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# Precambrian Evolution of the Avalon Terrane in the Northern Appalachians: A Review

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### Résumé de l'article

Dans les Appalaches septentrionales. la laniere d' Avalon forme une zone tectonostratigraphique distincte^ definie prIncipalement par la presence de rocnes volcano-sedimentaires et granitoldes tardi-precambriennes (circa 600 Ma) recouvertes par des series paleozoiques inferieures qui renferment une faune acado-baltique. La laniere d'Avalon affleure dans Vest de Terre-Neuve, dans l'ile du Cap-Breton, dans^le nord de la penInsule de Nouvelle-Ecosse, au Nouveau-Brunswick meridional, sur la cote sud-est du Maine et dans le sud-est de la Nouvelle-Angleterre.

Deux episodes tectonothermiques d'envergure ont affecte les socles continental et oceanique et ainsi conditionne l'evolution precambrienne de la laniere d'Avalon. Un magmatisme oceanique local et le metamorphisme regional d'une plate-forme a terrigenea et a carbonates du Proterozoique moyen (tardl-Helikien7), ainsi que de son socle

gneissique continental, pourraient tradulxe une phase de rifting ou de subduction et d'effondrement local de celle-cl a l'eo-Hadrynien (750-800 Ha). Un plutonisme granitoide calco-alcallne et un volcanisme largement repandu a caracteres tholeitique. calco-alcalln et, occasionellement. peralcalln tardHiadryiiiens (580-630 Ha) concretisent un episode de serrage qui semble englober des regimes d'arc enslallque. d'Inter arc oceanique et des conteites d,arrlere arc continental tant en extension qu'en transtenslon. Le serrage se conclut, peut-etre par des Jeux transformants, lore de 1'orogenle avalonlenne au ardi-Hadrynien et s'lliustra localenent par le developpement de flyschs qui portent la trace d'une glaclation vendlenne; apres quoi s'implanterent des >^p\*Hnfl successeurs de type nolasalque. Au flni-Hadrynien, des gres rouz volcanogeniques et un volcanisme blmodal, accompagnes d' une transtenslon largement repandue dans 1' arriere arc, peuvent annoncer la nalssance de 1'ocean Iapetus et le retour d'une plate-forme au tout debut du Paleozoique.

L'amorce de conditions de plate-formes du domaine atlantique qui s'empilerent au Cambro-Ordovicien, s'accompagna d'un faible volcanisme de type rift intra plaque. Des volcanites bimodales, des sediments marina d'eau peu profonde et de rares plutons peralcalins s'y grefferent par endroita au tardi-Ordovicien ou a l'eo-Devonien. On note localement les effets de 1'episode tectonothermique medio-ordovicien ("taconien") bien qu'ils se limitent pour une grande part a des discordances majeures. Le metajnorpbisme et la deformation acadiens tres repandus, d'age eo- a medio-devonien, pourraient representer l'accolement tectonique des lanieres d'Avalon et de Gander; lis a'accompagnerent par 1'emplacement dans ces deux lanieres de plutons granito'ides acadiens localement volundneux. Par contre, des portions de la laniere d'Avalon pourraient s'etre accretees plus tot. Au Devono-Carbonifere, la deposition extensive d'assises de type molassique associees A des volcanites blmodales tholeitiques a alcalines faconne une sequence de couverture sur le reste de l'orogene appalacbien bien que la\*position actuelle de la laniere d'Avalon soit infeodee localement A de grands accidents alleghaniens. Des lita roux et des volcanites mafiques non-deformees du Trias annoncent l'ouverture de 1\*ocean Atlantique actuel.

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# Precambrian Evolution of the Avaion Terrane in the Northern Appalachians: A Review

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The Avalon terrane of the Northern Appalachians forms a distinctive tectonostratigraphic belt defined primarily by the presence of late Precambrian (circa 600 Ma) volcano-sedimentary and granitoid rocks overlain by Lower Paleozoic sequences containing Acado-Baltic fauna. Avalon terrane is exposed in eastern Newfoundland, Cape Breton Island and the northern Nova Scotian mainland, southern New Brunswick, coastal southeastern Maine, and southeastern New England.

The Precambrian evolution of Avalon terrane includes two major tectonothermal events that involved both continental and oceanic basement. Local oceanic mafic magmatism and regional metamorphism of a mid-Proterozoic (late Helikian?) carbonate-clastic platform and its gneissic continental basement may record an early Hadrynian (750-800 Ma) rifting or subduction event and local platform collapse. Late Hadrynian (580-630 Ma) calc-alkaline granitoid plutonism and widespread volcanism of tholeitic, calcalkaline and occasionally peralkaline affinities record a closure event that is interpreted to involve ensialic arc, oceanic interarc, and both extensional and transtensional continental back-arc settings. Closure terminated, perhaps through transform interactions, in the late Hadrynian Avalonian orogeny and was locally heralded by molasse-like successor basins. Latest Hadrynian volcanogenic redbeds and bimodal volcanism, associated with widespread back-arc transtension, may herald inception of the Iapetus ocean and platform re-establishment during the earliest Paleozoic.

Onset of overlapping, Atlantic-realm platform conditions in the Cambro- Ordovician was accompanied by minor, within-plate rift volcanism and is in places succeeded by late Ordovian-early Devonian bimodal volcanic rocks, shallow marine sediments and rare peralkaline plutons. Effects of a mid-Ordovician ("Taconian") tectonothermal event occur locally but are generally limited to major unconformities. Widespread early to mid-Devonian Acadian metamorphism and deformation may record the tectonic juxtapositioning of Avalon and Gander terranes and was accompanied by the emplacement into both terranes of locally voluminous Acadian granitoid plutons. However, portions of the Avalon terrane may have accreted earlier. Deposition of widespread Devono-Carboniferous, molasse-like sequences associated with bimodal tholeiitic to alkaline volcanics form an overlap sequence with the remainder of the Appalachian orogen, although the present position of the Avalon terrane with respect to the orogen is locally controlled by major Alleghanian faulting. Undeformed Triassic redbeds and mafic volcanic rocks herald initial rifting of the present Atlantic Ocean.

Dans les Appalaches septentrionales, la lanière d'Avalon forme une zone tectonostratigraphique distincte définie principalement par la présence de roches volcano-sédimentaires et granitoides tardi-précambriennes (circa 600 Ma) recouvertes par des séries paléozoïques inférieures qui renferment une faune acado-baltique. La lanière d'Avalon affleure dans l'est de Terre-Neuve, dans l'île du Cap-Breton, dans le nord de la péninsule de Nouvelle-Ecosse, au Nouveau-Brunswick méridional, sur la côte sud-est du Maine et dans le sud-est de la Nouvelle-Angleterre.

Deux épisodes tectonothermiques d'envergure ont affecté les socles continental et océanique et ainsi conditionné l'évolution précambrienne de la lanière d'Avalon. Un magmatisme océanique local et le métamorphisme régional d'une plate-forme à terrigènes et à carbonates du Protérozoique moyen (tardi-Hélikien?), ainsi que de son socle

gneissique continental, pourraient traduire une phase de rifting ou de subduction et d'effondrement local de celle-ci à l'éo-Hadrynien (750-800 Ma). Un plutonisme granitoide calco-alcaline et un volcanisme largement répandu à caractères tholéitique, calco-alcalin et, occasionellement, peralcalin tardi-hadryniens (580-630 Ma) concrétisent un épisode de serrage qui semble englober des régimes d'arc ensialique, d'inter arc océanique et des contextes d'arrière arc continental tant en extension qu'en transtension. Le serrage se conclut, peut-être par des jeux transformants, lors de l'orogènie avalonienne au tardi-Hadrynien et s'illustra localement par le développement de flyschs qui portent la trace d'une glaciation vendienne; après quoi s'implantèrent des bassins successeurs de type molassique. Au fini-Hadrynien, des grès roux volcanogèniques et un volcanisme bimodal, accompagnès d'une transtension largement répandue dans l'arrière arc, peuvent annoncer la naissance de l'océan Iapetus et le retour d'une plate-forme au tout début du Paléozoïque.

L'amorce de conditions de plate-formes du domaine atlantique qui s'empilèrent au Cambro-Ordovicien, s'accompagna d'un faible volcanisme de type rift intra plaque. Des volcanites bimodales, des sédiments marins d'eau peu profonde et de rares plutons peralcalins s'y greffèrent par endroits au tardi-Ordovicien ou à 1'éo-Dévonien. note localement les effets de l'épisode tectonothermique médio-ordovicien ("taconien") bien qu'ils se limitent pour une grande part à des discordances majeures. Le métamorphisme et la déformation acadiens très répandus, d'âge éo- à médio-dévonien, pourraient représenter l'accolement tectonique des lanières d'Avalon et de Gander: ils s'accompagnèrent par l'emplacement dans ces deux lanières de plutons granitoïdes acadiens localement volumineux. Par contre, des portions de la lanière d'Avalon pourraient s'être accrétées plus tôt. Au Dévono-Carbonifère, la déposition extensive d'assises de type molassique associées à des volcanites bimodales tholéiltiques à alcalines façonne une séquence de couverture sur le reste de l'orogène appalachien bien que la position actuelle de la lanière d'Avalon soit inféodée localement à de grands accidents alleghaniens. Des lits roux et des volcanites mafiques non-déformées du Trias annoncent l'ouverture de l'océan Atlantique actuel. [Traduit par le journal]

# INTRODUCTION

In Atlantic Canada and New England, the southeastern margin of much of the Northern Appalachian orogen forms part of a distinctive tectonostratigraphic belt traditionally termed the "Avalon Platform" (Kay and Colbert 1965; Williams 1969) or the "Avalon Zone" (Rodgers 1968, 1972; "Zone H" of Williams et al. 1972, 1974). Although generally less deformed than belts within the Appalachian orogen. the Avalon zone is distinguished principally Ъ<del>у</del> the presence of 1ate Proterozoic (Hadrynian) volcanosedimentary sequences and co-genetic granitoid rocks of the circa 600 Ma Avalonian (Cadomian II) orogeny that overlain by Lower Paleozoic. shallow-marine platformal successions bearing Acado-Baltic (Atlantic-realm) trilobite fauna (Skehan et al. 1978; O'Brien et al. 1983; Rast and Skehan 1983).

As defined by Williams (1978, 1979), the Avalon zone forms one of five tectonostratigraphic zones within

the Northern Appalachian orogen (Fig. distinguished on the basis of contrasts in their pre-Silurian evolution and interpreted in terms of the creation and destruction of Precambrian - early Paleozoic Iapetus Thus the Humber zone records the development and destruction of the western miogeocline of Iapetus, whereas the Dunnage zone preserves vestiges of Iapetus and its island arc. remnant arc. and back-arc sequences. Gander zone has been interpreted to contain elements of the eastern active margin of Iapetus (Kennedy 1976: Williams 1978, 1979; Coleman-Sadd 1980) but is defined largely on its structural and metamorphic style, closely resembles the Dunnage zone in age and lithology, and lacks proven Grenvillian basement, rift-related magmatism, and a passive margin history (Williams and Hatcher 1983). However, relationships of the more easterly Gander, Avalon and Meguma zones to the Iapetus cycle are uncertain as the zones are structurally uncoupled and show no demonstrable connection to the ancient North

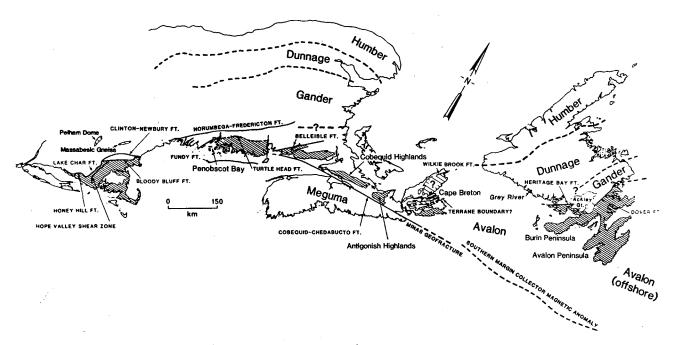


Fig. 1. Distribution of Precambrian Avalon terrane (broad-shaded where uncertain) and major terrane boundary faults in the Northern Appalachians (modified after Zen 1983). Zone boundaries from O'Brien et al. (1983) and Williams and Hatcher (1983); Collector Anomaly from Haworth and Lefort (1979); and Fundy Fault from Ballard and Uchupi (1975).

American miogeocline. More recent treatments of the Appalachian orogen have therefore viewed these, and other zones in the southern Appalachians and Burope, as a mosaic of suspect terranes (Keppie 1981; Williams and Hatcher 1982, 1983; Zen 1983; Keppie et al. 1985) that lie west of the eastern miogeoclinal margin of Iapetus preserved in the Scandinavian Caledonides, the Mauritanides of west Africa, and possibly Morocco (Williams 1984).

Within the Avalon terrane, wide variations in Proterozoic depositional settings and volcanic affinities suggest a composite terrane history during the Precambrian, the assembly of which may correspond to the late Precambrian Avalonian orogeny and must have occurred prior to the development of the Cambro-Ordovician overlap widespread sequence bearing distinctive Atlanticrealm fauna (Williams and Hatcher 1982, 1983). Cambro-Ordovician faunal contrasts (Cocks and Fortey 1982; Neuman 1984) and recently determined paleomagnetic discrepancies (Johnson and Van der Voo 1985; Van der Voo and Johnson 1985) suggest considerable separation of Avalon terrane from cratonic North America and place the Avalon terrane at

high paleolatitudes in the Lower Paleozoic. However, paleomagnetic evidence for the position of the Avalon terrane with respect to North America remains inconclusive for much of the Paleozoic (Kent and Opdyke 1978, 1979; Spariosu and Kent 1983; Scotese et al. 1984) and the mechanism and timing of the accretion of Avalon terrane to the remainder of the Appalachian orogen remains debated. Similar Devonian plutonic histories and local sedimentary linkages in Newfoundland suggest docking of the Avalon and adjacent Gander during the mid-Paleozoic terranes (Williams and Hatcher 1982, 1983) although portions of the Avalon terrane may have accreted earlier (Robinson and Hall 1980; Keppie et al. 1985; Currie et al. 1986: Ludman 1986). The absence of linkages between the Avalon and Meguma terranes across the Minas Geofracture (Fig. 1), other than a Carboniferous overlap sequence, suggests docking of the Meguma terrane during mid-to-late Paleozoic (Keppie 1982a; Williams and Hatcher 1982. 1983: Scotese et al. 1984; Spariosu et al. 1984). Kinematic and Ar/Ar geochronological analysis of ductile shear zones associated with this boundary (Keppie

et al. 1985) suggests accretion of the Meguma zone involved dextral displacements during the mid-to-late Devonian (360-373±7 Ma, Dallmeyer and Keppie 1984).

### DISTRIBUTION OF THE AVALON TERRANE

The Avalon terrane, whose type area the Avalon Platform 4 8 1 Newfoundland (e.g. King 1980), is the largest Appalachian suspect terrane extending more or less continuously from offshore northeastern Newfoundland to southern New Brunswick. and discontinuously to southern Maine, southeast-New England and possibly Carolinas (Williams 1978: Williams and Hatcher 1982, 1983). It may also form the basement to Florida (Smith 1982). However, the tectonostratigraphy of the Carolina terrane (Secor et al. 1983) may be sufficiently distinctive to warrant discontinuing the use of the term Avalon terrane in the southern Appalachians. Possible correlatives of the Avalon terrane occur outside the Appalachian orogen in the Caledonides of southern Britain (Rast et 1976), the Cadomian massifs of western and central Europe (Rast 1980), and the Pan-African belts of northwest Africa (Pique 1981; O'Brien et al. 1983; Rast and Skehan 1983).

Within the Canadian Appalachians 1), O'Brien et al. (1983) de-(Fig. fined the Avalon terrane as occupying Newfoundland east of the Hermitage Bay and Dover faults (Blackwood and Kennedy 1975: Dallmeyer et al. 1981a), at least southeastern Cape Breton (Barr and Raeside 1986) and the Antigonish and Cobequid Highlands of Nova Scotia north of the Minas Geofracture (Keppie 1982a), and New Brunswick south of the Wheaton Brook and Belleisle faults (McCutcheon 1981; Currie 1984, 1986a). Avalon terrane is continuous westward in the offshore to the edge of the continental shelf (King et al. 1985) and is bound to the southeast by the Collector magnetic anomaly (Haworth and Lefort 1979; Williams and Hatcher 1982, 1983). Deep seismic reflection results have recently shown the Dover fault to be a major crustal suture separating lower crust of markedly different seismic character (Keen et al. 1986).

In southern Maine. the Avalon terrane occupies the Coastal Volcanic Belt south of the Turtle Head fault (Stewart and Wones 1974: Keppie et al. 1985) and possibly north of the offshore Fundy fault (Ballard and Uchupi 1975) which may separate the Avalon and Meguma terranes beneath the Triassic Fundy Basin. Avalon terrane also forms much of southeastern New England where it is bound to the north and west by the Bloody Bluff, Lake Char, and Honey Hill faults (Skehan and Murray 1980; Skehan 1983: Zen 1983: Zartman and Naylor 1984).

However, due largely to its composite Precambrian nature, considerable discrepancy exists in the usage of the term Avalon terrane and further subdivision is undoubtedly necessary. O'Hara and Gromet (1985) have recently identified two Avalonian terranes (Esmond-Dedham and Hope Valley) in southeastern New England that are separated by the Hope Valley shear zone and are distinguished by contrasting Paleozoic accretionary histories and late Precambrian plutonic rocks (e.g. Hermes and Zartman 1985). However, as Hope Valley shear zone is an Alleghanian feature, the distinction between these terranes may reflect post-accretionary, Alleghanian reorganization. Gaudette and Olszewski (1986) have similarly suggested, on the basis of distinctive thermo-tectonic histories, that "Avalon" be restricted to the Avalon platforms of Newfoundland and Boston whereas inboard terranes. which show some similarities "Avalon" and are bound to the north and west by the Lake Char - Campbell River-Nonesuch River-Norumbega fault system, should be grouped separately as the Hope Valley-Cape Breton terrane. ever, such a subdivision ignores the distinction of Avalon terrane based on lowermost Paleozoic faunal criteria and presence of 1ate Precambrian volcano-plutonic suites.

Keppie et al. (1985) applied the "Avalon composite terrane" to the product of late Precambrian (Avalonian) terrane amalgamation and suggest that, prior to this time. "Avalon" comprised numerous terranes that probably evolved Thus the Avalon composite separately. terrane includes the Avalon Peninsula terrane of Newfoundland. the Cape Breton and Cheticamp terranes of Cape Breton, the Antigonish and Cobequid terranes of mainland Nova Scotia, the Caledonia Highlands and Kingston terranes of southern New Brunswick, the terrane (Coastal Volcanic Islesboro belt) of southern Maine, and the Cape Cod and Boston terranes of southeastern New England. However, given present uncertainties in the detailed geology of these areas and their degree of Precambrian separation, such extensive subdivision may be premature. Northern portions of the Cape Breton terrane, for example. may contain correlatives of the Gander and Humber zones and Raeside 1986: Raeside and Barr 1986), whereas the Cradle Brook fault of southern New Brunswick may tectonothe stratigraphically subdivide Caledonia Highlands terrane (McLeod 1986).

In an attempt to avoid confusion, Avalon terrane in the following summary is used in reference to the late Precambrian volcano-plutonic criteria and the early Paleozoic faunal distinctions and hence corresponds to the Avalon zone of O'Brien et al. (1983), the Avalonian terrain of Rast and Skehan (1983), and the Avalon composite terrane of Keppie et al. (1985).

# AVALON TECTONOSTRATIGRAPHY IN THE NORTHERN APPALACHIANS

Details of Avalon lithostratigraphy and tectonics have been presented by numerous authors, most recently Pique (1981), O'Brien et al. (1983), Rast and Skehan (1983), Skehan (1983), and Keppie et al. (1985). These are summarized and updated for the Northern Appalachians in the following review in order to provide a tectonostratigraphic framework within which more detailed geologic studies of individual segments of the Avalon terrane may be compared.

# Precambrian Components

The earliest known components of Avalon terrane are platformal quartzites and marbles of probable Middle Proterozoic (Helikian) (Hofmann 1974) but uncertain tectonic significance, that are spatially associated with granodioritic and tonalitic gneisses considered to represent an earlier basement (Currie 1983) to the clastic-carbonate cover (Fig. Candidates for basement include the Kelly Mountain complex of Cape Breton (Keppie 1979), the Mount Thom and Bass River (Great Village River gneiss) complexes in the Cobequid Highlands (Donohoe and Wallace 1985), and the Brookville gneiss of southern Brunswick (Wardle 1978; Currie et al. 1981). Potential candidates may also include gneisses of the Grey River complex in southern Newfoundland (Smyth 1981), although more recent work suggests these may be early Paleozoic (S.J. O'Brien, pers. comm. 1986), and recently recognized gneisses in central Cape Breton (S.M. Barr, pers. comm. 1986). Platformal successions include the George River Group of Cape Breton (Milligan 1970), the Gamble Brook and Folly River schists (Bass River. complex) of the Cobequid Highlands (Donohoe and Wallace 1985), and the Green Head Group of southern Brunswick (Wardle 1978). equivalent. medium-grade Precambrian carbonate and clastic rocks are reported (Stewart and Wones 1974; Osberg et al. 1985) from Penobscot Bay, Maine 1), and may also include quartzof the Westboro Formation in southeastern New England (Rast Skehan 1983).

Upper Proterozoic (Hadrynian) components of Avalon (Figs. 2) include submarine basalts and mafic intrusive rocks of ophiolitic affinity (Strong et al. 1978) that are largely restricted to the Burin Group of southern

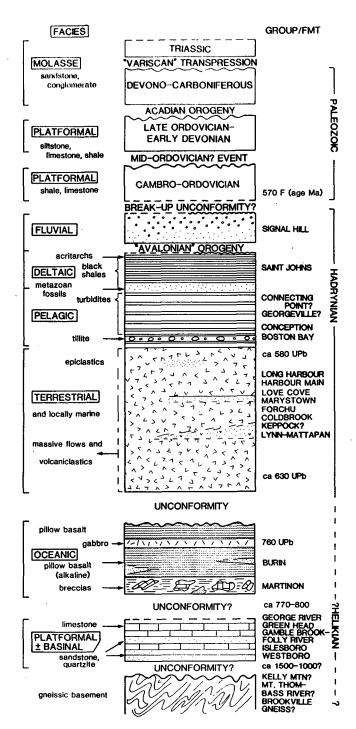


Fig. 2. Schematic composite section of the Avalon terrane in the Northern Appalachians showing facies development, tectonism and nomenclature (modified after O'Brien et al. 1983).

Newfoundland where their age, dated by an intrusive but probably comagmatic gabbro, is at least 762±2 Ma (U/Pb, Krogh et al. 1983). A gabbroic pegmatite in Boston has yielded a K/Ar age of 886±22 Ma (Zartman and Naylor 1984).

These and broadly similar Rb/Sr metamorphic ages obtained by Olszewski et (1981) and Gaudette et al. (1985) from the Kelly Mountain complex (701±66 Ma and 680-700 Ma respectively), by Olszewski and Gaudette (1982) from the Brookville gneiss  $(771\pm 55 \text{ Ma})$ , and by Brookins (1976) from gneisses of the Penobscot Bay region  $(750\pm100 \text{ Ma})$ . suggest that northwestern portions of the Avalon terrane underwent a period of regional metamorphism, deformation and, perhaps, ocean crust formation that is distinct from, and earlier than. the late Hadrynian Avalonian orogeny.

Late Hadrynian subaerial volcanic rocks occur throughout the Avalon terrane and show ages that cluster in the range 580-630 Ma (O'Brien et al. 1983: Rast Skehan and 1983). In Newfoundland. they occur in four distinct belts bound by Acadian faults (O'Brien et al. 1983) and include the calc-alkaline, bimodal basalt-rhyolite suite of the Connaigre Bay and Long Harbour Groups (O'Driscoll and Strong, 1979); the volcanic and volcaniclastic rocks, basalt flows, and rhyolitic pyroclastic rocks of the Marystown and and the 590±30 Ma (U/Pb, Dallmeyer et 1981b) Love Cove Groups: bimodal, spilite-keratophyre succession enigmatic Bull Arm Formation (Hughes and Malpas 1971); and bimodal pyroclastic rocks. lavas and volcaniclastic rocks of the mildly alkalic to tholeiitic Harbour Main Group (Papezik The Marystown Group displays a 1970). calc-alkaline/tholeiitic three-stage, to mafic alkalic to dominantly felsic evolution, the lower part of which may correlate with the Connaigre Bay Group upper part with the Lone (Strong et al. Harbour Group 1984). The Long Harbour Group is capped by peralkaline rhyolites that have been correlated with those of the Love Cove The Harbour Main Group, which is intruded by the 607±11 Ma (Rb/Sr. McCartney et al. 1966) Holyrood granite, is dominated by metaluminous rhyoand alkali basalts but outcrops in three belts of slightly differing ages (607 Ma, 622 Ma and 631 Ma; O'Brien. pers. comm. 1986). O'Brien et al. (1986) have proposed a three-stage, calc-alkaline/mildly tholeiitic to mildly alkaline to peralkaline magmatic evolution for the late Hadrynian volcanic rocks of Newfoundland, comparable to that of the Mesozoic-Cenozoic Basin and Range Province.

Late Hadrynian volcanic rocks in mainland Canada include felsic pyroclastic rocks, basalts, andesites and rhyolites with calc-alkaline, volcanic arc affinities in the Fourchu Group of Cape Breton (Keppie et al. 1979a. 1985); ignimbrites, felsic tuffs, andesites and basalts of the Keppoch Formation in the Antigonish Highlands (Murphy 1984); and felsic tuffs, volcanoclastic rocks and felsic to mafic flows of calc-alkaline chemistry (Rast et al. 1984) in the Coldbrook Group of New Brunswick (Giles and southern Ruitenberg 1977; Currie 1984). Hadrynian tholeiitic and ocean island basalts in the Georgeville Group, which overlies the Keppoch Formation (Murphy 1984), have been interpreted in terms of an interarc basin (Dostal et al. 1984). diabase and bimodal. A rhyolitic-dacitic dike complex (Rast 1979; Currie 1984), which locally intrudes the Coldbrook Group and shows tholeiitic to calc-alkaline affinities 1984), al. been (Rast has et attributed to the transtensional opening of the Iapetus Ocean (Rast 1979) and may be penecontemporaneous with the formation of the adjacent Pocologan mylonite zone in southern New Brunswick (Rast and Currie 1976: Rast and Dickson 1982).

In southeastern New England, felsic pyroclastic rocks, rhyolites, dacites and mafic flows of the bimodal Middlesex Fells and Lynn volcanic complexes are intruded by the co-genetic (Smith et al. 1985), 630±15 Ma Dedham granodiorite (K/Ar, Zartman and Naylor 1984) and may be equivalent to parts of the Coldbrook Group, from which Stukas (1977) obtained several ~630 Ma, Ar/Ar release spectra, and older portions of

the subaerial volcanic succession in Newfoundland (Rast and Skehan 1983). Mafic members of the Middlesex Fells complex are transitional alkali olivine basalts (Durfee Cardoza et al. 1985). The younger,  $602\pm3$  Ma Mattapan volcanics (U/Pb, Kaye and Zartman 1980), which comprise ash flow tuffs, lahars, volcanic breccias and andesites, have been interpreted as the product of collapse (Thompson 1985). caldera Highly deformed and metamorphosed, late Hadrynian rhyolitic and mafic volcanic be represented in may Ellsworth Formation of southern Maine (Osberg et al. 1985; Stewart 1986).

Intrusive into the late Hadrynian volcanic rocks of Newfoundland, Nova Scotia, New Brunswick, New England and perhaps Maine is a suite of probably co-genetic, granitoid plutons with ages in the range 540-620 Ma (e.g. O'Brien et al. 1983: Skehan 1983: Zartman and Naylor 1984: Hermes and Zartman 1985). Compositions are commonly granitic, granodioritic and tonalitic but range to minor diorites and gabbros, and chemistries are typically of calc-alkaline affinity (0'Driscoll and Strong 1979: Clarke et al. 1980: Skehan 1983; Hermes and Zartman 1985). significance of these granitoid rocks remains uncertain, but the presence in some of a strong primary foliation, their absence as intrusions into Cambrian strata, and their chemical similarity to Cenozoic Cordilleran and Andean granitoids suggest that those yielding older ages may be syntectonic with respect to the late Precambrian orogeny (O'Brien et Avalonian 1983). If so, early Cambrian radiometric dates from some of these granitoid rocks may be anomolously young (Poole 1980). Polyphase deformation greenschist to amphibolite facies metamorphism accompanied the Avalonian orogeny in Nova Scotia and Cape Breton (Keppie 1982b), and deformed late Precambrian volcanic clasts occur in basal Cambrian conglomerates in southern New Brunswick (McCutcheon et al. 1982). In Newfoundland, however, effects of the Avalonian event are largely restricted to the development of local unconformities (O'Brien et al. 1983), although cataclastic granitoid clasts occur in proximal fan facies of the late Precambrian Signal Hill Group (S.J. O'Brien, pers. comm. 1986).

the Ava1on Peninsula of In Newfoundland, thick sequences of flyschoid and molasse-like sediments overlie the late Hadrynian volcanic rocks (Fig. 2) and are locally overlain by Cambrian strata (King 1980). Conformably and locally unconformably overlying the Harbour Main Group, the late Hadrynian Conception Group comprises a thick, flysch-like sequence dominated by marine, turbiditic volcaniclastics carry well-preserved Ediacara fauna (Anderson and Misra 1968) contains a distinctive tillite unit (Gaskiers Formation) near its base (Bruckner and Anderson 1971). The group may correlate with the dominantly argillaceous Connecting Point Group to the west (King 1980). The Conception Group is, in turn, conformably overlain by black shales and siltstones of the St. John's Group which are conformably succeeded by molasse-like redbeds of the Signal Hill Group. King (1980) has interpreted this succession as reflecting shoaling of a marine turbidite basin followed by southward delta proand the gradation development alluvial plain conditions. Basal orthoquartzites of the unconformably overlying Cambrian sequence represent the onset of marine shelf conditions.

With the exception of Antigonish Highlands, where thick basinal shales and turbidites associated with intermittent oceanic-type volcanism dominate the Georgeville Group (Murphy 1984), extensive late Hadrynian sedimentation is not represented elsewhere within the Canadian Avalon terrane, nor is it present in the Burin Peninsula of Newfoundland. In Cape Breton and southern Brunswick. late Hadrynian volcanic rocks are respectively succeeded by redbeds of the Morrison River (Weeks 1954) and Ratcliffe Brook (Hayes and Howell 1937) Formations, and then by

Cambrian shales. However, recently identified "Kocambrian" feldspathic redbeds. felsic tuffs and basalts may disconformably separate the Cambro-Hadrynian Coldbrook and Ordovician Saint John Groups in southern New Brunswick (Currie 1984, 1986a; Tanoli et al. 1985). Hadrynian to Ordovician interbedded feldspathic and manganiferous metasedimentary rocks are in coastal southern (Stewart 1986).

In southeastern New England, however, the Boston Basin contains extensive conglomerates, argillites sandstones with occasional interbeds of tuffs (Billings 1976) and andesites of the Brighton volcanics (Durfee Cardoza et al. 1985). This thick succession. which constitutes the Boston Bay Group. nonconformably overlies the late Precambrian Dedham granodiorite, contains Vendian acritarchs and clasts of the Lynn and Mattapan volcanics (Skehan and Murray 1980, Skehan 1983), and, with the exception of the abundance of conglomerates, broadly resembles the Conception Group of Newfoundland. the Conception Group, the Boston Bay Group is overlain by Cambrian strata. contains a possible tillite (Squantum Member of the Roxbury Conglomerate), and has a late Precambrian volcanoplutonic provenance (Skehan 1983; Galli and Bailey 1986). Socci (1985) has suggested a glaciomarine origin for the Roxbury Conglomerate, while Socci and (1985) have interpreted Smith Boston Bay Group in terms of debris flows and turbidites of a submarine fan system, as originally implied by Dott Galli and Bailey (1986) sug-(1961).gested a back-arc setting for Boston Bay Group. Clastics and pillow basalts of the Blackstone Series in Rhode Island (Dreier and Mosher 1981) are of uncertain correlation but constitute an additional late Precambrian sedimentary sequence.

# Cambro-Ordovician Overlap Sequences

Shallow-marine, platformal Cambro-Ordovician overlap sequences with dis-

tinctive Atlantic-realm fauna unconformably to disconformably overlie the late Hadrynian volcano-sedimentary successions of Newfoundland. Nova Scotia. New Brunswick and southeastern (O'Brien et al. 1983: Massachusetts Skehan 1983) and late Precambrian granitoid rocks in Rhode Island and southeastern Maine (Skehan 1983; Stewart 1986). Typically floored by Lower Cambrian (Tommotian) orthoguartzites. these sequences are dominated by shale, limestone, sandstone and oolitic ironstone in Newfoundland (O'Brien and King 1982), and Lower Cambrian to Arenig shale, sandstone and limestone in Cape Breton (Keppie 1982b) and in the Saint John Group of southern New Brunswick (Hayes and Howell 1937; Pickerill and Tanoli 1985). Cambrian to Ordovician conglomerate, slate, quartzite. limestone and oolitic ironstone unconformably overlie the Georgeville in the Antigonish Highlands (Murphy et al. 1982; Murphy 1984; Keppie et al. 1985), whereas metamorphosed. late Proterozoic to Ordovician pelites and sandstones rest on sialic Precambrian rocks in the Coasta1 Volcanic Belt of Maine (Stewart 1986). Further north in Maine, quartzite. arenite and thick Ordovician sulfidic and carbonaceous shales of the Cambro-Ordovician Cookson and Penobscot Formations (Osberg et al. 1985) have been correlated with the Saint John Group (Ludman 1986), although their fauna may not be diagnostic of an Atlantic origin (Stewart 1986). In southeastern New England, Acado-Baltic trilobites are characteristic of the Lower Cambrian limestones and shales of the Weymouth and Hoppin Formations, and the Middle Cambrian shales and siltstones of the Braintree Argillite (Skehan and Murray 1980: Skehan 1983). Lower Cambrian lime mudstones and shales that resemble the Hoppin Formation and rest unconformably on late Precambrian turbidites have recently been identified in southeastern Rhode Island (Webster et al. 1986).

Minor but widespread bimodal volcanic rocks of early to middle Cambrian age accompany the Cambro-Ordovician successions of Newfoundland ((Greenough and Papezik 1985), Nova Scotia (Murphy et al. 1985), New Brunswick (Greenough et al. 1985), Maine (Stewart 1986) and possibly New England (Skehan 1983). Dominated by continental alkalic and less common tholeiitic basalts with compositions consistent with a withinplate setting, this episode of volcanism has been attributed both to mild tensional tectonics associated with the stages Iapetus rifting fina1 of (Greenough and Papezik 1985, 1986), and to faulting within local pull-apart basins formed during the last, transpressional stages of the late Hadrynian Avalonian orogeny (Murphy et al. 1985).

Late Ordovician and Younger Evolution

Although a detailed summary of the late Ordovician and younger evolution of the Avalon terrane is beyond the scope of this paper, the tectonostratigraphy of the Avalon terrane continues with an extensive post-Ordovician record. Late Ordovician and Silurian rocks occur in the eastern Newfoundland mainland Nova Scotia and southern Maine but are absent in other areas of the Avalon terrane. A 4000 m thick sequence of Ordovician-Silurian marine shales is unconformably overlain by Devonian redbeds in the Newfoundland offshore (Durling et al. 1986). the Antigonish Highlands, non-marine redbeds and 1ate Ordovician-early Silurian bimodal volcanic rocks unconformably overlie late Cambrian-early Ordovician black shales and pass upwards through Silurian shallow marine sedimentary rocks containing shellv Atlantic fauna (Watkins and Boucot 1975) into early Devonian nonmarine clastics (e.g. Keppie 1980; 1982b). Similarly, in the Cobequid Highland, bimodal but predominantly felsic volcanic rocks and fossiliferous quartz wackes, siltstones and shales that span Silurian overlie probable late Ordovician siltstones and tuffs, and are followed by early Devonian redbeds

(Donohoe and Wallace 1985). Thick, Lower Silurian to Lower Devonian bimodal volcanic and shallow marine sedimentary rocks bearing shelly Atlantic fauna also occur in southern Maine (Gates 1969; Brookins et al. 1973; Stewart 1986).

Although little concensus exists on the timing of assembly between the Avalon terrane and adjacent terranes and their docking with the margin of the Northern Appalachian orogen (e.g. Williams and Hatcher 1983; Zen 1983; Keppie et al. 1985), pre-Silurian accretion to more westerly terranes has been proposed by some for portions of the Avalon terrane in Canada (Keppie et al. 1985; Currie et al. 1986), Maine (Ludman 1986) and New England (Robinson and Hall 1980). Effects of the mid-tolate Ordovician Taconic orogeny, which are widely manifest in western portions of the Northern Appalachians Stanley and Ratcliffe 1985), are absent within the Avalon terrane. evidence for broadly contemporaneous deformation and metamorphism exists locally. Major unconformities separate early Ordovician from early Silurian in the Antigonish Highlands (Keppie et al. 1985) and southern Maine (Stewart 1986), and the late Precambrian from the early Silurian in the Cobequid Highlands (Donohoe and 1985). In the northern Antigonish Highlands, local polyphase folding and greenschist facies metamorphism occurred between the Arenig and latest Ordovician (Keppie 1982b, Keppie et al. 1985), while late Ordovician (470 Ma, Rb/Sr) pegmatitic sweats in the Brookville gneiss (Olszewski et al. 1980) may indicate locally elevattemperatures in southern Brunswick (Currie et al. 1981; Nance 1982). Evidence for Ordovician deformation and metamorphism may also exist in the form of metamorphosed fragments in the Lower Silurian of southern Maine and Wones 1974), (Stewart presence of pre-Silurian deformation in the Cookson Formation (Ludman 1985) if its correlation with the Saint John Group (Ludman 1986) is upheld and,

indirectly, in the single Ordovician faunal province envisaged for Maine by Neuman (1984).Bimodal. Silurian volcanic rocks in mainland Nova Scotia (Boucot et al. 1974) and similar Silurian-Devonian volcanics in southern Maine (Brookins et al. 1973) show anorogenic, within-plate characteristics (Moench and Gates 1976; Gates 1978; Keppie et al. 1979b) but have been interpreted as reflecting thermal and mechanical readjustments of Avalonian crust following Ordovician deformation (Currie et al., 1986). Ordovician and Silurian peralkaline plutons in southeastern New England (Zartman and Marvin 1971), such as the Quincy granite in the vicinity Boston, are also anorogenic (Zartman and Naylor 1984).

During the early to mid-Devonian Acadian orogeny, the Avalon terrane was affected by widespread regional metamorphism and polyphase deformation (e.g. Bradley 1983; Rast 1983). Metamorphic grades are typically of greenschist facies or lower and peaked at 370-390 Ma in Newfoundland where they are interpreted as accompanying tectonic juxtapositioning of the Avalon and (Dallmeyer et al. Gander terranes 1983: O'Brien et al. 1983). ciated Acadian granitoid plutons consist of locally voluminous suites of post-tectonic granites granodiorites that occur throughout the Avalon and adjacent terranes, although they have yet to be firmly demonstrated southern New Brunswick (Currie 1986b). Ages for these plutons cluster the interval 340-370 Ma (e.g. Brookins 1972; Clarke et al. 1980: Strong 1980; Donohoe and Barr 1981; Zartman and Naylor 1984; Hermes and Zartman 1985).

Following the Acadian orogeny, the tectonostratigraphy of the Avalon terrane is no longer unique, but is sedimentologically and plutonically linked to those of its bounding terranes (Williams and Hatcher 1982, 1983). In Newfoundland, Devonian conglomerates containing clasts derived from the adjacent Gander terrane unconformably

overlie Upper Cambrian shales on the eastern side of the Hermitage Bay Fault (Williams 1971), while the 355 Ma (Rb/Sr, Bell et al. 1977) Ackley granite straddles the same terrane boundary (Fig. 1).

Similar Devonian and Carboniferous molasse-like facies of alluvial. fluvial and locally lacustrine and shallow marine origin occur throughout most of the Avalon terrane, and are accompanied by subordinate bimodal volsequences of tholeiitic canic strongly alkaline affinities Newfoundland (O'Brien and King 1982), western Cape Breton (Blanchard et al. 1984), northern Nova Scotia (Keppie and the Magdalen Islands Dostal 1980). (Barr et al. possibly in 1985), southern New Brunswick (Strong et al. 1979) although the age of these has questioned (Currie and Nance been 1983). southeastern Maine and (Pickerill et al. 1978). In southeastern. New England molasse-like sequences are largely Pennsylvanian in age (Lyons 1979) and devoid of volcanics (Skehan 1983).

Carboniferous redbeds in coastal southern New Brunswick are additionally involved in extensive late Paleozoic (Alleghanian/Variscan) thrusting polyphase deformation (e.g. Rast and Grant 1973: Ruitenberg et al. 1973: Rultenberg and McCutcheon 1982) that, elsewhere in Canada, are largely confined to the Minas Geofracture (Keppie 1982a: Donohoe and Wallace 1985). Massachusetts and Rhode Island. Pennsylvanian sedimentary rocks are also affected by strong, polyphase deformation and metamorphism (Burks et al. 1981; Mosher and Rast 1986) and are intruded by the Permian Narragansett Pier granite (Hermes et al. 1981a, b). In both regions, late Paleozoic deformation appears to have involved an imcomponent of transpression portant (Mosher 1983; Nance and Warner 1986) that was associated with the final tectonic juxtapositioning of the Avalon and Meguma terranes (e.g. Spariosu et al. 1984).

Carboniferous redbeds deposited with-

in Alleghanian transtensional basins (e.g. Bradley 1982; Yeo and Ruixing 1986) are locally overlain unconformably by undeformed fluvial and alluvial sequences of Triassic age (Klein 1962). Triassic deposition occurred response to initial rifting of the present Atlantic Ocean (e.g. Ballard and Uchupi 1975; Nadon and Middleton 1984) and was locally accompanied reactivation extensional of the Alleghanian faults (Plint and Van Poll 1984).

# PRECAMBRIAN TECTONIC MODELS FOR THE AVALON TERRANE

Controlling factors for the late Precambrian evolution of the Avalon terrane remain largely equivocal due to uncertainties in regional correlation and local geologic development, coupled with the diverse geochemical signatures and scattered distribution of its Hadrynian igneous record. Earlier interpretations have consequently ranged from a pre-Iapetus, ensialic island arc (Rast et al. 1976) to an aborted Proterozoic rifting event, perhaps related to the opening of the Iapetus Ocean (Strong et al. 1978). More recent models advocate a genetic distinction between the development of Avalon terrane and that of Iapetus and point to the diversity in volcanic geochemistry as evidence of the composite nature of the Precambrian Avalon terrane, the accretion of which was largely complete prior to the inception of the Iapetus cycle.

Keppie (1982b) and Dostal et al. (1984), for example, have interpreted the late Precambrian development of the Avalon terrane of Nova Scotia in terms of a cratonic island arc (Cape Breton terrane), an interarc basin (Antigonish terrane, Murphy and Keppie 1985), and a remnant arc (Cobequid terrane) associated with the subduction of a major ocean lying to the southeast (Fig. 3). These terranes were assembled during the Avalonian orogeny to produce the Avalon composite terrane, and subsequently displaced by Devonian

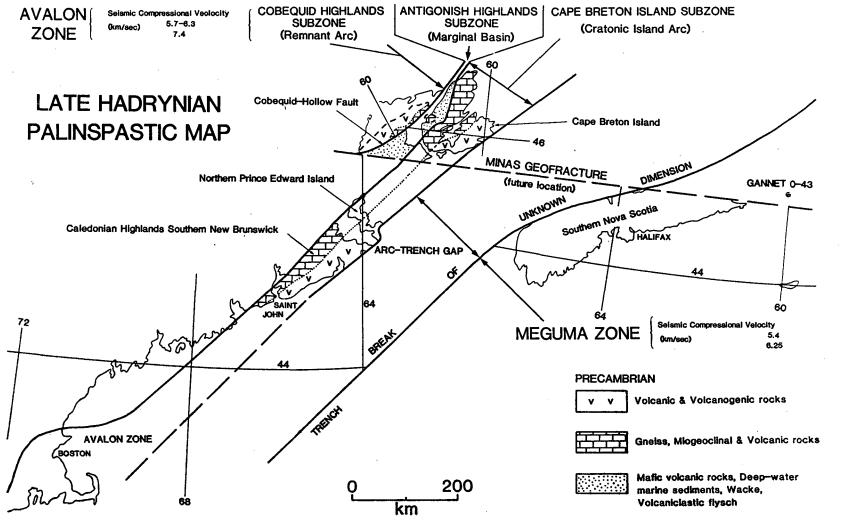


Fig. 3. Late Hadrynian palinspastic map for the Avalon and Meguma terranes in the Northern Appalachians as proposed by Keppie (1982b).

younger strike-slip movements on the Minas Geofracture. However, whether diversity in volcanic geochemistry that essentially reflects the spatial and temporal variation of volcanic regimes within a single arc complex is sufficient evidence to invoke separate suspect terranes is debatable and, with possible exception the Antigonish terrane (Murphy 1984), there is little evidence to suggest that processes of suspect terrane accretion occurred during the Avalonian orogeny.

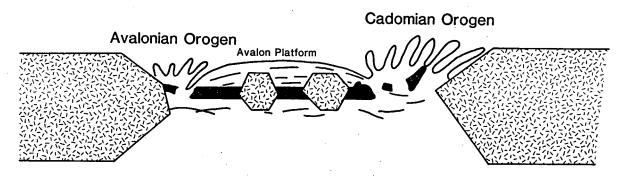
Rast and Skehan (1983) have similarly suggested that the Avalon terrane originated as a microcontinental arc which was separated from North America by a back—arc basin and from West Africa by the late Precambrian

Cadomian Ocean (Fig. 4a). They suggest that the arc developed as a volcanic carapace covering older sialic islands and ensimatic basins derived from fragmentation of an African supercontinent at 800-850 Ma. However, the processes by which this carapace developed are unclear in their model. Closure of this ocean and back-arc basin at 600-650 Ma (Fig. 4b) respectively produced the Cadomian (II) and Avalonian orogenies and accreted Avalon to the North American and West African cratons.

In contrast, O'Brien et al. (1983) employ the model of Kroner (1980) to propose a Pan-African origin for the Avalon terrane in Canada. In this model (Fig. 5), rifting of pre-Pan-African continental basement during

# North American Craton Back Arc Arc Cadomian Ocean The state of the s

# a. 700-800 Ma



b. 600 Ma

Fig. 4. Schematic cross sections from the North American (Laurentian) to the West African (Gondwanan) cratons proposed by Rast and Skehan (1983). (a) Section through Cadomian basins and microcontinent (ca. 700-800 Ma). (b) Section through Avalonian orogen, Avalon Platform and Cadomian orogen (ca. 600-650 Ma). No vertical scale implied.

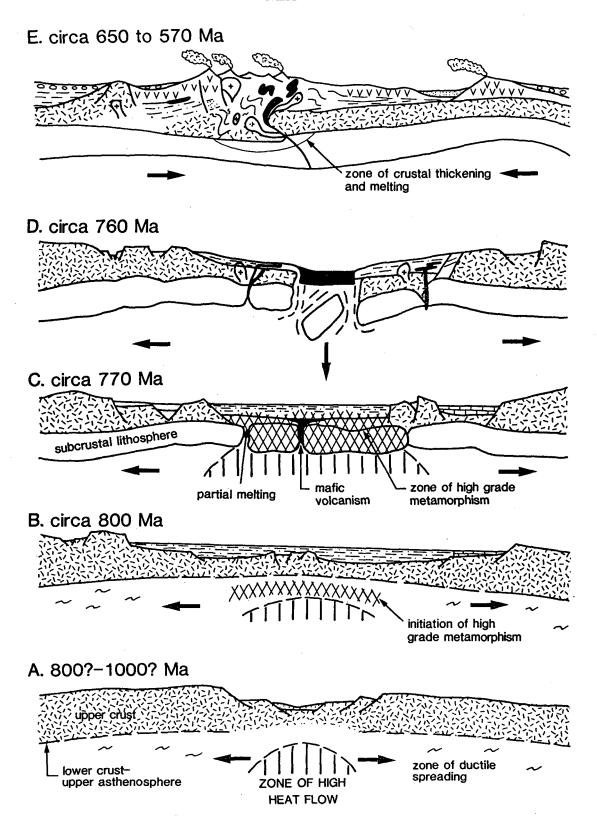


Fig. 5. Schematic cross sections illustrating major stages in the Precambrian development of the Avalon terrane in Canada as proposed by O'Brien  $et\ al.$  (1983).

interval 800-1000 Ma initially the produced ensialic basins (c.f. George River and Green Head Groups) that subsequently underwent partial melting and high-grade metamorphism during the period 750-800 Ma due to ductile spreading of the lower crust and local crustal separation (c.f. Burin Group). closure of these basins is attributed to limited subduction and collision during the interval 570-650 Ma, and produced late Hadrynian volcanic assem-Marystown Group blages (c.f. equivalents) and Avalonian orogenesis terminated with the intrusion that granitoids and marine-to-terrestrial clastic deposition (c.f. Concep-Johns/Signal Hill Groups). tion/St. The diversity of deformational intensity and volcanic geochemistry is attributed to the degree of crustal separation and ocean crust formation and, hence, subsequent variations in the style and amount of subduction and collision.

However, while there is general concensus on the existence of a late Hadrynian magmatic arc complex within the Avalon terrane, the nature of the early Hadrynian event is less certain. presence of similar carbonate slump-breccias in both the Martinon Formation (Wardle 1978), which may unconformably overlie the Green Head Group (Leavitt 1963; Currie 1984), and the Burin Group (Strong et al. 1978) suggests the early Hadrynian tectonothermal event coincided with collapse of the Green Head carbonate platform (Currie 1986b). The age of the Green Head Group is uncertain, but the Helikian age assigned to it on the basis of the stromatolite Archaeozoon acadiensis (Hofmann 1974) would suggest that, by the early Hadrynian, the platform was already well established and. hence, unrelated to the early Hadrynian rifting event proposed by O'Brien et al. (1983). Given the limited record of this event, a case could then be made for the collapse and regional metamorphism of a pre-existing platform during an early Hadrynian subduction event, perhaps coinciding with older

Rb/Sr ages (e.g. 750±80 Ma, Cormier 1969; 830±55 Ma, Stukas 1977) obtained from volcanic rocks of the Coldbrook Group. If so, the Burin Group, which rests on clastic rocks derived from a subaerial felsic volcanic terrane (Strong et al. 1978), could be reinterpreted as the product of a back—arc basin.

Furthermore, although the age of the basement of Avalon terrane remains controversial (e.g. Zen 1983), model age constraints on the protolith of the gneiss (Olszewski Brookville Gaudette 1982) are not inconsistent with the development of the Green Head Grenvillian basement, platform on tempting correlation of the Green Head Group with Helikian platform carbonates of the Grenville belt (e.g. Baer 1967). Considerations such as these have led Currie (1986b) to propose a tectonostratigraphic history for the Avalon terrane involving the repeated break-up and rewelding of a North American continental edge.

Evidence for a late Hadrynian subduction event is more compelling and forms a common theme to recent models the late Precambrian evolution of Avalon terrane. However, with respect to the present distribution of Avalon terrane within the Northern Appalachians, wide lateral and/or temporal variations must have existed within and, more importantly, behind the late Hadrynian magmatic arc complex. Newfoundland (O'Brien et al. 1986), at least three episodes of volcanic activity that span an interval of 50 Ma are locally separated by periods of sedimentation and follow an evolutionpath similar to that of the Basin Range Province. and Early alkaline granitoids and calc-alkaline to mildy tholeiitic basaltic and rhyolitic rocks are followed by epiclastic volcanism, rhyolites and mildly alkaline flood basalts and, finally, peralkaline, bimodal volcanism and rare plutonism associated with terrestrial sedimentation in strike-slip basins. Development of a magmatic arc in Newfoundland would therefore appear to

have been followed by continental backarc transtension that, in turn, may have heralded the inception of the lapetus ocean and the re-establishment of platformal conditions recorded in overlying Cambro-Ordovician succession.

A broadly similar pattern may also be recorded in the Boston Platform where calc-alkaline granitoid and cogenetic pyroclastic rocks (e.g. Dedham Granodiorite-Lynn Volcanic Volcanic North Smith et al. 1985) and Complex. volcanism (c.f. younger ash-flow Mattapan Volcanic Complex, Thompson 1985) are succeeded by intermediate to mafic flows, tuffs and agglomerates (c.f. Brighton Volcanics. Durfee-Cardoza et al. 1985) that are associated with the development of the Boston Basin. Acritarchs, facies associations and broadly northward paleoflow within the overlying Boston Bay Group suggest a marine, back-arc extensetting sional during the 1atest Hadrynian (A.D. Socci, pers. comm. 1986) that again yielded to platformal conditions recorded in the Early Cambrian Weymouth Formation.

In Nova Scotia. southern New Brunswick and coastal Maine the evolution of the magmatic arc remains uncertain as the precise age and correlation late Precambrian volcanic rocks is unknown. Nevertheless, significant variations in the processes operating in the back-arc would appear to have occurred in mainland Nova Scotia and New Brunswick. In the northern Antigonish Highlands (Murphy et al. 1980, Murphy 1984; Murphy and Keppie 1985), for example, basal ocean-floor tholeiites and thick shales and turbidites of the late Hadrynian Georgeville that are penecontemporaneous with continentally floored rhyolites, basaltic-andesites and basalts of the Keppoch Formation to the south. interpreted as forming in a narrow. deep-water basin floored by oceanic Thus back-arc processes within crust. the Antigonish Highlands involved the generation of ocean crust that has been attributed (Keppie 1982b, Murphy and Keppie 1985; Fig. 3) to the development of an interarc basin between a cratonic island arc. recorded in volcanics of the Fourchu Group of Cape Breton (Keppie et al. 1979a), and a remnant arc represented by presumed volcanics in the Highlands (e.g. Donohoe and Wallace 1985). Furthermore, alkalic basaltic volcanism that occurs first in the Keppoch Formation and later in the Georgeville Group has been related to the influence of "hot-spot" magmatism on a southerly moving, interarc lithosphere. Alaskitic and appinitic plutonism and deformation within the Georgeville Group during the latest Hadrynian are taken to reflect closure of the interarc basin and are followed by Cambrian rocks interpreted as a riftrelated succession (Murphy 1984).

In southern New Brunswick, contrast, back-arc processes appear to have involved significant strike-slip movements and the penecontemporaneous development of a 2 km-wide mylonite zone (Rast and Dickson 1982) and a mafic to bimodal dike complex (Rast 1979; Currie 1984, 1986a). Mafic and felsic flows, tuffs, ignimbrites and intercalated sedimentary rocks of the Coldbrook Group are generally considered equivalent to the Fourchu Group of Cape Breton (e.g. Keppie et al. 1985) although they are likely to exceed 10 Km in thickness (Giles and Ruitenberg 1977) and their stratigraphy, age range and geochemical affinities are poorly known. Nevertheless, a magmatic arc setting for the Coldbrook Group, and late Hadrynian and probably cogenetic granitoid plutons with which it is associated, seems likely. However, parallel and adjacent to the Belleisle fault, which here forms the northwestern boundary of the Avalon terrane. Coldbrook lithologies are affected by temporally and spatially associated zones of intense mylonitization and dike injection that are absent in the Cambro-Ordovician Saint John Group (Rast and Dickson 1982). complex, which extends laterally for consists of predominantly northeast-trending mafic dikes at its

narrower southwestern end (Rast 1979) but forms a 5 km-wide zone of alternating, north-south mafic and rhyolitic dikes to the northeast (Currie 1984. 1986a). Mylonitization affects earlier but not later dikes and has been attributed to left-lateral shear (Rast and Dickson 1982) compatible with the obliquity of north-south dike orientations to the regional trend of the In southern New Brunswick. complex. therefore. onset Cambrothe of Ordovician platformal conditions, which were heralded within the magmatic arc by minor, latest Hadrynian bimodal volcanics and redbeds that may unconformably overlie the Coldbrook Group (Currie 1986a), was preceded by focused back-arc transtension.

of the 1ate The evolution Hadrynian arc complexes thus followed quite separate paths within each of the areas in which the Avalon terrane is exposed. However, since the relative separation of these areas in the late Hadrynian is unknown, the variations in their evolution are not inconsistent with their development within a single subduction system as illustrated by the present complexity of the Indonesian region (Hamilton 1979). Interestingly, their is no evidence within the Avalon terrane that subduction was terminated through a major collisional event. fact, deformation and metamorphism associated with the Avalonian orogeny are both mild and localized, being most intense in the Antigonish Highlands where their development has been linked closure of the interarc (Murphy and Keppie 1985). Elsewhere. however, subduction was followed by back-arc rifting and shear suggesting that transform termination of oblique subduction, perhaps as a result of ridge-trench collision, was followed by back-arc relaxation and extension. Furthermore, in each of these areas, some form of back-arc extension heralded the onset of Cambro-Ordovician. Atlanticrealm platformal conditions that are widely attributed to the Lower Paleozoic location of the Avalon terrane near the southern margin of the Iapetus Ocean (e.g. Cooks and Fortey 1982; Neuman 1984). It would therefore appear that the locus of Iapetus rifting, at least on its southern margin, was initially provided by latest Hadrynian extension within the back—arc region of the late Precambrian Avalon terrane.

# SUMMARY AND DISCUSSION

The Avalon terrane of the Northern Appalachians, which can be defined by presence of 1ate Precambrian the volcano-sedimentary sequences and granplutons and Lower Paleozoic itoid Atlantic-realm platformal successions, forms a distinctive tectonostratigraphic belt that lies inboard of the Meguma terrane along much of the orogen's southeastern margin. The oldest known elements of the terrane are Helikian carbonates and mature clastic rocks that rest on gneissic continental basement and locally are overlain unconformably by slide breccias and turbidites interpreted to record collapse of the ensialic platform during an early Hadryian (ca. 750-800 Ma) tectonotherm-Associated high-grade real event. gional metamorphism and local oceanfloor tholeiitic magmatism affect the northwestern margin of the terrane and have been attributed to an abortive However, the associarifting event. tion is not inconsistent with an episode of subduction and back-arc spreading.

Voluminous late Hadrynian volcanic and associated granitoid rocks are the products of a magmatic arc system that appears to have involved a variety of back-arc processes including terrestrial and marine continental back-arc extension and transtension, and oceanic back-arc basin formation. Transform termination of oblique subduction coupled with local back-arc shortening is proposed for the 1ate Hadrynian Avalonian orogeny. Latest Hadrynian bimodal volcanic rocks may herald the inception, within the former back-arc, of the Lower Paleozoic Iapetus Ocean recorded in Cambro-Ordovician platformal successions and local within-plate

volcanism. Together, these form an overlap sequence, following a minor (break-up?) unconformity, to the Precambrian development of the Avalon terrane.

The Paleozoic evolution of Avalon terrane locally continues with latest Ordovician to early Devonian bimodal volcanism and shallow marine sedimentation that follows fragmentary evidence for a mid-Ordovician tectonothermal event and is followed by locally intense Acadian deformation, metamorphism and widespread plutonism. However, the significance of these events with resto the Paleozoic accretionary history of the Avalon terrane remains Widespread Devono-Carboniuncertain. ferous molasse and predominantly alkaline bimodal volcanic rocks form an overlap sequence with adjacent terranes and are locally affected by late Paleodeformation and metamorphism. Triassic redbeds and mafic volcanic rocks reflect the initial rifting of the present Atlantic Ocean.

Current tectonostratigraphic models for this complex evolutionary record have been strongly influenced by the recent proposition that much of the Northern Appalachians outboard of the North American miogeocline, rather than comprising a longitudinal system of zones, represents a mosaic of terranes with quite separate geologic histories. As a result. conventional tectonic models based on the overall bilateral symmetry of the belt that is implicit in the zonal scheme, although not without considerable merit, have proven to be over simplistic. Paradoxically, the broad regional correlations encouraged by zonal subdivisions permitted the construction of coherent tectonic models despite our rudimentary knowledge of their detailed geology. In contrast, successful terrane analysis requires a level of geologic sophistication beyond that currently available and, hence, the introduction of the terrane concept has temporarily compromised our understanding of the evolutionary history of the orogen. interpretation of Avalon (and other terranes within the Appalachian orogen) is therefore at an exciting stage of disarray. In particular, our knowledge of the Precambrian tectonic, magmatic and stratigraphic evolution of Avalon terrane, and the position and accretionary kinematics of Precambrian terrane boundaries that may exist within it, remain poorly understood, while the Paleozoic coherency and accretionary history of the terrane are poorly constrained. As a result, opinions diverge what is and is "Avalonian" this although it is author's view that the basic bipartite definition should hold.

Of the numerous additions to the geological data base needed to resolve these major uncertainties, precision radiometric ages and high-quality paleomagnetic data are essential for posterrane identification within sible Avalon and the determination of their displacement histories; further igneous and isotopic geochemistry is needed for evidence of tectonic setting, magmatic evolution and basement signatures; extensive metamorphic and deformational analysis is required to provide detailed thermotectonic models; coupled sedimentological and structural studies necessary to determine terrane tectonostratigraphy and the nature and location of source areas; refined palis required to establish eontology internal stratigraphies, the age of overlap sequences and evidence of faunal provinces; and both shallow and deep seismic reflection data would contribute valuable information concerning seismic stratigraphy. crustal structure, the nature of terrane boundaries. and the style of accretion. Much of this information is in the process of being collected since without doubt, the most valuable contribution of the terrane concept to date has been the revitalization of multidisciplinary bedrock mapping that is both rapidly expanding our knowledge of Avalon terrane and providing the data base needed for its eventual interpretation.

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