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H. V. Donohoe, J. Duncan Keppie et A. N. Halliday

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Résumé de l'article

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Le pluton de Hart Lake-Byers Lake a produit unftgede348±5Maetun rapport isotopique ^Sr/^Sr Initial de 0.7046 + 0.008. Cet age, siillalre a ceux des volcanltes felslques adjacentes de la formation de Byers Brook, suggere qu'lis sent comagnatlques, slgne evident qu'un episode igne carbon I fere forme I'm des falts lajeurs de I'hlstolre geologique des eonts Cobequld. Le rapport 87Sr/86Sr initial et la teneur en blot Ite et hornbIende de ces deux granites suggerent une genese crustale profonde du lapa avec une possible contribution nantelllque.

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Two Rb-Sr Whole Rock Isochrons from Plutons In the Cobequid Highlands, Nova Scotia, Canada

H.V. Donohoe*, A.N. Halilday**, J. Duncan Keppie***

* ***Department of Mines and Energy, P.O. Box 1087, Halifax

Nova Scotia, Canada B3J 2X1

**Scottish Universities Research and Reactor Centre East Kilbride, Glasgow, Scotland, Great Britain, G75 OQU

Two Rb-Sr whole rock isochrons have been obtained from plutons in the Cobequid Highlands of Nova Scotia: the Debert River and Hart Lake-Byers Lake Plutons. The undeformed, unfoliated Debert River Pluton yielded an age of 596 ± 70 Ma and an initial 87Sr/86Sr ratio of 0.7059 ± 0.0007 . This age places an upper age limit on the deformation recorded in the Folly River Schist (youngest unit of the Bass River Complex), the dioritic Frog Lake Pluton, and the sedimentary rocks immediately north and east of the Debert River Pluton.

The Hart Lake-Byers Lake Pluton yielded an age of 348 ± 5 Ma and an initial 87 Sr/ 86 Sr ratio of 0.7046 \pm 0.0008. The similarity of this age to that of the adjacent felsic volcanics of the Byers Brook Formation suggests that they are comagmatic and confirms that a Carboniferous igneous event forms an important part of the geological history of the Cobequid Highlands. The initial 87 Sr/ 86 Sr ratio and the biotite-hornblende content of the two granites suggest a mafic igneous lower crustal source or possible mantle component for the magma.

Les plutons de Debert River et Hart Lake-Byers Lake dans les monts Cobequid en Nouvelle-Ecosse ont ilvré deux isochrones Rb-Sr de roche globale. Ni déformé ni folié, le pluton de Debert River a donné un âge de 596 ± 70 Ma et un rapport 87 Sr/86 Sr initial de 0.7059 ± 0.0007 . Ceci place une limite supérieure sur l'âge de la déformation enregistrée dans le schiste de Folly River (l'unité la plus jeune du complexe de Bass River), le pluton dioritique de Frog Lake et dans les roches sédimentaires louxtant le pluton de Debert River au nord et à l'est.

Le piuton de Hart Lake-Byers Lake a produit un âge de 348 \pm 5 Ma et un rapport isotopique 87 Sr/ 86 Sr initial de 0.7046 \pm 0.008. Cet âge, similaire à ceux des volcanites felsiques adjacentes de la formation de Byers Brook, suggère qu'ils sont comagmatiques, signe évident qu'un épisode igné carbonifère forme l'un des faits majeurs de l'histoire géologique des monts Cobequid. Le rapport 87 Sr/ 86 Sr initial et la teneur en biotite et hornblende de ces deux granites suggèrent une genèse crustale profonde du magma avec une possible contribution mantellique.

INTRODUCTION

In this study, two plutons in the Cobequid Highlands, the Debert River Pluton and the Hart Lake-Byers Lake Pluton were chosen for Rb-Sr isotopic analysis (Fig. 1). These plutons and the geology of the Cobequid Highlands have been described by Donohoe and

Wallace (1980, 1982, 1985). dates allow limits to be set upon sedimentary of various ages voicanic units and deformational provide better also events. They definition of the ages of plutonic events and the origin of the magma.

TECHNIQUES

Rb-Sr isotopic analysis were performed at the Scottish Universities

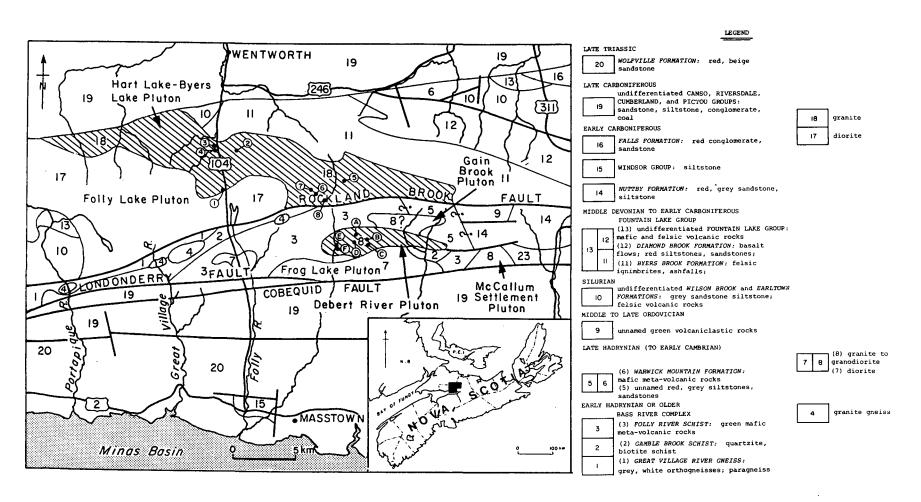


Figure 1. Geological map of the Cobequid Highlands (modified from Donohoe and Wallace, 1980) with sample locations.

Research and Reactor Centre using standard techniques described in detail 1979, elsewhere (Halliday <u>et at.</u>, 1983). Sample powders spiked with ⁸⁷Rb and 84Sr enriched isotopic tracers were dissolved using HF, HNO3 and HC1 acids. Rb and Sr were separated using convencation exchange tional resins. isotopic analyses were performed on a fully automated V.G. Isomass 54E mass spectrometer. Ratios of 87sr/86sr are reported normalized to 88sr/86sr = 87_{Sr/}86_{Sr} The average 8.37521. for NBS987 on this machine was 0.71027±1 N=79) at the time (2₀ mean. of analysis. Regression followed the method of York (1969). The decay constant for ⁸⁷Rb used is $1.42 \times 10^{-11} \text{y}^{-1}$. 87_{Rb/}86_{Sr} uncertainty in estimated to be $\pm 1.0\%$ (2 σ).

DEBERT RIVER PLUTON

The Debert River Pluton is located in the Bass Terrane (Keppie, 1982) between the Cobequid and Rockland Brook Faults (Fig. 1). On its west side, it is intrusive into the Folly River Schist which forms the upper part of the Bass River Complex. A minimum age for the Bass River Complex of 642 \pm 15 is provided by preliminary Rb-Sr data from granite gneisses intruding the Bass River Complex (Gaudette et al., 1983, 1984). On its east side, the pluton intruded and contact metamorphosed sedimentary rocks which have been correlated lithostratigraphically with the Devono-Carboniferous Nuttby Formation by Donohoe and Wallace (1980, To the south, it intruded the undated Frog Lake dioritic pluton. The undated Gain Brook Pluton intruded the Debert River Pluton along its northern margin.

The Debert River Pluton is a fresh, light grey to pale tan grey, medium to coarse grained, hypidiomorphic granular monzogranite to granodiorite. It consists of oligoclase (40-55%), quartz (20-30%), potassium-feldspar (5-30%), accessory biotite and hornblende, and traces of sphene, zircon, epidote and opaques.

The pluton is not foliated although local shear zones are present. pluton contains foliated xenoliths of green, metavolcanic rocks and quartzbiotite schist derived from the Folly River Schist and the Gamble Brook Schist respectively, both units of the Precambrian Bass River Complex. also contains xenoliths of other plutonic rocks such as biotite monzogranite, biotite granodiorite and diorite.

Seven whole rock samples of the Debert River Pluton were analyzed for Rb-Sr and the results are given in Table 1 and plotted on a conventional isochron diagram in Figure 2. The data fit scatter about а best line corresponding to an age of 596 Ma. The "scatter error" (York, 1969) for this determination is \pm 70 Ma (2 σ). The 87Sr/86Sr so determined is initial $0.7059 \pm 7 (2\sigma) (MSWD = 82).$

age of 596 \pm 70 Ma The consistent with all the geological constraints described above except for intrusive relationship sedimentary rocks previously correlated with the Devono-CarbonIferous Nuttby Formation (Donohoe and Wallace, 1980; These new data indicate that 1982). these sedimentary rocks are correlatives of the Nuttby Formation. They may be correlatives of the (?)Late Precambrian Warwick Mountain Formation and/or Jeffers Formation, which implies that a discontinuity is present between rocks and the dated Formation (Fig. 1). However, the age error limits also allow the possibility that the intruded sedimentary rocks could be Cambrian in age.

The age of the Debert River Pluton agrees with the 575 \pm 22 Ma age of the Settlement granite McCallum pluton (Gaudette et al., 1983, 1984) which is located just to the east (Fig. 1). The intrusive contact of the undeformed Debert River Pluton with the dioritic Frog Lake Pluton Indicates an older episode of dioritic plutonism. Elsewhere in the Cobequid Highlands this episode is represented by the Jeffers Brook Diorite (K-Ar on blotite:

				Aba Babasa	Distant Distant
Table 1.	Anaiyticai	and statistical	data for	the Debert	River Pluton.

Letter on Fig. 1	Sample Number	Rb (ppm)	Sr (ppm)	Rb/Sr (Welght)	87 _{Sr/} 86 _{Sr} (atomic)	$87_{Sr}/86_{Sr} \pm 2\sigma^{M*}$ (atomic)
A	JS75-46	91.00	190.4	0.4779	1.384	0.71733 ± 3
В	E11-4569	46.09	321.4	0.1434	0.4150	0.70931 ± 3
С	E11-4570	66.12	294.2	0.2247	0.6504	0.71195 ± 6
D	E11-4571	52.94	190.4	0.2781	0.8050	0.71309 ± 3
E	E11-4573	75.42	248.2	0.3039	0.8796	0.71326 ± 3
F	E11-4574	85.37	238.1	0.3586	1.038	0.71454 ± 4

*20 $^{\rm M}$ = two standard errors of the mean Age \pm 20 a priori (scatter error) Ma = 596 \pm 7 (70) Ma (87 Sr/ 86 Sr) initial \pm 20 a priori (scatter error) = 0.70593 \pm 9 (73) SUMS = 326 (SUMS = sum of the squares of the residuals) MSWD = 81.6 (MSWD = mean squared weighted deviates)

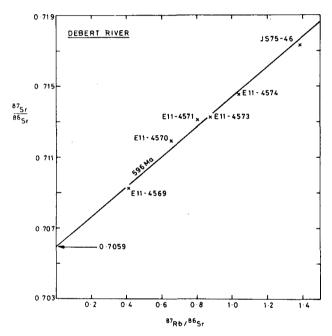


Figure 2. Rb-Sr whole rock isochron for the Debert River pluton.

 544 ± 22 Ma, 564 ± 22 Ma, 585 ± 23 Ma; and K-Ar on hornblende, 616 ± 28 Ma and 628 ± 28 Ma; Wanless et al., 1973, recalculated by Kepple and Smith, 1978, using the decay constants proposed by Stelger and Jager, 1977; and a Rb-Sr biotite isochron, 541 ± 25 Ma, Cormier, 1979, using $1.42 \times 10^{-11} \text{y}^{-1}$ for the 8^{7} Rb decay constant).

The initial 87Sr/86Sr ratio of 0.7059 \pm 0.0007 indicates that the

magma had a crustal component. The mafic mineralogy suggests that the Debert River Pluton has an I-type nature and therefore that the may have been mafic and igneous. mantle-derived component could also involved but its significance cannot be without knowing the exact nature of the lower crust and a complete petrologic study of pluton.

HART LAKE-BYERS LAKE PLUTON

This granitoid pluton occurs in the Cobequid Terrane (Keppie, 1982) lying north of the Rockland Brook Fault (Fig. 1). It has intruded and contact metamorphosed the fossiliferous Silurian Wilson Brook Formation and the Devono-Carboniferous Byers Brook Formation (Donohoe and Wallace, 1982) along its northeastern contact. Felsic volcanic rocks of the Byers Brook Formation have recently yielded a whole rock Rb-Sr isochron age of 341 ± 4 Ma (Cormier, 1982). The Hart Lake-Byers Lake Pluton is also intrusive into the Folly Lake Pluton to the dioritic southwest. It is cut and mylonized by lateral movement on the Rockland Brook Fault to the south and unconformably overlain by Late Carboniferous rocks to the north.

The Hart Lake-Byers Lake Pluton is

a pale pink, hypidiomorphic granular, occasionally porphyritic, hornblende biotite syenogranite to alkali feldspar granite. It consists of potassiumfeldspar (35-50%). quartz (20-35%). albite-oligoclase (10-15%), hornblende of and traces biotite. (5-10%)riebeckite, fluorite, sphene, zircon, allanite and opaques. apatite. internal foliation was observed.

Rb-Sr isotopic analyses of eight whole rock samples from the Hart Lake-Byers Lake pluton are given in Table 2 and plotted in Figure 3. All eight

samples define a reasonable linear array but sample JS76-57 plots slightly further off than the others. A best fit line for all eight samples gives an age of 355 ± 18 Ma (MSWD = 91) and exclusion of sample JS76-57 yields a better fit corresponding to an age of 348 ± 5 Ma with an initial 87Sr/86Sr of 0.7046 ± 8 (MSWD = 5.8). We accept the latter as the best estimate of the age of the pluton.

The Hart Lake and Byers Lake Plutons were previously dated by Cormier (1979), using Rb-Sr techniques.

Table 2. Analytical and statistical data for the Hart Lake-Byers Lake Pluton

	nber ig. 1)	Sample Number	Rb (ppm)	Sr (ppm)		87 _{Rb/} 86 _{Sr} 8 (atomic)	7 _{Sr/} 86 _{Sr} ± 2σ ^M * (atomic)
	1	JS75-22	308.7	129.8	2.378	6.900	0.73831 ± 30
	2	JS75-23	286.4	19.78	14.48	42.75	0.91707 ± 3
	3	JS75-32	161.8	22.71	7.122	20.80	0.80524 ± 6
	4	JS75-33	207.9	113.0	1.840	5.335	0.73117 ± 3
•	5	JS75-37	191.8	9.811	19.55	58.15	0.99391 ± 16
	6	JS75-57	225.5	35.85	6.291	18.38	0.80684 ± 3
	7	JS75-58	166.1	14.33	11.59	34.08	0.87452 ± 3
	8	J\$75-59	234.6	23.36	10.04	29.46	0.85036 ± 4
	2+	1075 03	202	20.1	14.58	43.0	0.0107
	2.	JS75-23	293	20.1	14.55	43.0	0.9127
	5 ⁺	JS75-37	195	12.6	15.48	45.6	0.9179

*20^M = two standard errors of the mean

*Data from Cormier (1979) for comparison (these data are not used in age calculations in this paper)

Sultes	N ¹	Age $\pm 2\sigma$ a.p. (s.e) $(Ma)^2$	$(87_{Sr}/86_{Sr})_{ }$ ± 20 a.p.(s.e) ²	sums ³	MSWD4
Hart Lake Byers Lake	8	355 ± 2 (± 18)	0.70419 ± 32 (± 306)	546	91
ditto minus 76 - 57	7	348 ± 2 (± 5)	0.70457 ± 30 (± 76)	29	5.8

^{1.} N = number of samples used in calculations of age and intial $87_{ST}/86_{ST}$

^{2.} \pm 20 a.p. (s.e.) = \pm two standard errors a priori (scatter error)

^{3.} SUMS = sum of the squares of the residuals

^{4.} MSWD = mean squared weighted deviates

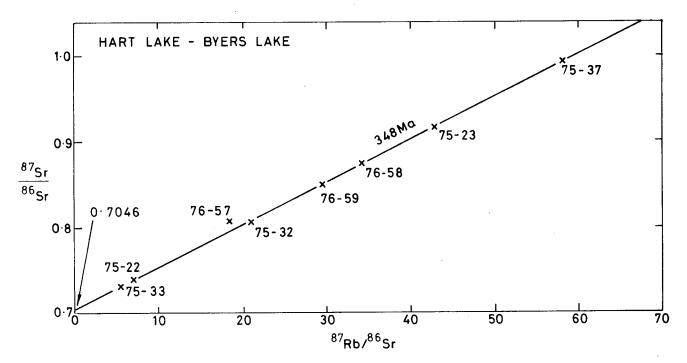


Figure 3. Rb-Sr whole rock isochron for the Hart Lake-Byers Lake pluton.

and yielded results of 331 \pm 17 Ma (initial 87Sr/86Sr ratio = 0.7076 ± 0.0096) and 331 \pm 27 Ma (initial ⁸⁷sr/⁸⁶sr ratio $= 0.7097 \pm$ 0.0093),respectively. The errors were sufficiently large that was impossible to determine whether the plutonism correlates with the widespread Late Devonian (375-360 Ma) granitoid plutonism in the Meguma Terrane of Nova Scotia (Clarke and Halliday, 1980), or represents a later Carboniferous thermal event. The result from the present study (348 \pm 5 Ma) suggests that the plutonism is part of a Carboniferous event.

suite for sample this study includes two samples previously analyzed by Cormier (1979)whose included in Table results are 2 for Discrepancy between the comparison. two laboratories for sample JS75-23 is only 0.6% but is larger for the other sample (JS75-37) (8% for 87Sr/86Sr, and $87_{Rb}/86_{Sr}$). 25% for Analysis of a standard (NBS987) along with these samples at the Scottish Universities Research and Reactor Centre indicates an analytical uncertainty of \pm 0.1% (σ) 87_{Rb/}86_{Sr} in this for laboratory. Cormier (1979) does not report analyses of a standard run along with samples so it is impossible to assess source of the large discrepancies. Two possibilities are inaccurate analyses by Cormier (1979)or inhomogeneity in sample JS75-37. The fact that the ages and initial ratios from the two laboratories agrees within the reported errors may suggest than the latter possibility more likely.

The 348 \pm 5 Ma age of the Hart Lake-Byers Lake Pluton is similar to the age of 339 \pm 4 Ma obtained from granitoid rocks in drill core of the Hart Lake-Byers Lake north Pluton (Cormier, written communication, Furthermore, the similar ages 1984). of the Hart Lake-Byers Lake Pluton and the juxtaposed felsic volcanics of the Byers Brook Formation (341 \pm 4 Ma, Cormier, 1982; and 343 \pm 5 Ma, Cormier, written communication, 1984) suggests that they are comagmatic.

The low initial ⁸⁷Sr/⁸⁶Sr ratio of 0.7047 ¹ 0.0010 suggests a possible lower crustal or mantle source for the magma. More detailed petrologic studies of the plutonic and volcanic rocks are necessary in order to interpret their petrogenesis.

CONCLUSIONS

The 596 \pm 70 Ma age of the Debert Indicates that River Pluton Late Precambrian-Cambrian plutonism was more widespread in the Cobequid Highlands than formerly known (Donohoe Wallace, 1980: 1982). It places an limit on the age of the rocks It intrudes and on the age of deformation observed in the Bass River Complex. contrast, the 348 \pm 5 Ma age of Hart Lake-Byers Lake Pluton clearly documents the existence of plutonism Carboniferous the Cobequid Highlands which corroborates the 339 \pm 22 Ma age for the Cape pluton the western Chignecto in Cobequid Highlands (Cormier, 1979). Carboniferous plutonism was synchronous with adiacent felsic volcanism suggesting that they are comagmatic.

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