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BAFFIN BAY CONTINENTAL SHELF CLAY MINERALOGY

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INTRODUCTION

Little is known about either the sources or mode of deposition of surficial sediment on the continental shelves in Baffin Bay. Interpretation of depositional processes is discussed by Pelletier *et al* 1974, and Marlowe (1968), Grant (1971) and Baker and Friedman (1973) have identified source areas of the coarser sediment fraction. Surficial sediments are in part ice-rafted, and are partly derived by weathering or reworking of Pleistocene till. River sediment is probably mostly trapped in fjords.

We have used X-ray diffraction analysis of the carbonate-free fraction ($<50\mu$) of about 50 sediment samples to try to distinguish sources of fine sediment. A preliminary interpretation of the first 14 of these analyses is given by Piper and Slatt (1976) who describe the methods used, which in general follow those of Biscaye (1965). About 35 of the samples are surficial, the remainder are resedimented material (mostly mudflows) from the central basin of the Bay, believed derived from shelf sediments during Pleistocene glaciation.

RESULTS

Almost all samples contain significant amounts of illite-mica, chlorite, montmorillonite, kaolinite, quartz and feldspars in the $<2\mu$ fraction. The percentage of montmorillonite and kaolinite in surficial samples, expressed as a percentage of the total clay minerals, is shown in Figure 1. Chlorite content usually lies between 10 and 20%. Typical X-ray diffractograms of glycolated samples are shown in Figure 2.

Montmorillonite is commoner in the northern part of the Bay. Low kaolinite contents are found close to the Greenland coast, and in the southern part of the Bay. Illite-rich samples, with high feldspar/quartz ratios, are commonest on the Greenland Shelf, especially close to land. Low feldspar/quartz ratios, and higher percentages of montmorillonite and kaolinite, are found close to shore in areas of known Mesozoic-Cenozoic sediments at Pond Inlet and around the eastern end of Lancaster Sound. Chlorite is most abundant in the northwestern part of the bay. Clay mineralogy appears independent of both coarse-fraction petrology and microfossil abundance (Fig. 3).

DISCUSSION

Most sediments can be interpreted as representing mixtures from four main sources:

(1) A clastic sedimentary source, with low feldspar/quartz ratios and little or no montmorillonite. Illite, chlorite and kaolinite are all present. Sample C170 (Fig. 2), representing this source, is red in colour. This source consists of sediments including red beds.

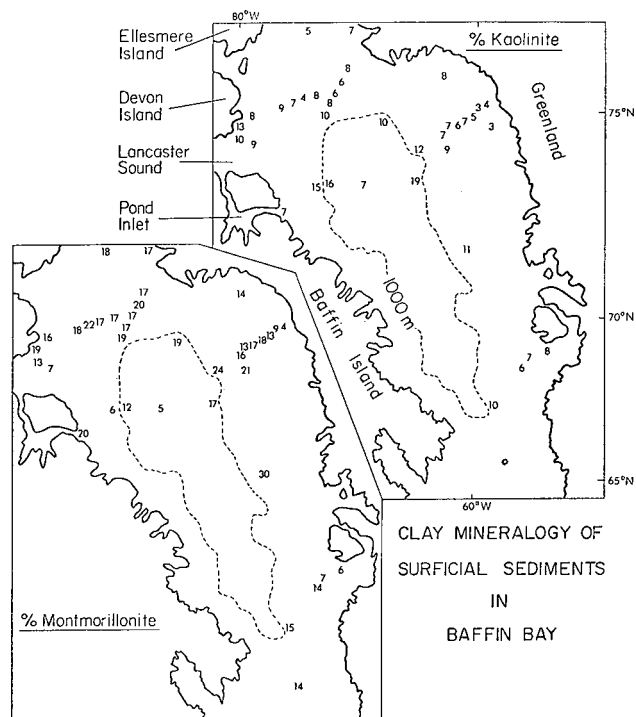


FIG. 1 Percentage of montmorillonite and kaolinite in surficial sediments from Baffin Bay.

(2) A source rich in carbonate, which is found throughout the whole range of grain sizes. This is associated with low feldspar/quartz ratios, an overall low quartz content, and some montmorillonite and illite. Sample 2A (Fig. 2) is dominated by this source material. The Paleozoic limestones and shales of Devon and Ellesmere Islands constitute this source.

(3) A source with abundant illite/mica, a high feldspar to quartz ratio, and crystalline rock fragments abundant in the sand fraction. Sample 194 (Fig. 2) is a good example of a sediment dominated by this source material. The Precambrian areas of Greenland and Baffin Island constitute this source.

(4) A fourth source is inferred, although it is not well represented by any single sample. It is a clastic sedimentary source rich in montmorillonite and probably also chlorite. In sample 52 (Fig. 2), it is mixed with material from source (1), and in 12 (Fig. 2) with material from source (3).

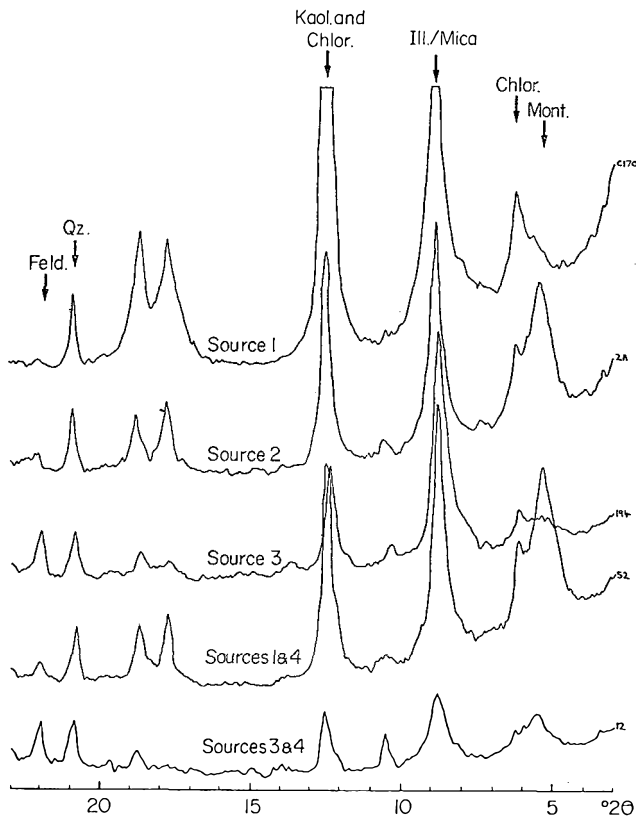
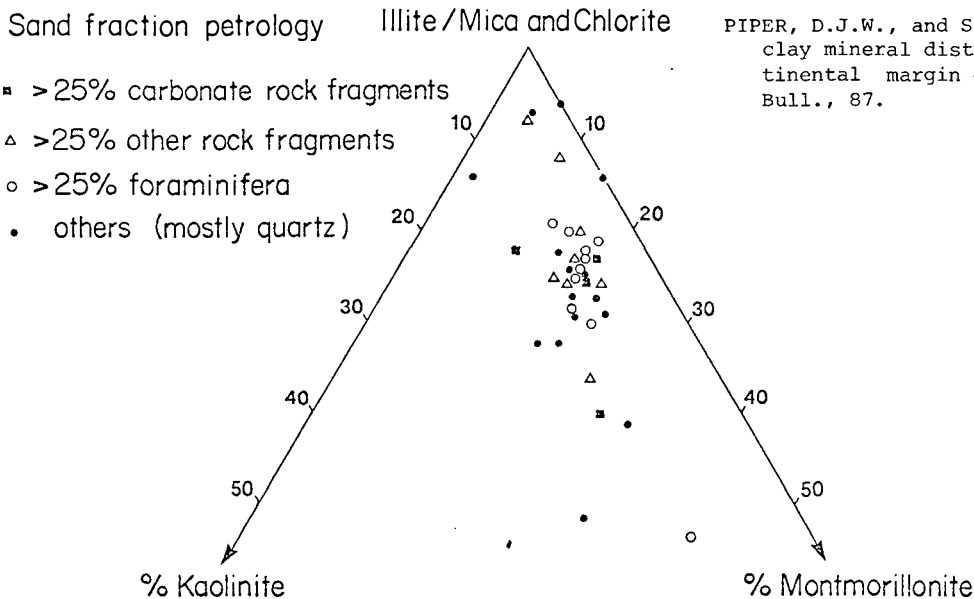


FIG. 2 X-ray diffractograms of selected samples of the glycolated $<2\mu$ fraction from Baffin Bay, illustrating inferred source areas.

FIG. 3 Ternary plot showing relationship between clay mineralogy and sand-fraction petrology.



The lack of correlation between clay mineralogy and both microfossils and coarse-fraction petrology suggests that contemporary transport of sediment plays a minor role in determining clay mineralogy. Local reworking of till is probably the dominant process. Sediments from sources (2) and (3) were probably transported to the continental shelves by grounded ice (nearshore) and icebergs; sediments from sources (1) and (4), which are presumably submerged Cenozoic or Mesozoic continental-shelf strata, were probably eroded by the grounded ice that eroded the glacial valleys found on the shelves.

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