ratios appear to pro-

FIGURE 4. Radiocarbon age vrs. D/L ratios for aspartic acid in the Mount Kenya Quaternary sequence.

Datations au radiocarbone et rapports D/L de l'acide aspartique de la séquence quaternaire du mont Kenya.

TABLE II

Particle distributions for the buried paleosol horizons in the Mount Kenya Quaternary Sequence

— not available

a, Sites are located on Figures 1a and ab.

vide relative age dates for buried soils. In addition, D/L ratios can provide a useful criterion for acceptance or rejection of radiocarbon dates.

Other amino acids (Table I) show variable trends with time that appear useful in age determination. In particular, alanine, glutamic acid, and leucine produced D/L ratios that increased in progressively older samples. Still others, such as phenylalanine and valine, increase slightly or not at all in older buried paleosols, suggesting little utility as a geochronometer.

PARTICLE SIZE

We investigated the possibility that particle size might influence amino acid values either by clay adsorption or by influencing moisture retention. The data in Table Il reveal rather wide variations in percentages of sand, silt and clay, reflecting differences in origin. Even percentages of clay vary widely among the buried A horizons in the paleosol sequence, and although some amounts are essentially nil (e.g. GOR48 and 41) we found no apparent relationship with the height of amino acid peaks.

CONCLUSIONS

Paleosols on Mount Kenya have been dated by radiocarbon and amino acid dating techniques. D/L ratios of amino acids, as determined for these samples, proved useful for relative age determinations and as an important crosscheck on the validity of radiocarbon dates. Aspartic acid proved remarkably reliable as an age indicator giving progressively higher ratios for most samples of increasing age. Several other amino acids including alanine, glutamic acid, and leucine showed distinct age-dependent trends, although not as convincing as aspartic acid. The aspartic acid ratio/age relationships, verified by radiocarbon controls, allow construction of an aspartic acid time diagram that can be used to date samples ranging from a few hundred to $> 42,000$ yrs BP. Further work is needed to add more radiocarbon control points and aspartic acid values to this diagram.

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APPENDICES

APPENDIX I

Paleosol Profile TV42^ª in a debris flow, Upper Teleki Valley

a. Soil site is located on Figure 1a.

b. Soil descriptions follow Birkeland (1984) and Soil Survey Staff (1951, 1975), Colors were taken from Oyama and Takehara (1970) in the moist(m) and dry(d) states. Consistence is given in the moist state.

APPENDIX Il

Soil Profile TV4a^a in late Holocene alluvium, upper Teleki Valley

a. Soil site is located on Figure 1a.

b. Parent material (Cu) is based on criteria established by Hodson (1976).

APPENDIX III APPENDIX IV

Soil Profile TV33^ª in a middle Holocene alluvial fan, Upper Teleki Valley

Elevation: 4160 m Age: postglacial

a. Soil site is located on Figure 1a. b. See Appendix II.

Ab 140-144 Brownish black (1OYR 3/2m) fine sandy

plastic and slightly sticky Coxb 144+ Dull yellow orange (10YR 6/3m) pebbly

plastic and nonsticky

loam, massive, friable consistence, slightly

sandy loam, friable consistence, slightly

Coxb 78-109

Coxb 118-140

Paleosol Profile TV39^ª in a middle Holocene alluvial fan, Upper Teleki Valley

See Appendix II.

Horizon

Depth

APPENDIX V

Paleosol Profile GOR8^a in Liki I Till, Gorges Valley

Elevation: 3475 m Age: Liki Till Vegetation: Ericaceous Zone Horizon Depth (cm) Description" A1 0-11 Grayish yellow brown (10YR 4/2m) silty loam, granular structure, friable, nonplastic and nonsticky IIB2 11-31 Brown (10YR 4/6m) loamy sand, massive structure, friable, nonplastic and nonsticky IIC 31-40 Dull yellowish brown (10YR 5/3m) sandy loam, massive structure, friable, nonplastic and slightly sticky IIIAb 40-45 Black (10YR 2/1m) silty loam, granular structure, friable, nonplastic and nonsticky IVD $45 +$ Till consisting of an open network of pebbles, cobbles, and boulders containing basalt, phonolite and kenyte

a. Soil site is located on Figure 1b.

b. See Appendix II.

APPENDIX Vl

Paleosol Profile GOR52^ª in outwash, Gorges Valley

Grayish brown (7.5YR 4/2m) and grayish yellow brown (10YR 6/2d) sandy loam, massive structure, friable, nonplastic and nonsticky B21bm 69-83

Brownish gray (7.5YR 4/1m), grayish brown (7.5YR 5/2m) and grayish yellow brown (10YR 6/2d) sandy loam, massive structure, loose, nonplastic and nonsticky B22bm 83-97

C1boxm 97-127 Grayish brown (7.5YR 4/2m), grayish yellow brown (10YR 6/2d) silty loam, massive structure, friable, nonplastic and nonsticky

C2boxm 127-142 Grayish brown (7.5YR 4/2m) and dull yellow orange (10YR 6/3d) sandy loam, massive structure, friable, nonplastic and nonsticky

C3boxm 142-163 Brownish gray (7.5YR 4/1m), dull brown (7.5YR 5/3m) and dull yellow orange (10YR 6/4d) silt, massive structure, friable, nonplastic and nonsticky

Cub $163 +$ Grayish brown (7.5YR 5/2m), dull brown Lacustrine (7.5YR 5/3m) and grayish yellow brown ediment (10YR 6/2d) massive structure, firm, plastic and sticky

Soil site is located on Figure 1b.

See Appendix II.

structure,

yellowish structure,

a. Soil site is located on Figure 1b.

b. See Appendix I.

APPENDIX VII

Elevation: 3050 m Age: Teleki Till Vegetation : Lower Afroalpine

Soil Profile LE3^a in Teleki Till and outwash, Lake Ellis

APPENDIX VIII

Soil Profile TV23^a in Teleki and pre-Teleki Tills, lower Teleki Valley

Soil Profile LE4^a in pre-Teleki Till and outwash, Lake Ellis

a. Soil site is located on Figure 1b.

b. See Appendix I.

a. Soil site is located on Figure 1a.

b. See Appendix I.

APPENDIX X

Paleosol Profile GOR48^ª in pre-Teleki Till, Gorges Valley

a. Soil site is located on Figure 1b.

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b. See Appendix I.

APPENDIX Xl

Paleosol Profile GOR41^ª in pre-Teleki outwash, Gorges Valley

a. Soil site is located on Figure 1b.

b. See Appendix I.