

# The pedogenesis of tin- and sulfide-lode mineralization at True Hill, southwestern New Brunswick

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## Article abstract

The subvolcanic True Hill granite porphyry in southwestern New Brunswick is genetically related to the Beech Hill series of granites of Devonian-Carboniferous age. Three True Hill (TH) greisenised granite porphyry cupolas host Bi-Sn-Mo-W mineralized zones that have similarities to the nearby W-Mo-Bi orebodies at Mount Pleasant (MP).

Tin-bearing lodes (0.03 to 0.66 wt, % Sn) are developed along faults and fractures that cut both mineralized (Bi-Sn-Mo-W) and unmineralized True Hill granite porphyry and enclosing metasedimentary rocks of the Waweg Formation. Within these lodes, particularly the central Main lode, two types of mineralization are present, an early chlorite-sulfide assemblage (Stage 1) and a later hematite-phengite-cassiterite assemblage (Stage 2) that replaces, in part, the earlier Stage 1 assemblage. The Stage 1 assemblage consists of Fe-rich sepiotrichite, pyrite, low-Fe sphalerite, magnetite, chalcocite, galena, Pb-Bi sulfides, tennantite, arsenopyrite and native Ag. Fe-Mg-Mn metasomatism responsible for chloritization was coincident with leaching of alkali and alkali-earth elements during feldspar hydrolysis. The Stage 1 assemblage (chl-sulfide) was formed at low temperature (<200°C) and low oxygen fugacity. The Stage 2 assemblage includes hematite, phengite, cassiterite and monazite with minor late quartz and fluorite. This assemblage was deposited at slightly higher temperatures (200°–300°C) higher oxygen fugacity (above hematite magnetite buffer) and at moderate pH's (<musc-Kfs buffer). The overprinting of Stage 2 assemblages on portions of Stage 1 locally produces complex textures within the lodes at True Hill. The composition of both Stage 1 and 2 assemblages in the lodes is controlled by the (i) host rock composition and (ii) the mineralizing hydrothermal fluids. The deposition of Sn and LREE from Stage 2 fluids formed cassiterite and monazite along the reaction front between Stage 2 assemblages and chlorite-sulfide (Stage 1) and granite. Deposition of cassiterite probably occurred in response to (i) an increase in the Fe-Cl complexing at the expense of Sn- and LREE-C1 complexes during interaction with Stage 1 assemblages, (ii) an increase in pH due to the feldspar hydrolysis in the granite and (iii) cooling of the fluid by the host rock. Although uneconomic, tin-sulfide lodes at True Hill have similarities with lodes developed within the North Zone of Mount Pleasant (New Brunswick) and Cornwall (England). However, the association of hematite and phengite with cassiterite mineralization is unique to this deposit.

## The petrogenesis of tin- and sulfide-lode mineralization at True Hill, southwestern New Brunswick

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The subvolcanic True Hill granite porphyry in southwestern New Brunswick is genetically related to the Beech Hill series of granites of Devonian-Carboniferous age. Three True Hill (TH) greisenised granite porphyry cupolas host Bi-Sn-Mo-W mineralized zones that have similarities to the nearby W-Mo-Bi orebodies at Mount Pleasant (MP).

Tin-bearing lodes (0.03 to 0.66 wt. % Sn) are developed along faults and fractures that cut both mineralized (Bi-Sn-Mo-W) and unmineralized True Hill granite porphyry and enclosing metasedimentary rocks of the Waweg Formation. Within these lodes, particularly the central Main lode, two types of mineralization are present, an early chlorite-sulfide assemblage (Stage 1) and a later hematite-phengite-cassiterite assemblage (Stage 2) that replaces, in part, the earlier Stage 1 assemblage. The Stage 1 assemblage consists of Fe-rich septeclorite, pyrite, low-Fe sphalerite, magnetite, chalcopyrite, galena, Pb-Bi sulfides, tennantite, arsenopyrite and native Ag. Fe-Mg-Mn metasomatism responsible for chloritization was coincident with leaching of alkali and alkali-earth elements during feldspar hydrolysis. The Stage 1 assemblage (chl-sulfide) was formed at low temperature (<200°C) and low oxygen fugacity. The Stage 2 assemblage includes hematite, phengite, cassiterite and monazite with minor late quartz and fluorite. This assemblage was deposited at slightly higher temperatures (200°-300°C), higher oxygen fugacity (above hematite-magnetite buffer) and at moderate pH's (<muscovite-Kfs buffer). The overprinting of Stage 2 assemblages on portions of Stage 1 locally produces complex textures within the lodes at True Hill. The composition of both Stage 1 and 2 assemblages in the lodes is controlled by the (i) host rock composition and (ii) the mineralizing hydrothermal fluids. The deposition of Sn and LREE from Stage 2 fluids formed cassiterite and monazite along the reaction front between Stage 2 assemblages and chlorite-sulfide (Stage 1) and granite. Deposition of cassiterite probably occurred in response to (i) an increase in the Fe-Cl complexing at the expense of Sn- and LREE-Cl complexes during interaction with Stage 1 assemblages, (ii) an increase in pH due to the feldspar hydrolysis in the granite and (iii) cooling of the fluid by the host rock. Although uneconomic, tin-sulfide lodes at True Hill have similarities with lodes developed within the North Zone of Mount Pleasant (New Brunswick) and Cornwall (England). However, the association of hematite and phengite with cassiterite mineralization is unique to this deposit.

Le porphyre granitique subvolcanique de True Hill (Nouveau-Brunswick méridional) est lié par sa genèse au cortège granitique dévonien-carbonifère de Beech Hill. Trois coupôles de porphyre granitique de True Hill (TH) altéré en greisen renferment des zones minéralisées en Bi-Sn-Mo-W et qui montrent des similitudes avec les gisements voisins de W-Mo-Bi présents à Mount Pleasant (MP).

Des faisceaux filoniens porteurs d'étain (0,03 à 0,66% pds. Sn) s'installent le long de failles et de cassures recoupant les porphyres granitiques de True Hill, tant minéralisés que stériles, ainsi que leur encaissant métasédimentaire (Formation de Waweg). Au sein de ces faisceaux, et surtout du faisceau central principal, on note deux types de minéralisation: un assemblage précoce à chlorite-sulfure (stage 1) et un assemblage tardif à hématite-phengite-cassitérite (stage 2) qui remplace en partie le premier. L'assemblage représentant le stage 1 consiste en de la septeclorite riche en Fe, de la pyrite, de la blende pauvre en Fe, de la magnétite, de la chalcopyrite, de la galène, des sulfures à Pb-Bi, de la tennantite, du mispickel et de l'argent natif. La métasomatose des Fe-Mg-Mn, dont provient la chloritisation, s'accompagne d'un lessivage des éléments alcalins et alcalino-terreux durant l'hydrolyse des feldspaths. L'assemblage dénotant le stage 1 (chlorite-sulfures) se forma à basse température (<200°C) et faible fugacité d'oxygène. L'assemblage traduisant le stage 2 comporte de l'hématite, de la phengite, de la cassitérite, ainsi que de la monazite avec de faibles quantités de quartz et de fluorine tardifs. Cet assemblage fait intervenir des températures un peu plus élevées (200°C à 300°C), une fugacité d'oxygène plus élevée (au-dessus du tampon hématite-magnétite) et des pH modérés (en-dessous du tampon des feldspaths potassiques de la muscovite). La surimpression des assemblages représentant

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le stage 2 sur des portions du stage 1 produit localement des textures complexes au sein des faisceaux filoniens de True Hill. La composition des assemblages dénotant tant le stage 1 que le stage 2 au sein de ces faisceaux est régie par: (i) la composition de l'encaissant et (ii) les fluides minéralisateurs hydrothermaux. Le dépôt d'étain et de T.R. légères à partir des fluides du stage 2 engendra de la cassitérite et de la monazite le long du front réactionnel, entre les assemblages témoins du stage 2, les chlorite-sulfures et le granite. Le dépôt de la cassitérite est probablement le fruit de: (i) une augmentation de la formation de complexes à Fe-Cl aux dépens des complexes à Sn et à Cl-T.R. légères durant l'interaction avec des assemblages du stage 1, (ii) une élévation du pH par suite de l'hydrolyse des feldspaths du granite, et (iii) le refroidissement du fluide par l'encaissant. Bien que de nature non économique, les faisceaux filoniens à étain et à sulfures présents à True Hill ne sont pas sans rappeler les faisceaux développés dans la Zone Nord de Mount Pleasant (Nouveau-Brunswick) et à Cornwall (Angleterre). Cependant, ce gisement se singularise par l'association d'hématite et de phengite avec la minéralisation en cassitérite.

[Traduit par le journal]

## INTRODUCTION

In southwestern New Brunswick several Sn-W-Mo deposits are genetically associated with Late Devonian to Early Mississippian peraluminous granites (Ruitenberg and Fyffe, 1982; Ruitenberg and McCutcheon, 1985). Of these the deposit at Mount Pleasant is the most well documented (Fig. 1). At Mount Pleasant W-Mo-Bi ore zones are associated with fine-grained granite that formed in a shallow, subvolcanic environment (Ruitenberg, 1963; Dagger, 1972; Parrish and Tully, 1978; Pouliot *et al.*, 1978; Davis and Williams-Jones, 1985; Kooiman *et al.*, 1986). Structurally-controlled cassiterite-sulfide lodes crosscut fine-grained granite and associated W-Mo-Bi mineralization in the North Zone and Fire Tower Zone at MP (Kooiman *et al.*, 1986). They interpreted these lodes to be related to granite porphyry that intruded older fine-grained granite and associated W-Mo orebodies. Some sulfide-poor, cassiterite-rich zones are also associated with this younger granite. In the Fire Tower Zone (MP; Fig. 1), these zones are poorly defined but are associated with irregularly developed chlorite, topaz, fluorite, biotite alteration (Kooiman *et al.*, 1986). In the North Zone, near-surface replacement deposits of cassiterite, stannite, pyrrhotite, pyrite,

arsenopyrite, sphalerite, and chalcopyrite with quartz, fluorite and topaz have reserves of approximately 2500 tonnes with an average grade of 0.6 wt. %  $\text{SnO}_2$ , 2.3 wt. % Zn, 0.3 wt. % Cu and 0.36 wt. % Pb (Mulligan, 1975). Reserves in the Deep Tin Zone below the North Zone (MP) are approximately 2.14 million tonnes grading 0.45 wt. %  $\text{SnO}_2$ , 0.8 wt. % Zn and 0.06 wt. % W (Sinclair *et al.*, 1988). Total reserves in the Deep Tin Zone, Contact Crest, Contact Flank and Endozone deposits in the North Zone at Mount Pleasant contain approximately 5.9 million tonnes of 0.79 wt. % Sn (The Northern Miner, October 1985).

Located 4.5 km west of Mount Pleasant, the True Hill granite porphyry has intruded Silurian (?) metasedimentary rocks and locally metamorphosed them to biotite-andalusite-cordierite grade hornfelses (Figs. 2, 3). The True Hill granite is petrographically and chemically similar to the Mount Pleasant fine-grained granite (Lentz *et al.*, 1988). Endogranitic and exogranitic mineralized zones are associated with three exposed granite cupolas (TH; Fig. 1). Low grade Bi-Sn-Mo-W mineralized zones (TH), similar to the W-Mo-Bi orebodies at Mount Pleasant, are associated with greisen alteration in two of the three cupolas (Lentz *et al.*, 1988). Cassiterite-sulfide lodes occur along faults that crosscut the True Hill granite and associated greisen.

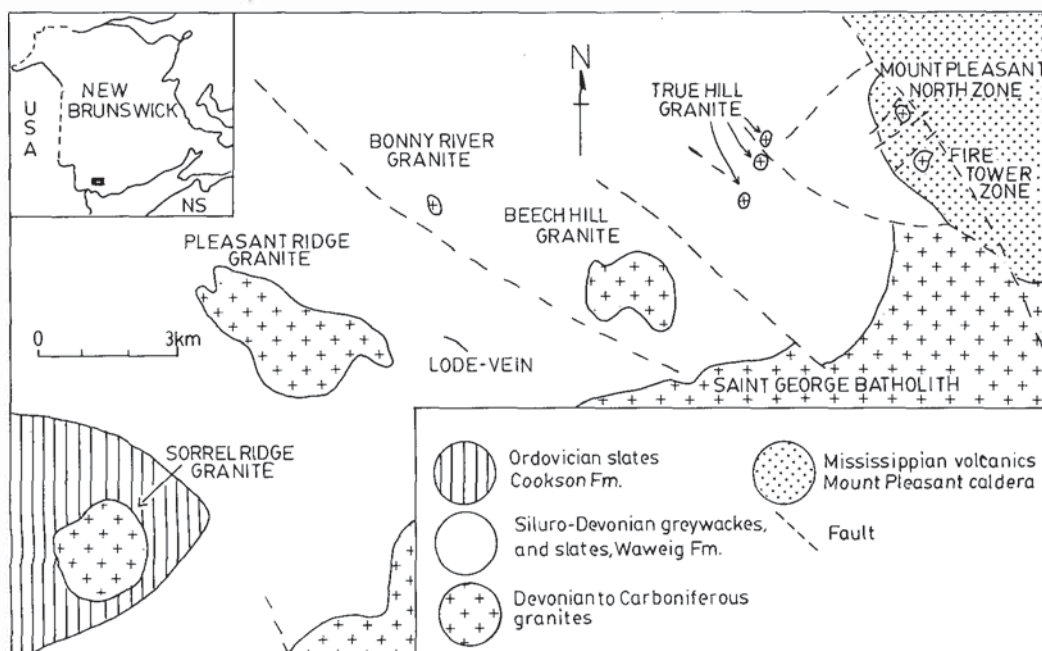


Fig. 1. Geological map of the Mount Pleasant area, southwestern New Brunswick (modified after Taylor *et al.*, 1985).

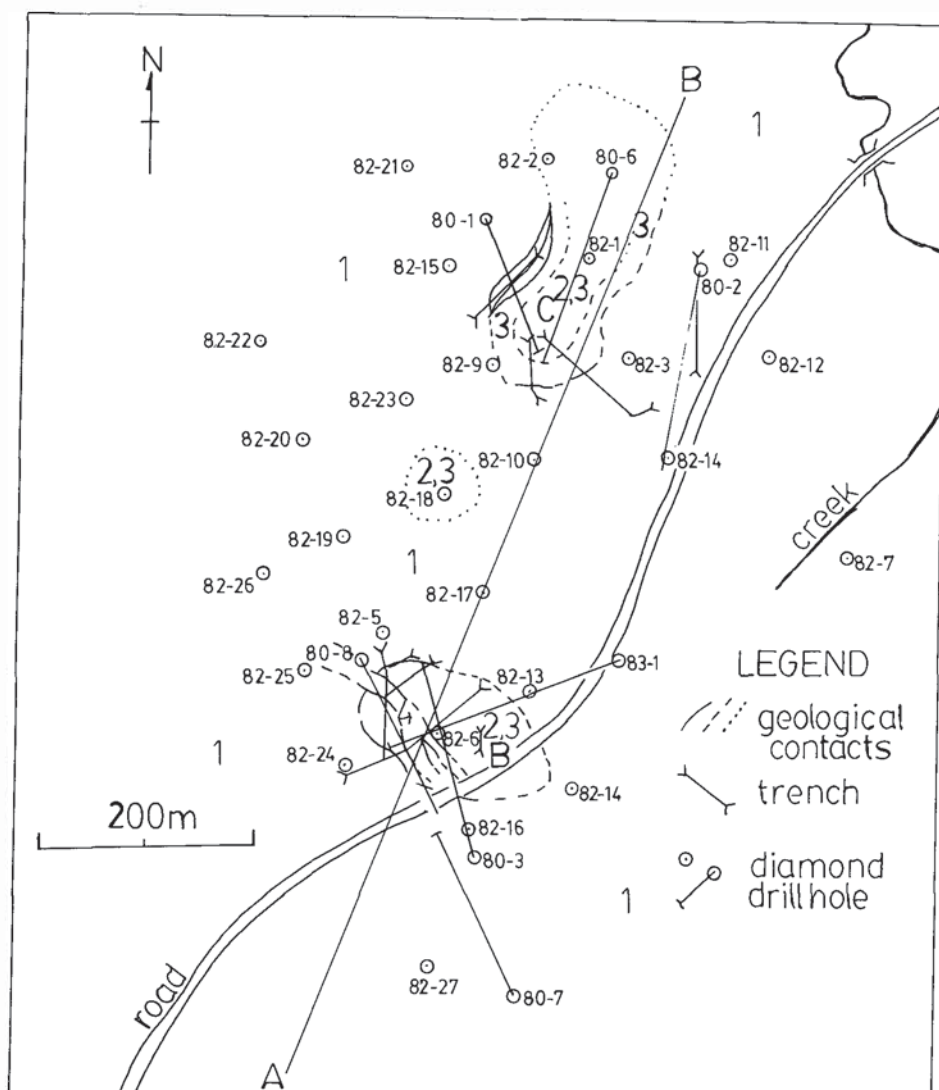


Fig. 2. True Hill geological map of cupolas B and C with locations of diamond drill hole (DDH) and trench locations: 1 - Metasedimentary rocks, 2 - Granite porphyry, 3 - Greisen, 4 - Tin-sulfide lodes. (A-B cross section see Fig. 3.)

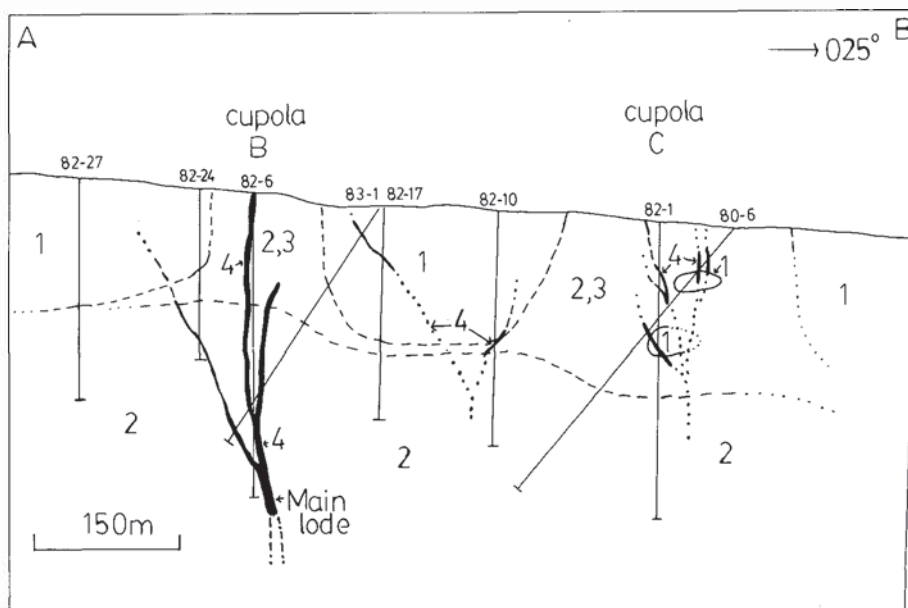


Fig. 3. True Hill geological cross section (A-B) of cupola B and C (see Fig. 2) with DDH locations.



The purpose of this study is two-fold; to evaluate the relationship of lode mineralization at True Hill to earlier Bi-Sn-Mo-W greisen mineralized zones associated with the True Hill granite and to compare the tin lode mineralization of True Hill with the North Zone at Mount Pleasant. In addition, this study attempts to evaluate the mechanisms responsible for the transport and deposition of tin within the True Hill lode. The good geologic control of the Main lode provided by surface exposure and drill intersections, forms the basis for most of the petrographic and geochemical analyses.

### ANALYTICAL TECHNIQUES

Ten samples of various types of mineralized lodes were analysed in duplicate by X-ray Assay Limited (Don Mills, Ontario) for major and trace elements. Analytical precision is  $\pm 3\%$  for major elements and  $\pm 5\%$  for trace elements having abundances three times their detection limit. A selected True Hill granite composition was obtained by averaging several analyses (Lentz *et al.*, 1988).

The sulfur isotopes were determined on  $\text{SO}_2$  gas prepared by oxidizing pyrite with  $\text{Cu}_2\text{O}$  at  $1050^\circ\text{C}$  for 15 minutes. Analyses were performed on the micromass 602E mass spectrometer (Derry Laboratory, University of Ottawa), calibrated with an internal sulfur standard. The analyses are expressed relative to the Canyon Diablo Troilite (CDT) and the accuracy and precision are  $\pm 0.1$  ‰ based on replicate analyses of the standard.

A Cambridge S4-10 stereoscan (SEM) was used for quantitative energy dispersive (SEM-EDS) analyses of minerals on carbon-coated polished thin sections at the University of New Brunswick. Both natural and synthetic sulfide and silicate standards were used to calibrate the EDAX system. Reliable analytical totals for chlorite were considered to be between 85 to 90 wt. % due to the uncertainty in hydroxyl content of septechnorites and chlorites. Only one phengite analysis was considered reliable with a total within 2 wt. % of 95 wt. % (Table 1).

Sphalerite, cassiterite and rutile were renormalized to 100 wt. % if within 2 wt. % for comparative purposes.

### GEOLOGY OF THE TRUE HILL AREA

Three cupolas of True Hill granite porphyry designated A, B and C are exposed in outcrop and 36 drill holes (Figs. 2, 3). The granite porphyry contains quartz (30-35 vol. %), K-feldspar (25-30 vol. %), plagioclase (30 vol. %) and biotite (5 vol. %) phenocrysts with a compositionally similar matrix. Accessory minerals occurring with biotite are titanite, zircon, magnetite and ilmenite. Major- and trace-element whole-rock analyses indicate that the True Hill granite is highly fractionated (Lentz *et al.*, 1988). Lentz *et al.*, (1988) determined that the True Hill granite crystallized at approximately 100 MPa  $P_{\text{H}_2\text{O}}$  and between  $700^\circ$  and  $730^\circ\text{C}$ .

The Silurian (?), greenschist-grade, metagreywackes and slates of the Waweig Formation have upright open folds with undulatory northeast trending plunges. They are compositionally heterogeneous, containing variable proportions of sericite, chlorite, epidote, quartz, plagioclase, tourmaline, zircon, titanite, ilmenite and sulfides. Locally, biotite-cordierite-andalusite hornfels occur proximal to the intrusion. Compositional variation within the metasedimentary rocks precludes a detailed analysis of the major- and trace-element compositions of the metasediment-hosted lodes.

Bi-Sn-Mo-W greisen zones are located in cupolas B and C of the True Hill granite porphyry (Lentz *et al.*, 1988). Three types of greisen alteration are present: (i) quartz-sericite, (ii) topaz-quartz-sericite and (iii) chlorite-quartz-sericite. Within the metasedimentary rocks, chlorite is more common with anatase and tourmaline. Commonly, greisen assemblages have associated fluorite, pyrite, magnetite, hematite, sphalerite, chalcopyrite, galena, bismuthinite, cassiterite, molybdenite and wolframite.

Table 1. Chemical composition of micas.

wt. %	1	2	3	4	5	6	7	8	9	10
SiO <sub>2</sub>	23.8	24.6	23.3	22.1	23.0	22.1	21.8	22.2	22.0	47.3
Al <sub>2</sub> O <sub>3</sub>	19.0	19.7	18.9	17.5	18.1	18.4	21.7	27.0	27.6	29.5
FeO <sub>T</sub>	35.9	38.2	39.4	40.5	33.0	32.9	43.0	35.8	34.4	4.91
MgO	2.47	2.35	2.73	2.23	3.55	3.28	1.32	1.0	4.7	0.23
MnO	4.11	3.06	3.02	3.06	8.82	8.75	1.22	0.4	-	0.21
CaO	0.00	0.00	0.10	0.00	0.00	0.00	0.07	1.3	-	0.09
Na <sub>2</sub> O	0.00	0.00	0.00	0.00	0.00	0.00	0.00	-	-	0.27
K <sub>2</sub> O	0.00	0.00	0.00	0.06	0.00	0.00	0.08	-	-	10.5
Cr <sub>2</sub> O <sub>3</sub>	-	-	0.00	0.17	0.00	0.00	0.10	-	-	-
NiO	-	-	0.11	0.00	0.00	0.00	0.00	-	-	-
Total	85.2	88.7	87.6	85.6	86.4	85.4	89.3	89.1	89.2	93.0

\*Samples 1 @ 83-1-317.4 m; 2, 3, 4 @ 82-13-53.2 m; 5 and 6 @ 82-13-47.7 m; 7 @ 80-5-78.4 m, 8 from Mount Pleasant also includes 32.5 wt. % FeO and 3.7 wt. % Fe<sub>2</sub>O<sub>3</sub> (Petruk, 1964); 9 has 30.2 wt. % FeO and 4.7 wt. % Fe<sub>2</sub>O<sub>3</sub> from Cornwall, England (Petruk, 1964); 10 is phengite from Main lode 83-1-317.4 m.



## GEOLOGY OF TIN LODES

Tin lodes at True Hill comprise zones consisting of veins, veinlets and disseminated mineralization occurring along faults, fractures and/or foliations in metasedimentary rocks and granite porphyry. They are best developed in cupola B where the Main lode crosscuts greisenized granite porphyry near the surface and unaltered granite porphyry at depth (Figs. 2, 3, 4, 5A). The Main lode is 0.5 m to 3 m wide, has a strike length of at least 500 m and exceeds 300 m in depth. The Main lode is predominantly a replacement body that was controlled by a vertical fault and associated fracture zones in the granite porphyry. However, intense alteration has all but obliterated evidence of faulting except for the rare occurrence of silicified breccia within the Main lode. Linear magnetic anomalies are coincident with the Main lode but discontinuous along strike.

An early chlorite-sulfide assemblage (Stage 1) completely replaced granite and metasedimentary rocks and was subsequently overprinted by the Stage 2 hematite-phengite-cassiterite assemblage (Fig. 4). In some portions of the lode, Stage 2 fluids have partially replaced the chlorite-sulfide assemblage to produce a chlorite-bearing phengite-hematite-cassiterite assemblage. The irregular distribution of Stage 1 and 2 assemblages within samples of the lode produces complex chemical compositions. Cassiterite, introduced by the Stage 2 fluid, is deposited adjacent to recrystallized and altered chlorite-sulfide assemblages (Stage 1) and, to a lesser extent, unaltered granite (Fig. 4). No Sn-bearing phases have been identified in the chlorite-sulfide assemblage.

### Chlorite-Sulfide Assemblage (Stage 1)

The chlorite-sulfide alteration consists dominantly of chlorite, with lesser amounts of sulfides and rare magnetite. XRD analysis revealed low intensity (001) reflections characteristic of poorly crystalline septeclorite. The (060) (1.56 to 1.60 Å) reflection is indicative of high ferrous iron content (Carroll, 1970; Lentz, 1986). Analyses of 7 septeclorites from the metasediment-hosted lodes, indicate that they are iron-rich brunsvigite to ripidolite (Hey, 1954; Table 1). Analyses of chlorite from tin lodes at Mount Pleasant and Cornwall, England (Petruk, 1964) are similar to the True Hill chlorite analyses in having high ferrous Fe content (8 and 9 in Table 1). However, these chlorites have lower  $\text{SiO}_2/\text{Al}_2\text{O}_3$  ratios, and higher MnO contents than those of Mount Pleasant and Cornwall, England (Table 1).

Pyrite is the most common sulfide within the lodes. It is idiomorphic (2-10 mm) and usually evenly disseminated in a dominantly chloritic groundmass. Sphalerite is very common, occurring with galena, chalcopryrite, native silver, bismuth sulfides and pyrite, and typically contains abundant fine inclusions of chalcopryrite. Seven analyses of sphalerite from 3 localities have between 0.5 to 2 wt. % Fe with minor Cu, Cd and Mn (Table 2). Tennantite was identified in a metasediment-hosted lode where it is associated with pyrite and chalcopryrite. Galenobismutite and cosalite occur with sulfides in the Main lode and arsenopyrite (0.2-0.05 mm) was found in metasediment-hosted

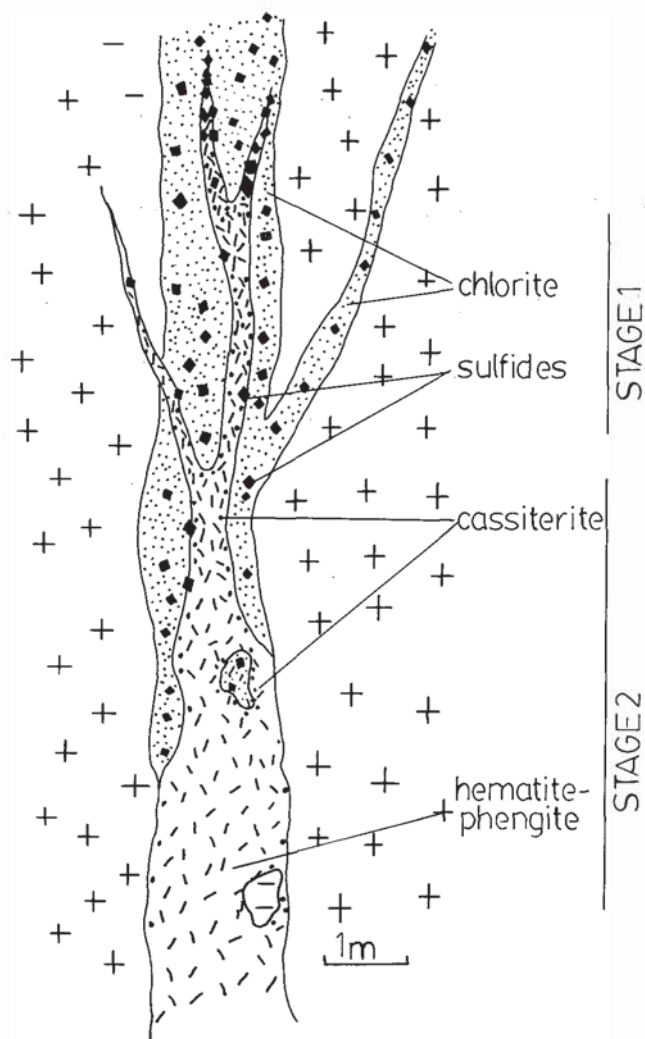


Fig. 4. Schematic section of a portion of the Main lode at True Hill illustrating the relationship between the chlorite-sulfide lode (Stage 1), the cassiterite-hematite-phengite lode (Stage 2), and the True Hill granite.

lodes. Native Ag (1-10 µm) is hosted within sphalerite, galena, pyrite and more rarely in chlorite. Magnetite (0.2 mm) is rare but occurs with pyrite and sphalerite as euhedral, finely disseminated grains (Fig. 5B). Rhodochrosite with a composition of  $(\text{Mn}_{0.83}\text{Fe}_{0.07}\text{Ca}_{0.05})\text{CO}_3$  occurs in a metasediment-hosted lode (Lentz, 1986).

### Hematite-Phengite-Cassiterite Assemblage (Stage 2)

The second mineralizing episode (Stage 2) at True Hill consists of cassiterite with hematite and phengite. Phengite, specularite, quartz and fluorite are the predominant minerals of this assemblage. As with the chlorite-sulfide assemblage, brecciated host rocks are almost entirely replaced, except where silicification has locally preserved breccia within the Main lode. Cassiterite is weakly disseminated throughout the zone of hematite-phengite alteration with more abundant cassiterite near the chlorite-sulfide alteration.

The phengite is generally coarse grained (2 to 5 mm) and occurs as radiating aggregates superimposed on early fine-



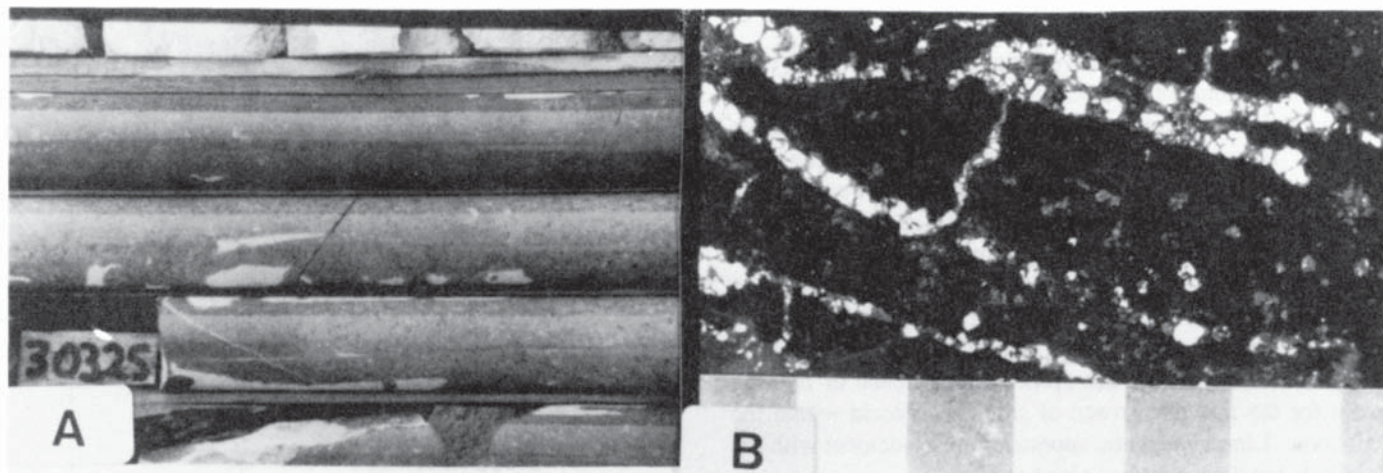


Fig. 5. (A) Unaltered True Hill granite porphyry beside Main lode (DDH 83-01-303 m; 13 cm across). (B) Stage 1 chlorite-magnetite lode (dark) veined by pyrite (light) which is in turn replaced by hematite (grey) (Stage 2; Trench 7, cupola C).

Table 2. Compositions of sphalerite.

wt. %	1	2	3	4	5	6	7
Zn	67.0	66.3	64.3	65.4	66.6	67.5	68.2
Fe	1.50	2.00	0.50	1.43	1.48	1.27	1.3
Cu	0.00	0.27	0.27	0.34	0.00	0.31	0.84
Cd	0.31	0.27	0.44	0.33	0.61	-	-
S	31.2	31.2	34.5	32.5	31.3	30.9	29.7
Total	100.0	100.0	100.0	100.0	100.0	100.0	100.0

\*1 - rim; 2 - core of sphalerite (82-13-53.2 m); 3 - (82-24-162 m); 4 to 7 - (80-15-38.1 m). Renormalized samples within 2% of 100%.

grained chlorite (Figs. 6A, B, D, 7A, B, C). A single analysis of phengite from the Main lode is presented in Table 1. Two samples have a  $d(001)$  spacing of  $9.8 \text{ \AA}$  which is characteristic of Fe-rich muscovites and phengites. In addition, the micas are transitional between the 1M and 2M polytypes. Carroll (1970) found the transition from the 1M to 2M Fe muscovite to occur between  $200^\circ$  and  $300^\circ\text{C}$ .

Specularite is fine to medium grained ( $<0.1 \text{ mm}$ ) and commonly pseudomorphs and/or replaces pre-existing sulfides (Figs. 6A, D, E, 7A, B, D). Fluorite and quartz occurs mainly in the interstices of phengite and along reaction margins adjacent to the chlorite-sulfide assemblage (Fig. 6D). Medium- to fine-grained hydrothermal quartz is minor and is variably present in portions of the Main lode. Fluorite is most closely associated with the

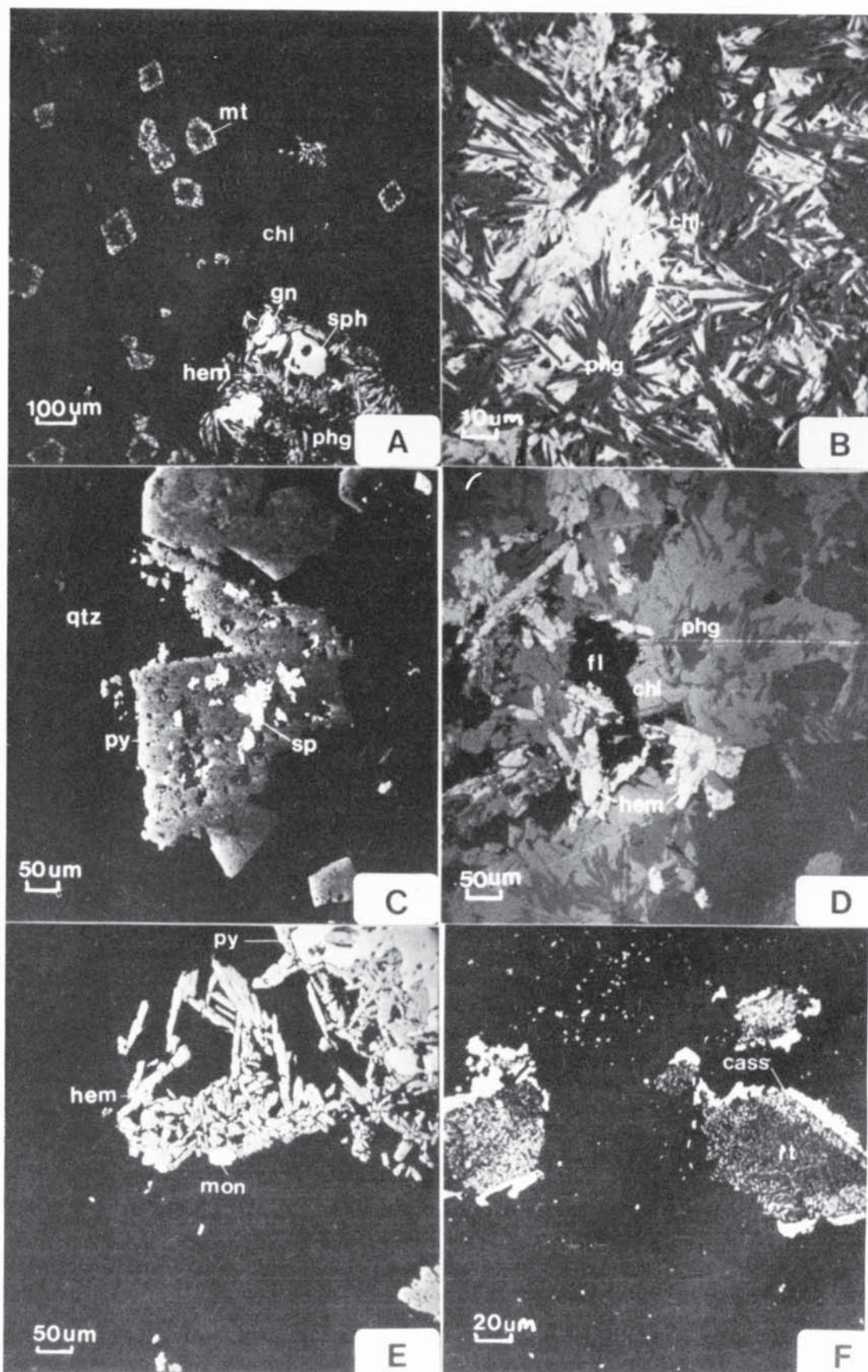
metasediment-hosted lodes due to the high CaO content which favours the deposition of fluorite.

Schorl occurs as very fine- to coarse-grained acicular crystals with chlorite in altered metasedimentary rocks. Coarse-grained, dark blue-green, zoned schorl forms acicular splays with quartz, phengite, rutile and cassiterite. However, boron was not found in the few samples analysed. Rutile is common in lodes hosted by metasedimentary rocks and was formed by oxidation of Ti-rich biotite and ilmenite. Cassiterite mantles fine-grained, euhedral rutile (Fig. 7F). Rutile contains 1.3 wt. %  $\text{SnO}_2$  and 5.4 wt. %  $\text{Fe}_2\text{O}_3$  (Table 3). At Mount Pleasant, rutile and cassiterite are hosted within greisenized porphyry (Petruck, 1964, 1973; Dagger, 1972; Pouliot *et al.*, 1978).

Cassiterite is predominantly concentrated along the reaction

Fig. 6. (A) Disseminated magnetite (mt) + (pyrite (py)), galena (gn), and sphalerite (sph) in chlorite matrix (chlorite-sulfide assemblage) being replaced by hematite (hem) and phengite (phg) (S.E.M. photo; DDH 80-3-315 ft.). (B) Phengite (phg) replacing chlorite (chl) during interaction of the later (hematite-phengite) fluid with the pre-existing chlorite-sulfide assemblage (DDH 83-1-317.4 m). (C) Remobilized pyrite (py) enclosing sphalerite (sph) in a quartz (qtz) matrix related to hematite-phengite assemblage (S.E.M. photo; DDH 83-1-289.5 m). (D) Fluorite (fl), phengite (phg), and hematite (hem) replacing chlorite (chl) from pre-existing chlorite-sulfide assemblage (S.E.M. photo; DDH 80-3-350 ft.). (E) Hematite (hem) replacing pyrite (py) with monazite (mon) and hosted in phengite (S.E.M. photo; Cupola C Trench 7b). (F) Rutile (rt) mantled by cassiterite (cass) in hematite-phengite and tourmaline-rich metasediment-hosted lode (S.E.M. photo; DDH 80-5-561 ft.).







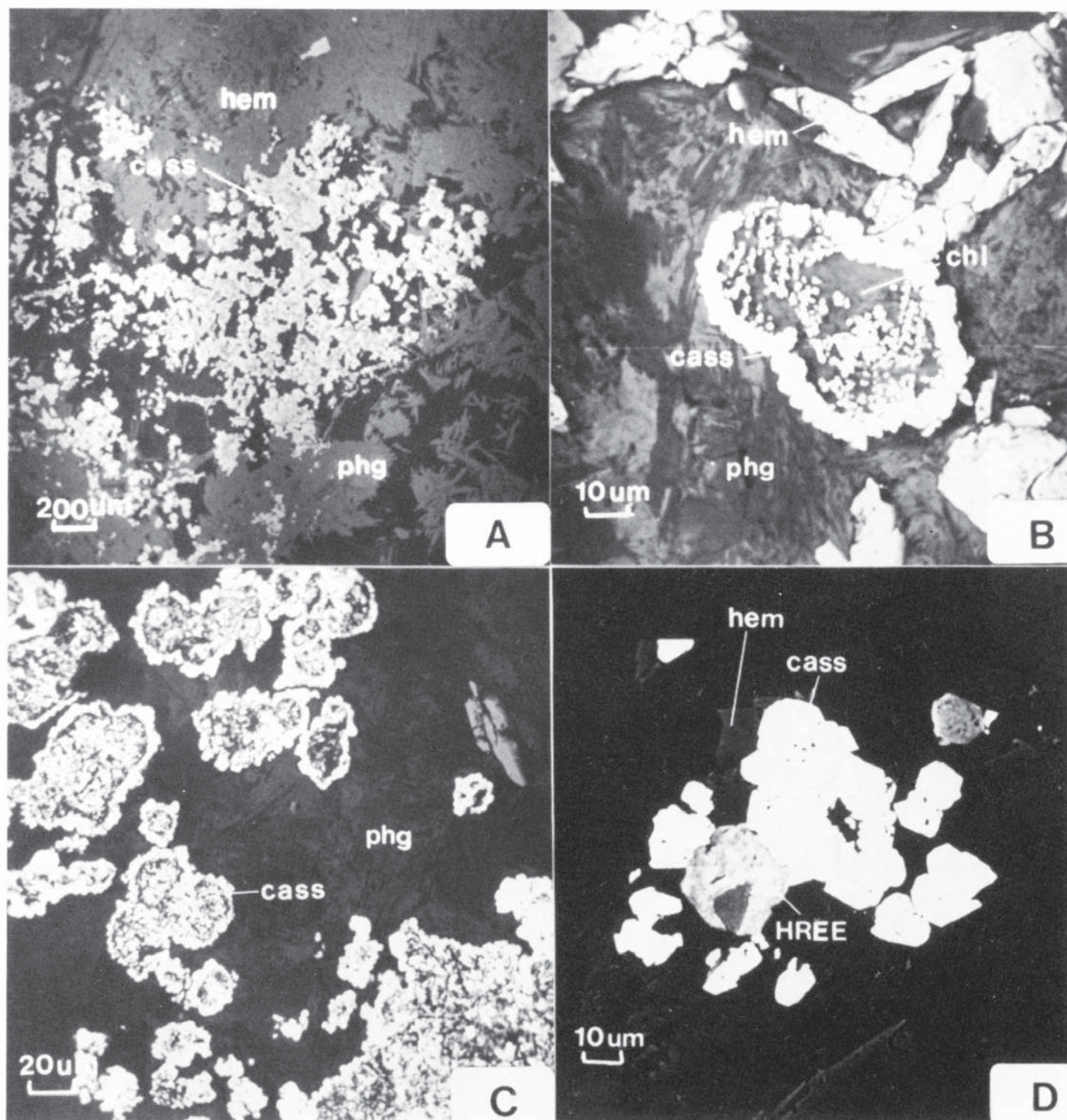


Fig. 7. (A) Cassiterite (cass) hosted in hematite (hem) and phengite (phg) (S.E.M. photo; DDH 83-1-316.5 m). (B) Cassiterite (cass) with chlorite (chl) in hematite (hem) and phengite (phg) assemblage (S.E.M. photo; 83-1-284 m). (C) Cassiterite (cass) mantling cassiterite crystals and coalescing to form glomeroporphyritic crystals in phengite (phg) (S.E.M. photo; DDH 83-1-284 m). (D) HREE-phosphate mineral (xenotime?) associated with hematite (hem), phengite (dark) and cassiterite (cass) assemblage (S.E.M. photo; DDH 83-1-290.9 m).

front between Stage 1 and 2 alteration, where it occurs as finely disseminated colliform crystals, in places coalescing to form dense aggregates (Fig. 7A, B). Metasediment-hosted cassiterite has higher  $\text{TiO}_2$  (3.0 wt. %) and lower Fe whereas a granite-hosted cassiterite from the Main lode has only 0.9 wt. %  $\text{TiO}_2$  and higher Fe indicative of the low Ti content of the granite (Table 3).

Other trace minerals associated with the second stage of mineralization include monazite (Fig. 6E) and a HREE phosphate mineral (xenotime?; Fig. 7D). Both of these minerals commonly occur with cassiterite in most portions of the lodes and are also concentrated along the reaction margin.



Table 3. Compositions of rutile and cassiterite.

wt. %	1	2	3
SnO <sub>2</sub>	97.0	96.5	1.30
TiO <sub>2</sub>	2.98	0.90	93.8
Fe <sub>2</sub> O <sub>3</sub>	0.00	2.27	4.89
MnO	0.00	0.24	-
Total	100.0	100.0	100.0

\*1 - Cassiterite (metasediment-hosted with rutile; Fig. 6F; 80-5-561 ft); 2 - Cassiterite (granite-hosted with hematite-phengite; 83-1-284 m); 3 - Rutile (metasediment-hosted mantled by cassiterite; Fig. 6F; 80-5-561 ft).

### Remobilization Textures

It is evident that fracturing precedes Stage 2 and postdates Stage 1 mineralization (Fig. 5B); hence, two discrete mineralizing episodes are indicated. The boundary between the two stages may be sharp (near surface) to diffuse (at depth) depending on the extent of chemical interaction. This "reaction front" is represented by an irregular zone with variably replaced chlorite-sulfide assemblages (Fig. 6A, B, C, D). Complex textural features were formed by the alteration of the Stage 1 assemblage to the Stage 2 assemblage. Septechlorites, for instance, increase in grain size with increasing proximity to Stage 2 assemblages (Fig. 6B) and they approach an equilibrium 14 Å-chlorite structure (increase 001 reflection) at temperatures >200°C. Chlorite has been identified as part of the hematite-phengite assemblage in areas of considerable reaction with the chlorite-sulfide assemblage and may represent a stable phase.

There is abundant textural evidence for alteration and replacement of septechlorite and sulfides (reaction rims, coronas, embayments) during Stage 2 mineralization along the reaction front between the two assemblages (Figs. 5B, 6A, C, E). In some areas, pyrite has apparently been remobilized (dissolved and redeposited) by Stage 2 fluids, as indicated by coarse grains of pyrite in fractured chlorite-magnetite lode (Fig. 5B) or hosted in silicified lode. Commonly it is subsequently replaced by hematite during desulfidation (Fig. 5B) and hematite often replaces pyrite (Fig. 6E).

### GEOCHEMISTRY OF LODES

Mineralogical differences between Stage 1 chlorite-sulfide and Stage 2 hematite-phengite-cassiterite assemblages are reflected by major chemical differences between the two types of mineralization (Table 4).

The major- and trace-element composition of the lodes reflect the dominant minerals in each type. Identification of the distinguishing chemical characteristics between the two types of mineralization is problematic because samples analyzed include

variable proportions of the chlorite-sulfide and hematite-phengite-cassiterite assemblage. Only sample 34 (metasediment-hosted) represents Stage 1 assemblages and sample 24 represents Stage 2 assemblages (granite-hosted). The remaining samples have both Stage 1 and 2 assemblages. Chemical analyses illustrated in Figures 8, 9 and 10 result from the combined effects of the two assemblages; thus, subsequent discussion will try to emphasize the individual roles of each type of mineralization.

Within the granite-hosted lode samples, minor volume increases are expected due to hydration and have densities between 2.85 and 3.68 g/cm<sup>3</sup> in comparison to the granite (2.54 g/cm<sup>3</sup>). In addition, the most immobile elements (e.g., TiO<sub>2</sub>) are mobile (enriched slightly) even after accounting for density increases. Lentz *et al.* (1988) reported that TiO<sub>2</sub> varied from 0.07 to 0.11 wt. % which is within the 0.08 wt. % TiO<sub>2</sub> from the average True Hill granite (Table 4). This is similar to the TiO<sub>2</sub> variation in the 7 granite-hosted lodes (0.05 to 0.15 wt. %) hence a rigorous mass balance analysis is not warranted but comparative analysis with the unaltered granite should be examined with caution.

### Composition of Chlorite-Sulfide Assemblage (Stage 1)

The depletion of Si, K, Na, Ca, Rb and Sr and enrichment of Fe, Mg and Mn in the granite-hosted lode (Figs. 8, 9; Table 4) is due to the chloritization of feldspars during hydrothermal alteration. The minor depletion of other trace elements (Th, U, Nb, Ta, Zr and Y) is probably produced at this stage as well. The strong enrichment in S, Cu, Zn, Cd, and Bi and moderate enrichment of Pb, Ag and Mn relative to the True Hill granite (Fig. 9) is consistent with the presence of sulfide established by petrographic observations.

### Composition of Hematite-Phengite-Cassiterite Assemblage (Stage 2)

The hematite-phengite-cassiterite assemblage is formed by alkali and hydrogen metasomatism coincident with oxidation of iron based on petrographic observation. The variation in the amount of Fe<sub>2</sub>O<sub>3</sub>, MgO and MnO versus K<sub>2</sub>O and Na<sub>2</sub>O (Table 4) reflects the relative abundance of the two types of assemblages. The enrichment of alkali elements (Cs, Rb, Ba and K) in pre-existing depleted chloritized lode resulted in only moderate element depletions when the data are normalized to the average True Hill granite (Fig. 8).

### Rare-Earth-Element Compositions of the Lodes

Of the ten mineralized samples analysed, three are hosted in metasedimentary rocks and the remainder in granite (Table 4). Chondrite-normalized (CN) REE abundances (Wakita *et al.*, 1971) have a range of distributions (Fig. 10); (La/Lu)<sub>CN</sub> ranges from 0.7 to 8.2, of which 8.2, 7.0 and 6.4 correspond to the metasediment-hosted lodes. The granite-hosted lodes have a more restricted REE distribution in the highly fractionated True Hill granite (Lentz *et al.*, 1988). The consistently higher Eu<sub>CN</sub> content in the metasediment-hosted lode is indicated by values of Eu<sub>CN</sub> of 20.5, 27.4 and 35.6 whereas the granite-hosted lodes



Table 4. Chemical analyses of selected True Hill granite and lodes.

Sample wt. %	TH G	20 LG	24 LG	26 LG	30 LG	32 LG	36 LG	38 LG	22 LS	28 LS	34 LS
g/cm <sup>3</sup>	2.54	2.85	3.35	3.09	3.68	2.97	3.37	3.32	3.16	3.33	3.13
SiO <sub>2</sub>	76.6	30.1	32.9	49.1	20.2	48.4	20.6	32.5	42.3	22.5	37.5
TiO <sub>2</sub>	0.08	0.08	0.05	0.11	0.15	0.10	0.08	0.09	0.96	1.01	1.39
Al <sub>2</sub> O <sub>3</sub>	12.82	6.24	9.96	14.65	13.15	14.0	17.5	15.4	14.9	18.15	11.65
Fe <sub>2</sub> O <sub>3</sub>	1.05	44.65	48.9	27.35	59.35	25.7	50.45	41.8	32.6	45.75	22.2
MgO	0.09	0.13	0.21	0.18	0.35	0.705	0.69	0.57	0.84	1.77	1.56
MnO	0.05	0.87	0.65	0.55	0.36	0.61	2.51	0.25	0.64	1.56	0.48
CaO	0.61	0.19	0.06	0.12	0.09	0.33	0.05	0.01	0.19	1.38	0.84
Na <sub>2</sub> O	3.39	0.04	0.06	0.08	0.08	0.08	0.02	0.08	0.04	0.06	0.01
K <sub>2</sub> O	5.29	2.10	4.42	4.89	4.26	6.67	1.20	5.91	2.10	2.45	1.58
P <sub>2</sub> O <sub>5</sub>	0.01	0.02	0.01	0.02	0.03	0.03	0.03	0.03	0.02	0.03	0.03
LOI	-	15.2	2.00	2.77	1.81	2.59	6.89	2.85	15.2	4.39	7.08
Total	100.	99.6	99.2	100.1	99.8	99.2	100.1	99.5	99.5	99.0	84.25
ppm											
Cr	2	20	10	<10	10	<10	15	<10	20	90	130
Rb	534	300	440	760	645	700	150	845	300	270	185
Sr	26	<10	<10	<10	10	10	10	<10	<10	10	10
Y	83	140	<10	130	55	155	20	55	140	35	10
Zr	145	25	15	130	150	125	60	35	25	125	175
Nb	100	55	20	50	80	35	75	30	55	80	90
Ag	0.5	3.5	5.75	780	2.0	1.0	61	1.0	4.8	2.0	61
As	-	2700	5	41.5	6	36.5	110	6	2700	3.5	375
Auppb	-	80	<20	<20	<20	<20	<20	<20	80	<20	<20
Ba	86	400	<150	300	200	200	<150	200	400	<150	450
Bi	1.5	720	100	86.0	40.5	15.0	3.45	11.5	720	13.5	435
Cd	0.2	3.9	0.70	<0.2	0.7	0.7	<0.2	0.4	3.9	1.7	350.0
Co	44	53	24	19	11	12	17	10	53	24	85
Cr	2	<10	4	110	9	2	2	3.5	<10	110	155
Cs	6.1	20	22.5	37.5	29.5	32.0	7.9	38	20	27.5	22.0
Cu	8	130	40.5	25.5	405	10.0	74.0	8.25	130	210	2350
Hg	-	3.5	2	6	8.5	6	6	4	3.5	5	9.5
Mo	5	50	12	<5	13.5	<5	43	6	50	13	67
Ni	-	<1	<1	<1	<1	<1	<1	<1	<1	22.5	91.5
Pb	84	1300	155	70	610	22450	185	35	395	730	22450
Sb	0.25	30.0	3.15	6.05	3.2	2.0	1.25	2.2	30.0	1.3	13.5
Se	-	5	<3	<3	<3	<3	<3	<3	5	<3	56
Ta	10	3.5	2	6.5	9.5	3	8	4	3.5	2	1
Zn	61	1050	365	245	320	425	1600	200	1050	1550	59950
Sc	3.95	4.35	4.95	2.35	6.2	3.35	7.7	6.3	4.35	11.5	19.5
La	44.25	229	15.15	33.5	126	38.75	26.3	50.25	229	81	89.0
Ce	103.5	428	46	105.5	251	92.5	57	108	428	155	170
Nd	30	135	20	50	80	20	20	30	135	65	55
Sm	10.0	37.4	8.0	17.9	16.5	9.25	6.15	8.5	37.4	12.65	12.95
Eu	0.2	0.5	0.2	0.25	0.3	0.2	0.95	0.35	0.5	2.0	2.65
Yb	23	32	6	24.5	17	34	10	10	32	7.5	8.5
Lu	3.8	4.66	1.24	3.83	2.79	5.34	1.59	1.64	4.66	1.15	1.42
U	32	13.5	30.75	24.6	38.5	24.3	12.0	19.5	13.5	6.2	5.5
Th	69.6	24.0	6.75	47.0	74.5	77.0	41.0	28.0	24.0	47.0	25.5
B	-	10	<10	<10	<10	<10	<10	<10	10	<10	<10
Cl	-	<50	<50	<50	<50	<50	<50	<50	<50	<50	<50
F	-	1600	630	1700	1250	2750	860	1700	1600	1700	7050
In	-	8.0	10.0	8.0	17.5	14.5	11.0	12.5	8.0	8.0	24.0
Li	34.5	49.5	73	90.5	56.5	155	275	91	49.5	90.5	255
S%	-	16.6	0.05	0.485	0.06	0.575	1.58	0.485	16.6	0.485	5.02
Sn	6.8	6600	970	780	230	550	2300	4900	6600	780	1200

\*LG is granite-hosted lode, LS is metasediment-hosted lode. (X-ray Assay Laboratories Limited, 1885 Leslie Street, Don Mills, Ontario M3B 3J4). Majors elements, Cl and S (XRF), As, Au, Ba, Co, Cr, Cs, Hf, Mo, Sb, Se, Ta, Sc, La, Ce, Nd, Sm, Eu, Yb, Lu, U, Th by INAA, Ag, Cd, Cu, Ni, Pb, Zn, B, In by D. C. Plasma spectrometry, F and Li by wet chem, and Sn by EMS. All duplicate analyses were within several percent of each other so were averaged.



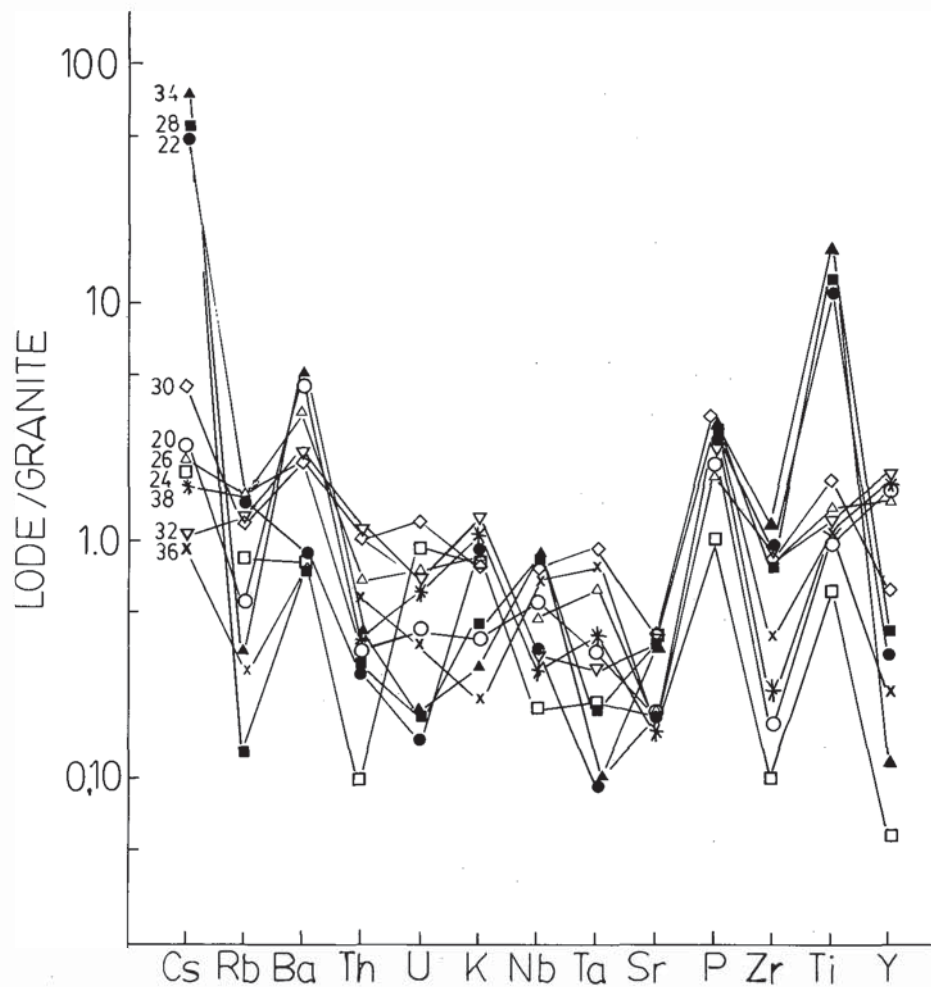


Fig. 8. Compositions of the lodes normalized to the "average" True Hill granite (Table 1). Elements are arranged according to their field strength.

range from 13.0 to 2.7  $\text{Eu}_{\text{CN}}$ . The high Eu content may, in part, be inherited from the metasediment during fluid interaction with metasedimentary rocks. This is substantiated by the strong positive correlation coefficients between Eu and Cr (0.96), Ti (0.96), and Sc (0.95). The Eu content of the granite-hosted lodes indicates an enrichment of 1 to 4.8 relative to the True Hill average granite (Fig. 10).

The absolute abundances of both LREE and HREE in the lodes are high; La and Lu concentrations being, respectively, 77 to 674 and 21 to 157. To facilitate comparison with the granite, the absolute abundances of REE in the lode samples also have been normalized to True Hill granite (Table 4; Fig. 10). It is important to understand that in most cases, excluding a few relict zircons and rutile, REE-bearing minerals have not been identified in association with the Stage 1 assemblage. Only three of the seven granite-hosted lode samples are higher in LREE (La, Ce, Nd and Sm) as compared to the granite, whereas four are weakly to moderately lower than the granite (Fig. 10). The granite-normalized, low Yb and Lu (HREE) distribution indicates that the HREE have not been enriched in the lodes to the extent that the LREE have been enriched. Relative to the granite HREE, Yb and Lu, range from 0.2 to 1.5 and 0.3 to 1.4, respectively, and 4

of the 7 granite-hosted samples give results below the granite. The lode samples are variably enriched in LREE relative to the HREE with mean values of granite-normalized La, and Lu of 1.7 and 0.6, respectively.

As previously mentioned, monazite commonly occurs with cassiterite in the hematite-phengite assemblage whereas xenotime occurs more rarely. This association explains the moderate positive correlation coefficients for Sn with La (0.65), Ce (0.57), Nd (0.53), and Sm (0.58) respectively (above 90% confidence level) whereas correlation coefficients for Sn with Yb (0.23), and Lu (0.18) are positive but low (Lentz, 1986).

#### Physiochemical Parameters

Stage 1 hydrolysis of K-feldspar to chlorite is estimated to occur at a pH between 5 (200°C) and 6 (100°C) at  $\log(K^*) = -1.5$  based on equilibria between chlorite and K-feldspar (Lentz, 1986). Similarly, Stage 2 is associated with the hydrolysis of K-feldspar to muscovite (phengite) which occurs at a pH of 5.1 (300°C at  $\log(K^*) = -1$ ; Montoya and Hemley, 1975) and a protolysis alteration of chlorite to phengite (equation 1).



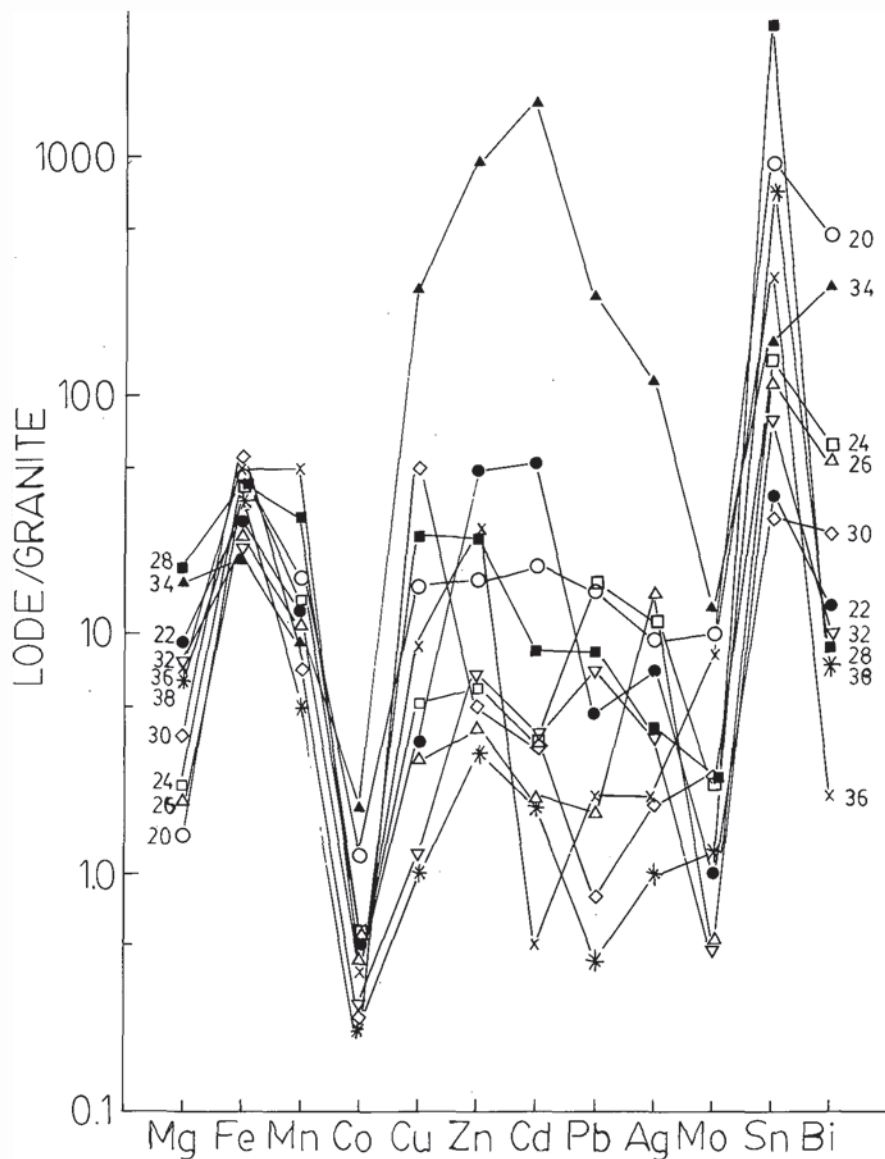


Fig. 9. Ore-element compositions of the lode relative to the "average" True Hill granite. Elements are in no particular order.

The coexistence of low-Fe sphalerite, pyrite and native silver in a Stage 1 assemblage indicates low temperatures (< 200°C; Lentz, 1986). The coexistence of low-Fe sphalerite (1 wt. % FeS) with pyrite indicates a low sulfur fugacity near the pyrite-pyrrhotite buffer under reducing conditions ( $H_2S$ ;  $\log(S_2) = -16$  to  $-25$  @ 100° to 200°C; Barton and Skinner, 1979). This coincides with the low temperature stability of septeclorite (<200°C) and their high ferrous iron content (XRD and Petruk, 1964). Stage 2 fluids were oxidizing, above the hematite-magnetite buffer ( $hm-mt \log f(O_2) = -39$  to  $-31$  @ 200° to 300°C). Septeclorite recrystallizes to chlorite during Stage 2 alteration. Phengite is transitional between the characteristics of 1M and 2M structures which indicates a depositional temperature between 200° and 300°C.

The deposition of fluorite and quartz during Stage 2 occurred either in response to a decrease in fluid temperature or a decrease in geostatic pressure during fluid ascent in the lode which is

compatible solubility studies (Holland and Malinin, 1979) and disposition of Stage 2 mineralization in fractures.

#### Sulfur Isotopes

Seven sulfur isotope analyses of pyrite (Stage 1) from various portions of the True Hill lodes have an average  $\delta^{34}S = 0.54$  per mil with a range of  $-2.8$  to  $3.6$  per mil (CDT; Table 5). Sulfides in the chlorite-sulfide lode were deposited under reducing conditions from an  $H_2S$  dominated fluid. At temperatures of approximately 200°C the  $\delta^{34}S$  of  $H_2S$  would be 2 to 3 per mil lighter than the pyrite (Ohmoto and Rye, 1979). The isotopic compositions and their narrow distribution between  $-3$  and  $0$  per mil  $\delta^{34}S$  favour chemically uniform conditions with a uniform  $^{34}S$  fluid composition. The deposition of sulfides (mainly pyrite) from a reduced  $H_2S$ -bearing Stage 1 fluid would decrease the sulfur content and the  $\delta^{34}S$  of the fluid.



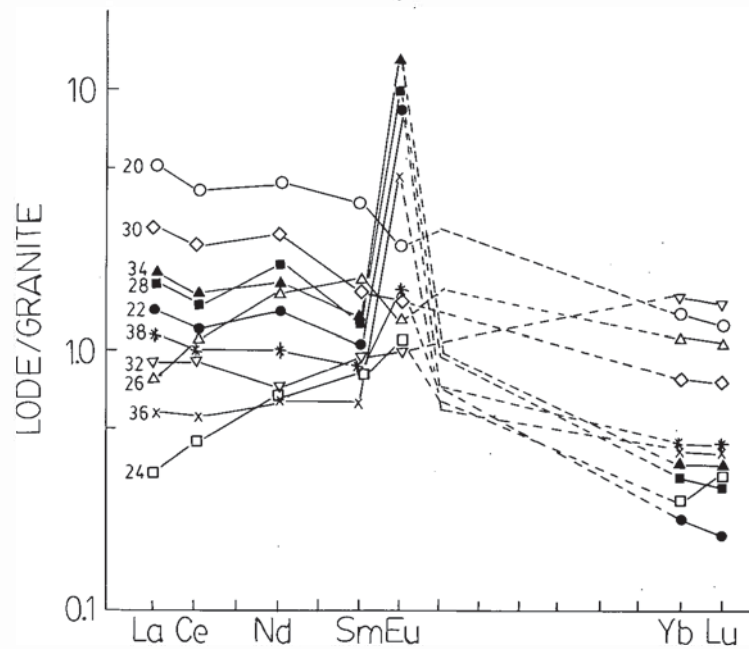
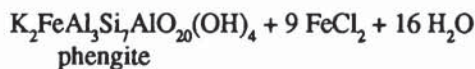
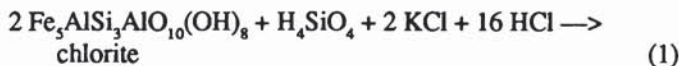


Fig. 10. True Hill granite-normalized REE distribution for lode samples (Table 1) (Gd values are extrapolated from a straight line between Sm and Yb). Solid symbols are metasediment-hosted lode (22, 28, 34) and the remainder are granite-hosted lodes.

### Deposition of Cassiterite

The deposition of LREE-bearing minerals with Sn under moderately acid pH's suggests similar mechanisms responsible for their deposition. The mobilization of LREE as chloride complexes (Flynn and Burnham, 1978; Alderton *et al.*, 1980; Taylor and Fryer, 1982; Humphris, 1984) indicates Sn-chloride rather than Sn-fluoride complexes are responsible. In order of relative importance based on petrographic observations, the deposition of Sn resulted from several processes; (1) decrease in effective chlorinity, (2) increase in pH due to feldspar alteration to K-mica and (3) decrease in temperature resulting from interaction with cool wall rocks. Merits will be discussed in turn.

(1) The most important reaction boundary in the lode is that of chlorite to phengite (equation 1) which involves removal of some Fe in solution as chloride complexes, with exchange, dehydration, and oxidation reactions. High concentrations of aqueous KCl and  $H_4SiO_4$  ( $SiF_4$ ) in the second hydrothermal fluid drives reaction 1 to the right.



The interaction of  $SnCl_3^-$  ( $SnCl_2$ ,  $SnCl^+$ ) with  $Fe^{2+}$  released during phengitization of chlorite may reduce the effective chlorinity of the hydrothermal fluid (equation 2). The shift in equilibrium of reaction 2, caused by Fe-chloride complexing, will decrease the solubility of Sn due to a decrease in Sn-chloride complexing. Oxygen balance considerations can be accommodated by the reduction of Fe.

Table 5. Sulfur isotopes.

(DDH-depth)	$\delta^{34}S$ (CDT)
80-03-283 ft	-1.6
80-03-314 ft	2.5
80-03-386 ft	0.3
82-5-55.3 m	0.4
82-06-293 m	0.4
82-13-238 m	-2.8
Trench 7b	3.6



$$\log K_1 = -3.69 = \log K_1 - \log K_2$$



$$\log K_1 = -8.2 \text{ Patterson } et al. (1981)$$



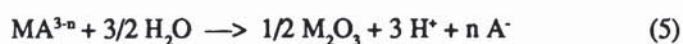
$$\log K_2 = -4.53 \text{ Crerar and Barnes (1976)}$$

The exact Sn-Cl ligation number is not known, but calculations by Patterson *et al.* (1981) indicate that  $SnCl^+$ ,  $SnCl_2$  and  $SnCl_3^-$  are the dominant chloride species in acidic chloride solutions acidic pH's (Jackson and Helgeson, 1985a; Eugster, 1985; Eadington, 1988), but neutral species are probably most stable at higher temperatures because of the decreasing dielectric constant



of water with higher temperatures. The deposition of LREE along the chlorite-sulfide/hematite-phengite reaction front resulted from, in part, a decrease in the chloride ligands available for complexing with LREE due to an increase in Fe-chloride complexing upon interaction with the chlorite-sulfide assemblage.

(2) Hydrolysis reactions may be responsible for the deposition of cassiterite along the granite-lode margin during the second mineralizing event. The increase in pH caused by hydrolysis results in the decreased solubility of Sn-chloride complexes in the fluid (Alderton *et al.*, 1980; Patterson *et al.*, 1981; Eugster, 1985; Jackson and Helgeson, 1985a, b; Eadington, 1988). Alderton *et al.* (1980) proposed a pH mechanism for the deposition of REE which may be applicable to this study (equation 5).



M = REE (or Sn<sup>2+</sup>), A = F, Cl, n = 0, 1, 2, 3

The deposition of REE along the boundary between the hematite-phengite assemblage and the granite may be in response to an increase in pH during feldspar hydrolysis.

(3) A decrease in temperature of Stage 2 fluids while interacting with (cooler) granite porphyry could be responsible for the deposition of the Sn with LREE and, to a lesser extent, with HREE. The Sn<sup>2+</sup>/Sn<sup>4+</sup> boundary is postulated to be 10 to 40 log f(O<sub>2</sub>) units above the SO<sub>4</sub><sup>2-</sup>/H<sub>2</sub>S boundary at 350°C (Patterson *et al.*, 1981) but moves to lower oxygen fugacities at lower temperatures resulting in the very low solubility of tin below 250°C (Jackson and Helgeson, 1985a).

## DISCUSSION

The Stage 1 chlorite-sulfide and Stage 2 hematite-phengite-fluorite-cassiterite assemblages are also present in Cornish tin-lode deposits such as the Wheal Jane (Rayment *et al.*, 1971), Mount Wellington tin deposit (Kettaneh and Badham, 1978), and the Levant tin-sulfide deposit (Jackson *et al.*, 1982). The mineral paragenesis and inferred zonation found at True Hill is similar but less complex than the tin lode deposits in Cornwall, England (Rayment *et al.*, 1971; Kettaneh and Badham, 1978; Hosking, 1979; Taylor, 1979; Jackson *et al.*, 1982; Strong, 1988) and in Nova Scotia (Pitre and Richardson, 1989).

Fluid inclusion and isotope studies from other tin deposits indicate moderate temperatures (250° to 350°C) and moderate salinities (10 wt. % salts) (Kelly and Turneaure, 1970; Kelly and Rye, 1979; Patterson *et al.*, 1981; Jackson *et al.*, 1982; Norman and Transcotchasan, 1982; Jackson and Helgeson, 1985b; Davis and Williams-Jones, 1985; Samson, 1990). Fluid inclusions in tin-sulfide lodes at Mount Pleasant are dominated by liquid-vapour inclusions with temperatures ranging from 100°C to 350°C and 0 to 10 wt. % eq. NaCl (Samson, 1990). This may be compatible with either a supercritical magmatic fluid (low salt) or a condensed fluid interacting with cooling meteoric water. Commonly, magmatic fluids have moderate salinities (15-17 wt. % eq. NaCl) upon fluid-phase separation (Burnham and Ohmoto,

1980). The development of Stage 1 and Stage 2 assemblages are postulated to be derived from a crystallizing magma at depth which intruded the True Hill granite porphyry. The lack of hydrothermal alteration at depth in the True Hill granite would obstruct meteoric fluid influx from the margins of the lode; however, this does not preclude meteoric or formation waters derived at a considerable depth below True Hill.

The enrichment of Fe, Mg, Mn, and other chalcophile elements and sulfur in the fluid responsible for stage 1 (chlorite-sulfide) may represent early magmatic degassing at depth. Strong partitioning of chalcophile elements into a supercritical fluid from a granite (Holland, 1972; Candela and Holland, 1984; Whitney *et al.*, 1985), as well as H<sub>2</sub>S and H<sub>2</sub> with H<sub>2</sub>O, could be responsible for the chlorite-sulfide assemblage. Alternatively, alteration of ferromagnesian phases in the True Hill granite at considerable depth could enrich the Stage 1 fluid in Fe, Mg and Mn. The Stage 1 fluids have characteristic moderately acidic pH's and low f(O<sub>2</sub>) and f(S<sub>2</sub>) similar to other tin-sulfide lodes. Stage 2 mineralization is characterized by alkali and hydrogen metasomatism which produced enrichment of trace elements (Sn, Th, U, Ta, Nb, and Y). These siliceous, fluorine and alkali fluids are characteristics of magmatic-hydrothermal fluids and may have been derived from a volatile-saturated magma. Reaction between the Stage 2, Sn- and REE-bearing fluid and the pre-existing chlorite-sulfide assemblage and granite are responsible for the deposition of cassiterite and monazite. The Stage 2 fluid is estimated to have been oxidized (>hm-mt buffer) not far from the Sn<sup>2+</sup>/Sn<sup>4+</sup> boundary. The low solubility of Sn in high f(O<sub>2</sub>) fluids may account for the subeconomic nature of this deposit.

The Stage 1 fluid apparently leached the REE-bearing phases from the granite, hence the overprinting by the Stage 2 fluid has introduced almost all of the REE except possibly some of the HREE tied up in trace zircon and rutile. The REE enrichment (Stage 2) may be due to partitioning into a hydrothermal fluid during second boiling of a granitic magma at depth (Flynn and Burnham, 1978) or leaching the granites at depth. The alteration of plagioclase (Fowler and Doig, 1983) may be responsible for the enrichment Eu in the hydrothermal fluid although the high Eu in metasediment-hosted lodes seems to be inherited. Incongruent alteration of LREE-bearing phases and plagioclase by weakly acidic, high-Cl fluids may be responsible for their enrichment in the Stage 2 fluids. Tin may be partitioned into a volatile phase during magmatic degassing similar to Zn, Mn, Fe and Cu (Holland, 1972) or Eu<sup>2+</sup> (Flynn and Burnham, 1978). However, leaching of Sn during alteration of biotite at considerable depth may not be precluded.

Although there are many similarities between True Hill lodes and lodes at Mount Pleasant, there are a few differences: (i) there is no documentation of two stages of mineralization within the Mount Pleasant lodes, (ii) the MP lodes host a single stage of cassiterite, stannite, pyrite, pyrrhotite, sphalerite and Cu sulfide mineralization similar to those in Cornwall contrast to that of True Hill and (iii) the oxidation state of the lodes in the North Zone (MP) is very low compared to the Stage 2 assemblages at True Hill. The cassiterite-bearing hematite-phengite assemblage is unique with respect to Sn lode mineralization world wide.

The change in f(O<sub>2</sub>) from Stage 1 to Stage 2 may be due to



either magmatic-oxidation resulting from degassing of  $H_2$  and  $H_2S$  fluids (Stage 1) from a granite emplaced at depth below True Hill or two unrelated fluids sources (different T, pH,  $f(O_2)$ , etc.).

## CONCLUSIONS

(1) Field observations indicate that lode mineralization at True Hill occurred along structures (fractures and faults) transecting the granite and metasedimentary rocks.

(2) Two distinct mineralizing episodes occurred within the lodes, represented by an early chlorite-sulfide assemblage (Stage 1) and a later hematite-phengite-cassiterite assemblage (Stage 2).

(3) The Stage 1 assemblage contains Mn- and  $Fe^{2+}$ -rich septeclorite (XRD), pyrite, sphalerite, galena, chalcocopyrite, Pb-Bi sulfides, and native Ag which formed at low temperatures ( $<200^\circ C$ ). Intense Fe, Mg, Mn and  $H_2S$  metasomatism, as well as the enrichment of chalcophile elements, implies high chloride concentrations. Chloritization (weakly acidic) within the fault-controlled lode was intense and leached alkali and alkali-earth elements from the host granite and metasedimentary rock. The coexistence of low-Fe sphalerite indicates low  $f(O_2)$  and low  $f(S_2)$  conditions for the Stage 1 fluid. The narrow variation of  $\delta^{34}S$  near 0 per mil for pyrite indicates uniform conditions during deposition from a fluid phase with a  $\delta^{34}S$  of between -3 and 0 per mil.

(4) Stage 2 (hematite-phengite-cassiterite) mineralization overprints the Stage 1 chlorite-sulfide assemblage. Interaction of weakly acidic, alkali-rich fluids with Stage 1 assemblages produced complex alteration. These fluids caused feldspar hydrolysis in the selvages adjacent to granite. Late quartz and fluorite veins and alteration are present, albeit uncommon. The high  $f(O_2)$  of the fluid led to oxidation and desulfidation of the pre-existing chlorite-sulfide (Stage 1) assemblage, as well as oxidation of biotite and ilmenite to produce rutile in the metasediment-hosted portions of the lodes. The association of LREE minerals with cassiterite in the vicinity of the reaction front of Stage 1 and Stage 2 assemblages and the granite lode contact indicates chloride complexing under weakly acidic conditions.

(5) In order of relative importance, deposition of cassiterite along a reaction front between Stage 2 assemblage and the chlorite-sulfide assemblage and the granite is probably related to (i) a decrease in effective chlorinity of the fluid, (ii) an increase in pH due to feldspar hydrolysis at the granite/lode contact and (iii) a decrease in temperature which may have been caused by interaction with the cooler granite porphyry.

(6) Interpretation of geologic observations, whole rock and mineral composition and sulfur isotopes supports magmatic-hydrothermal origin for the two stages of mineralization. The evidence favours partitioning of tin into a fluid phase during second boiling of a crystallizing granitic magma, as a stannous species, but the higher  $f(O_2)$  of the fluid inhibited more significant enrichment due to the lower solubility of Sn under these conditions. The spatial relationship and origin of the chlorite-sulfide and hematite-phengite assemblages is much more problematical, but a two-stage or progressive devolatilization of the underlying magma may explain both the change in redox conditions and chemical compositions of the fluids which formed the lode. Alternately, the two assemblages may reflect different fluid sources.

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ALDERTON, D.H.M., PEARCE, J.A., and POTTS, P.J. 1980. Rare-earth element mobility during granite alteration: evidence from southwest England. *Earth and Planetary Science Letters*, 49, pp. 149-165.

BARTON, P.B. and SKINNER, B.J. 1979. Sulfide Mineral Stabilities. *In Geochemistry of Hydrothermal Ore Deposits*. Second edition. Edited by H.L. Barnes. John Wiley and Sons Limited, New York, New York, pp. 71-136.

BURNHAM, C.W. and OHMOTO, H. 1980. Late stage processes in felsic magmatism. *In Granitic magmatism and related mineralization*. Edited by S. Ishihara and S. Takenouchi. Mining Geology, Special Issue No. 8, Society of Mining Geologists, Japan, pp. 1-11.

CANDELA, P. and HOLLAND, H.D. 1984. The partitioning of copper and molybdenum between silicate melts and aqueous fluids. *Geochimica et Cosmochimica Acta*, 48, pp. 373-380.

CARROLL, D. 1970. Clay Minerals: A Guide to their X-Ray Identification. Geological Society of America, Special Paper 126, 80 p.

CRERAR, D.A. and BARNES, H.L. 1976. Ore solution chemistry v. solubilities of chalcocopyrite and chalcocite assemblages in hydrothermal solution at  $200^\circ$  to  $350^\circ C$ . *Economic Geology*, 71, pp. 772-794.

DAGGER, G.W. 1972. Genesis of the Mount Pleasant tungsten-molybdenum deposit, New Brunswick. Institute of Mining and Metallurgy, Transactions Section B, 81, pp. 73-102.

DAVIS, W.J. and WILLIAMS-JONES, A.E. 1985. A fluid inclusion study of porphyry-greisen, tungsten-molybdenum deposit at Mount Pleasant, New Brunswick, Canada. *Mineralium Deposita*, 49, pp. 7-23.

EADINGTON, P.J. 1988. The solubility of cassiterite in hydrothermal solutions in relation to some lithological and mineral associations in tin ores. *In Recent Advances in the Geology of Granite-Related Mineral Deposits*. Edited by R.P. Taylor and D.F. Strong. CIM Special Volume 39, pp. 25-32.

EUGSTER, H.P. 1985. Granites and hydrothermal ore deposits: A geochemical framework. *Mineralogical Magazine*, 49, pp. 7-23.

FLYNN, R.T. and BURNHAM, C.W. 1978. An experimental determination of Rare Earth Partition Coefficients between a chloride containing Vapour Phase and Silicate Melts. *Geochimica et Cosmochimica Acta*, 42, pp. 685-701.

FOWLER, A.D. and DOIG, R. 1983. The significance of europium anomalies in the REE spectra of granites and pegmatites, Mont Laurier, Quebec. *Geochimica et Cosmochimica Acta*, 47, pp. 1131-1137.

HEY, M.H. 1954. A new review of the chlorites. *Mineralogical Magazine*, 30, pp. 277-292.

HOLLAND, H.D. 1972. Granites, Solutions, and Base Metal Deposits. *Economic Geology*, 67, pp. 281-301.

HOLLAND, H.D. and MALININ, S.D. 1979. The Solubility and Occurrence of Non-Ore Minerals. *In Geochemistry of Hydrothermal Ore Deposits*. Edited by H.L. Barnes. John Wiley and Sons



- Limited, New York, New York, pp. 461-508.
- HOSKING, K.F.G. 1979. Tin Distribution Patterns. *Geological Society of Malaysia Bulletin*, 11, pp. 1-70.
- HUMPHRIS, S.E. 1984. The Mobility of the Rare Earth Elements in the Crust. In *Rare Earth Element Geochemistry*. Edited by P. Henderson. Elsevier, New York, New York, pp. 317-342.
- JACKSON, K.J. and HELGESON, H.C. 1985a. Chemical and thermodynamic constraints on the hydrothermal transport and deposition of tin: I. Calculation of the solubility of cassiterite at high pressures and temperatures. *Geochimica et Cosmochimica Acta*, 49, pp. 1-22.
- . 1985b. Chemical and thermodynamic constraints on the hydrothermal transport and deposition of tin: II. Interpretation of phase relations in the Southeast Asian Tin Belt. *Economic Geology*, 80, pp. 1365-1378.
- JACKSON, N.J., HALLIDAY, A.N., SHEPPARD, S.M.F., and MITCHELL, J.G. 1982. Hydrothermal activity in the St. Just Mining District, Cornwall, England. In *Metallization Associated with Acid Magmatism*. Edited by A.M. Evans. John Wiley and Sons Limited, pp. 137-179.
- KELLY, W.C. and RYE, R.O. 1979. Geologic, Fluid Inclusion, and Stable Isotope Studies of the Tin-Tungsten Deposits of Panasqueira, Portugal. *Economic Geology*, 74, pp. 1721-1822.
- KELLY, W.C. and TURNEAURE, F.S. 1970. Mineralogy, paragenesis and geothermometry of the tin and tungsten deposits of the Eastern Andes, Bolivia. *Economic Geology*, 65, pp. 609-680.
- KETTANEH, Y.T. and BADHAM, J.P.N. 1978. Mineralization and Paragenesis at the Mount Wellington Mine, Cornwall. *Economic Geology*, 73, pp. 486-495.
- KOOIMAN, G.J., McLEOD, M.J., and SINCLAIR, W.D. 1986. Porphyry Tungsten-Molybdenum Orebodies, Polymetallic Veins and Replacement Bodies, and Tin-bearing Greisen Zones in the Fire Tower Zone, Mount Pleasant, New Brunswick. *Economic Geology*, 81, pp. 1356-1373.
- LENTZ, D.R. 1986. Geology and Depositional Conditions of Tin lodes at True Hill, New Brunswick. Unpublished M.Sc. thesis, University of New Brunswick, 280 p.
- LENTZ, D.R., LUTES, G.G., and HARTREE, R. 1988. Bi-Sn-W-Mo greisen mineralization associated with the True Hill granite, southwestern N.B. *Maritime Sediments and Atlantic Geology*, 24, pp. 321-338.
- MONTOYA, J.W. and HEMLEY, J.J. 1975. Activity relations and stabilities in alkali feldspars and mica alteration reactions. *Economic Geology*, 70, pp. 577-583.
- MULLIGAN, R. 1975. Geology of Canadian tin occurrences. Geological Survey of Canada, Economic Geology Report No. 28, 155 p.
- NORMAN, D.I. and TRANSCOTCHASAN, Y. 1982. Mineralization and Fluid Inclusion Study of the Yod Nam Tin Mine, Southern Thailand. In *Mineralization Associated with Acid Magmatism*. Edited by A.M. Evans. John Wiley and Sons Limited, New York, New York, pp. 261-272.
- OHMOTO, H. and RYE, R.O. 1979. Isotopes of Sulfur and Carbon. In *Geochemistry of Hydrothermal Ore Deposits*. Edited by H.L. Barnes. John Wiley and Sons Limited, New York, New York, pp. 509-567.
- PARRISH, I.S. 1977. Mineral catalog for the Mount Pleasant Deposit of New Brunswick Tin Mines. *Canadian Mineralogist*, 15, pp. 121-126.
- PARRISH, I.S. and TULLY, J.V. 1978. Porphyry tungsten zones at Mt. Pleasant, N.B. *Canadian Institute of Mining and Metallurgy Bulletin*, 71, pp. 93-100.
- PATTERSON, D.J., OHMOTO, H., and SOLOMON, M. 1981. Geologic setting and genesis of cassiterite-sulfide mineralization at Renison Bell, western Tasmania. *Economic Geology*, 76, pp. 393-438.
- PETRUK, W. 1964. Mineralogy of the Mount Pleasant Tin Deposit in New Brunswick. *Canada Mines Branch Technical Bulletin*, 56, 37 p.
- PETRUK, W. 1973. The tungsten-molybdenum-bismuth deposit of Brunswick Tin Mines Limited: Its mode of occurrence, mineralogy, and amenability to mineral beneficiation. *Canadian Institute of Mining and Metallurgy Bulletin*, 66, pp. 113-130.
- PITRE, C.V. and RICHARDSON, J.M. 1989. Paragenesis of veins of the Duck Pond tin prospect, Meguma Group, East Kempville, Nova Scotia. *Canadian Journal of Earth Sciences*, 26, pp. 2032-2043.
- POULIOT, G., BARONDEAU, B., SAUVE, P., and DAVIS, M. 1978. Distribution of alteration minerals and metals in the Fire Tower zone at Brunswick Tin Mines Limited, Mount Pleasant area, New Brunswick. *Canadian Mineralogist*, 16, pp. 223-237.
- RAYMENT, B.D., DAVIS, G.R., and WILSON, J.D. 1971. Controls to mineralization at Wheal Jane, Cornwall. *Transactions of the Institute of Mining and Metallurgy, Section B*, 80, pp. B224-B237.
- RUITENBERG, A.A. 1963. Tin Mineralization and Associated Rock Alterations at Mount Pleasant, Charlotte County, New Brunswick. Unpublished M.Sc. thesis, University of New Brunswick, 172 p.
- RUITENBERG, A.A. and FYFFE, L.R. 1982. Mineral Deposits associated with granitoid intrusions and related subvolcanic stocks in New Brunswick and their relationship to Appalachian tectonic evolution. *Canadian Institute of Mining and Metallurgy Bulletin*, 75, pp. 83-97.
- RUITENBERG, A.A. and McCUTCHEON, S.R. 1985. Tungsten, Molybdenum, and Tin Deposits in New Brunswick. Geological Association of Canada/Mineralogical Association of Canada, Field Trip Guidebook, Excursion 13, 33 p.
- SAMSON, I.M. 1990. Fluid evolution and mineralization in a subvolcanic granite stock: The Mount Pleasant W-Mo-Sn deposits, New Brunswick. Geological Association of Canada/Mineralogical Association of Canada, Program with Abstracts, 15.
- SINCLAIR, W.D., KOOIMAN, G.J.A., and MARTIN, D.A. 1988. Geological setting of granites and related tin deposits in the North Zone, Mount Pleasant, New Brunswick. In *Current Research, Part B*, Geological Survey of Canada, Paper 88-1B, pp. 201-208.
- STRONG, D.F. 1988. A review and model for granite-related mineral deposits. In *Recent Advance in the Geology of Granite-Related Mineral Deposits*. Edited by R.P. Taylor and D.F. Strong. CIM Special Volume 39, pp. 424-445.
- TAYLOR, R.G. 1979. *Geology of Tin Deposits*. Elsevier, New York, New York, 217 p.
- TAYLOR, R.P. and FRYER, B.J. 1982. Rare-earth Element Geochemistry as an Aid to Interpreting Hydrothermal Ore Deposits. In *Metallization Associated with Acid Magmatism*. Edited by A.M. Evans. John Wiley and Sons Limited, New York, New York, pp. 357-365.
- TAYLOR, R.P., SINCLAIR, W.D., and LUTES, G.G. 1985. Geochemical and Isotopic Characterization of Granites Related to W-Sn-Mo Mineralization in the Mount Pleasant Area, New Brunswick. In *Granite Related Mineral Deposits*. Edited by R.P. Taylor and D.F. Strong. Canadian Institute of Mining and Metallurgy, pp. 265-273.
- WAKITA, H., REY, P., and SCHMITT, R.A. 1971. Abundances of the 14 rare-earth elements and 12 other trace elements in Apollo 12 samples: five igneous and one breccia rocks and four soils. *Proceedings of the second Lunar Science Conference*, pp. 1319-1329.
- WHITNEY, J.H., HEMLEY, J.J., and SIMON, F.O. 1985. The Concentration of Iron in Chloride Solutions Equilibrated with Synthetic Granite Compositions: The Sulfur-Free System. *Economic Geology*, 80, pp. 444-460.



## APPENDIX : Sample locations

## 1a. Chemical Analyses

<u>SAMPLE #</u>	<u>DDH#-DEPTH</u>	<u>STAGE</u>	<u>HOST ROCK</u>
20	82-06-285.6 m	Stage 1 and 2	(granite-hosted)
22	82-13-050.1 m	Stage 1 and 2	(metasediment-hosted)
24	82-06-288.5 m	Stage 2	(granite-hosted)
26	82-06-317.4 m	Stage 1 and 2	(granite-hosted)
28	82-15-062.3 m	Stage 1 and 2	(metasediment-hosted)
30	82-24-165.9 m	Stage 1 and 2	(granite-hosted)
32	83-01-284.9 m	Stage 1 and 2	(granite-hosted)
34	82-15-038.1 m	Stage 1	(metasediment-hosted)
36	Trench 7b	Stage 1 and 2	(granite-hosted)
38	82-27-214.6 m	Stage 1 and 2	(granite-hosted)